

# The IITM Earth System Model (ESM)

## Development and Future Roadmap

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Centre for Climate Change Research (CCCR)  
Indian Institute of Tropical Meteorology, Pune

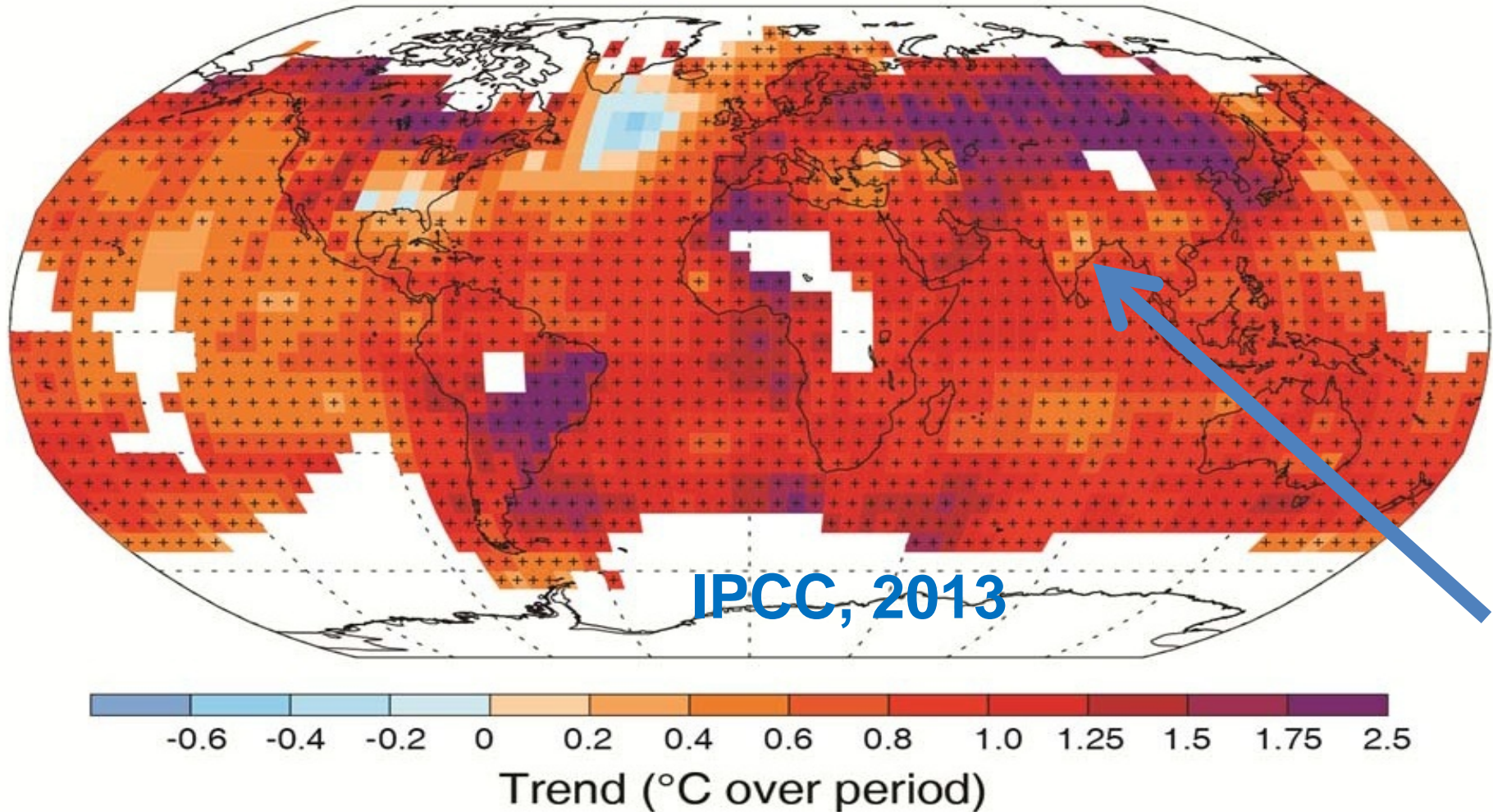
ESM Team: P. Swapna, D.C. Ayantika, Prajeesh, Sandeep Narayansetti, Manmeet Singh, M.K. Roxy, A. Modi, Ramesh Vellore

Diagnostics: M. Mujumdar, B. Preethi, Sabade

**INTROSPECT 2017:** *International Workshop on Representation of Physical Processes in Weather and Climate Model*  
**13 – 16 February 2017, IITM, Pune**

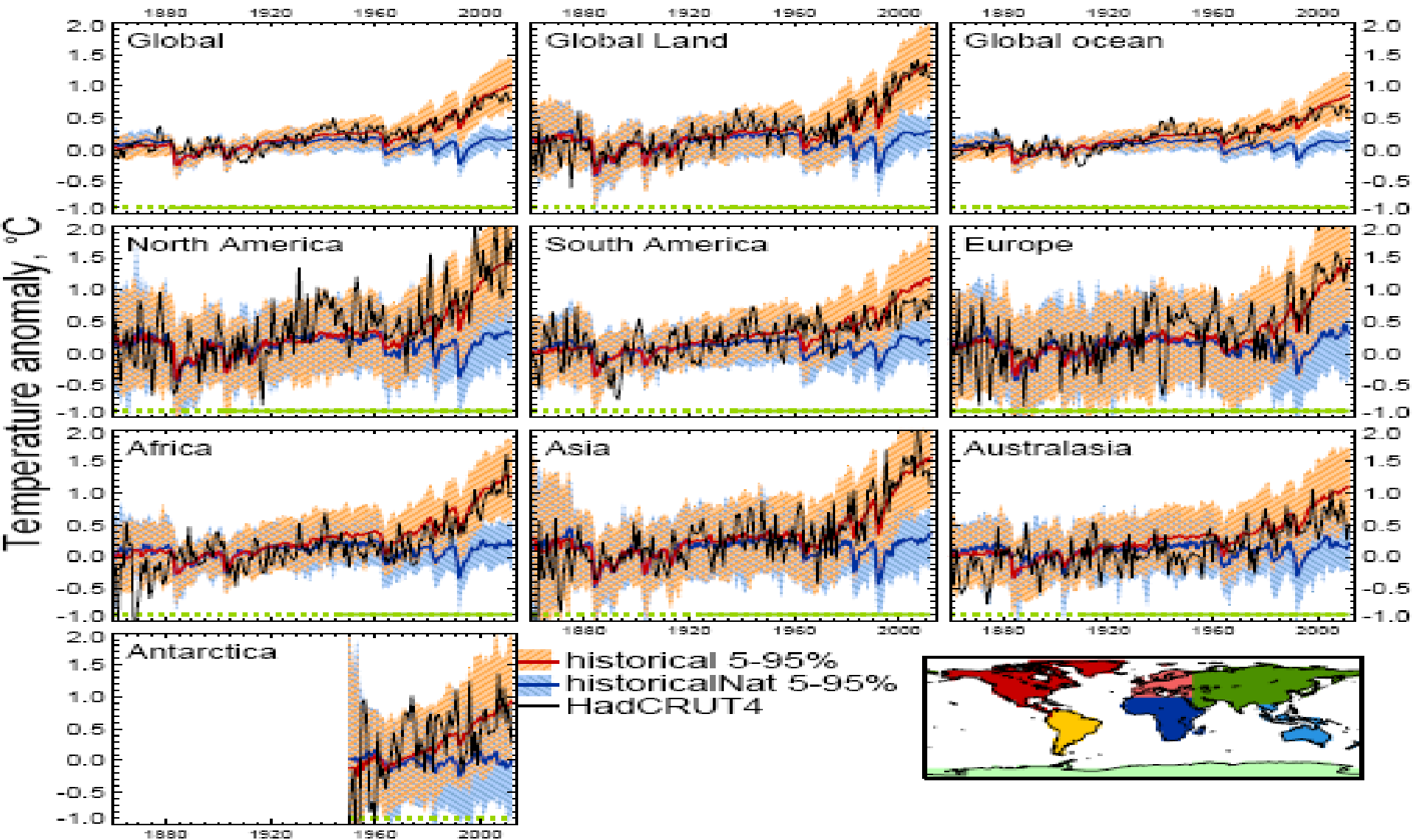
# Recent climate change report

Observed change in average surface temperature 1901–2012



Planet has warmed by 0.85 K over 1880–2012

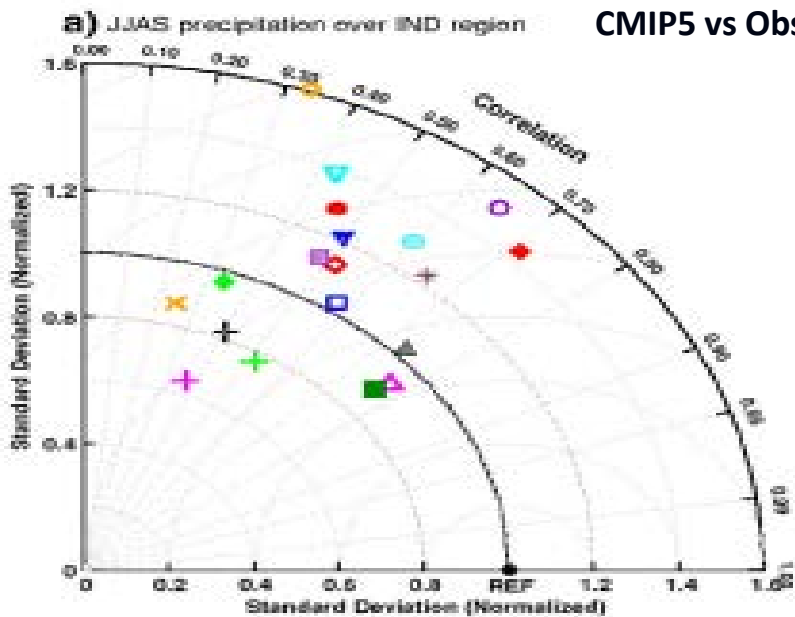
# Climate Change 2013: WG1 contribution to IPCC Fifth Assessment Report



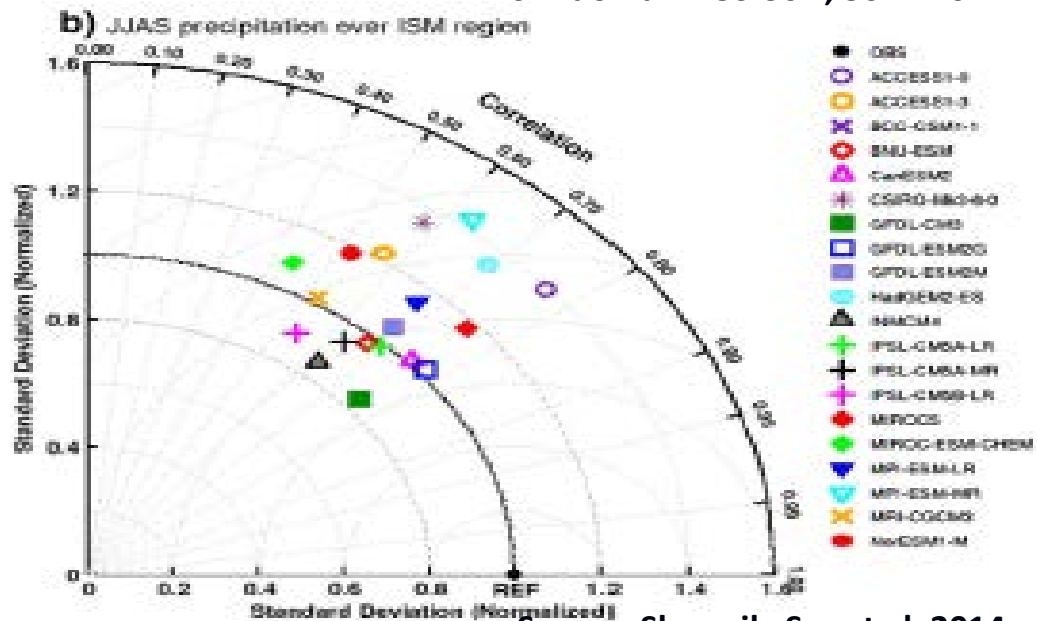
**Figure 10.7:** Global, land, ocean and continental annual mean temperatures for CMIP3 and CMIP5 historical (red) and historicalNat (blue) simulations (multi-model means shown as thick lines, and 5–95% ranges shown as thin light lines) and for HadCRUT4 (black). Mean temperatures are shown for Antarctica and six continental regions formed by combining the sub-continental scale regions defined by Seneviratne et al. (2012). Temperatures are shown with respect to 1880–1919 for all regions apart from Antarctica where temperatures are shown with respect to 1950–2010. Adapted from Jones et al. (2013).

## Wide variations among CMIP5/ CMIP3 models in capturing the South Asian monsoon

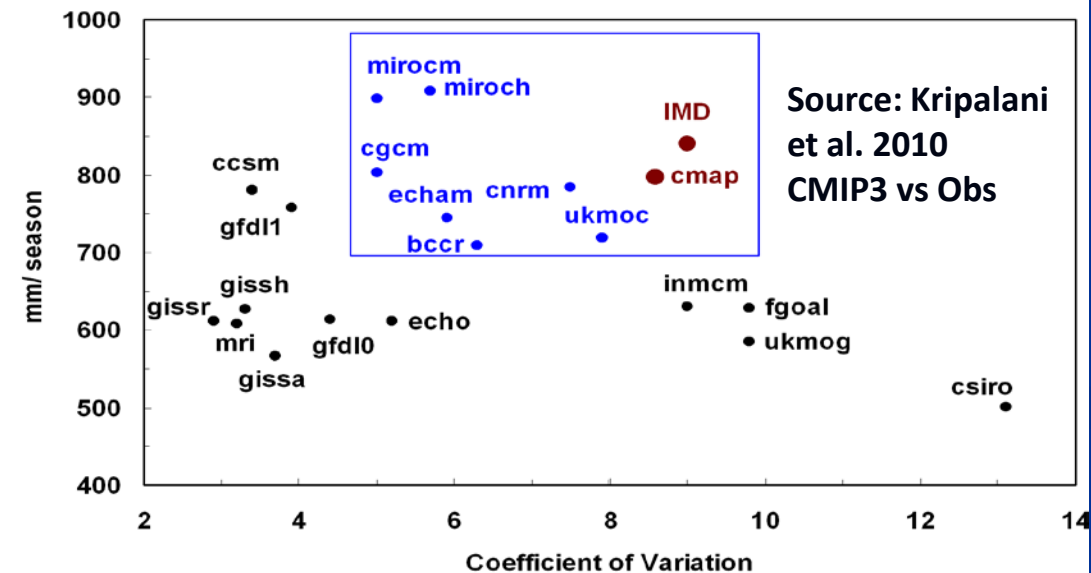
## Indian Land: CMIP5 vs Obs



ISM domain 15S-30N, 50E-120E



Source: Sharmila Sur et al. 2014



Realism of present-day  
climate simulation is an  
essential requirement for  
reliable assessment of future  
changes in monsoon



# Science of climate change

Detection, attribution & projection of global climate and regional monsoons, variability and change

## Roadmap for Earth System Model (ESM) development

- Start with an atmosphere-ocean coupled model with realistic mean climate
  - Fidelity in capturing the global and monsoon climate
  - Realistic representation of monsoon interannual variability
  - Features of ocean-atmosphere coupled interactions
  - ...
- Include components / modules of the ESM
  - Biogeochemistry
  - Interactive Sea-ice
  - Aerosol and Chemistry Transport
  - ...

# Basic modeling framework: Coupled Forecast System (CFS-2) T126L64

**Formal agreement for collaboration:** The Ministry of Earth Sciences, Govt. of India and NOAA, USA in 2011. Implement the NCEP CFS-2 model at IITM, Pune for seasonal prediction of the Indian monsoon.

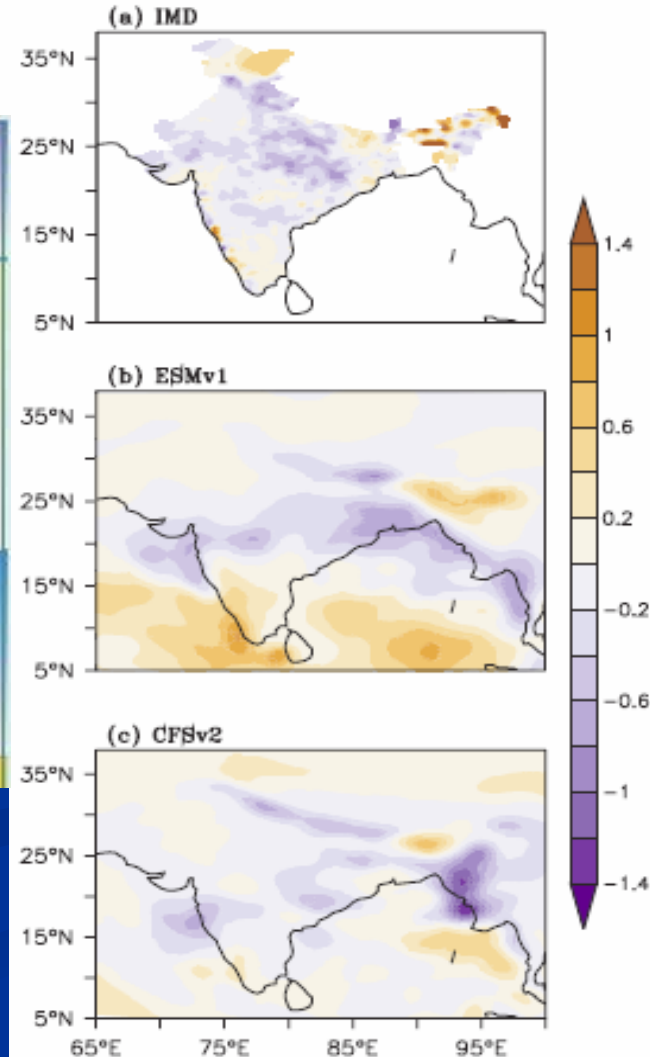
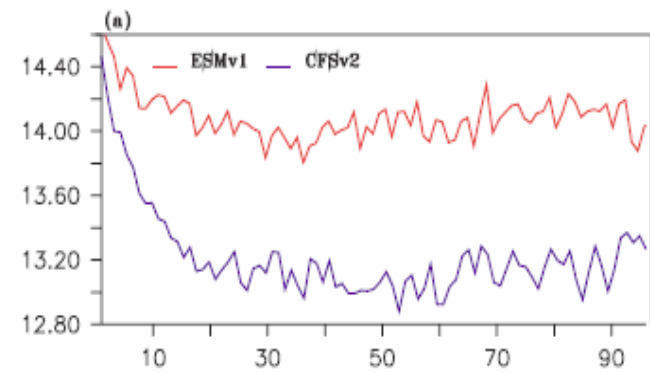
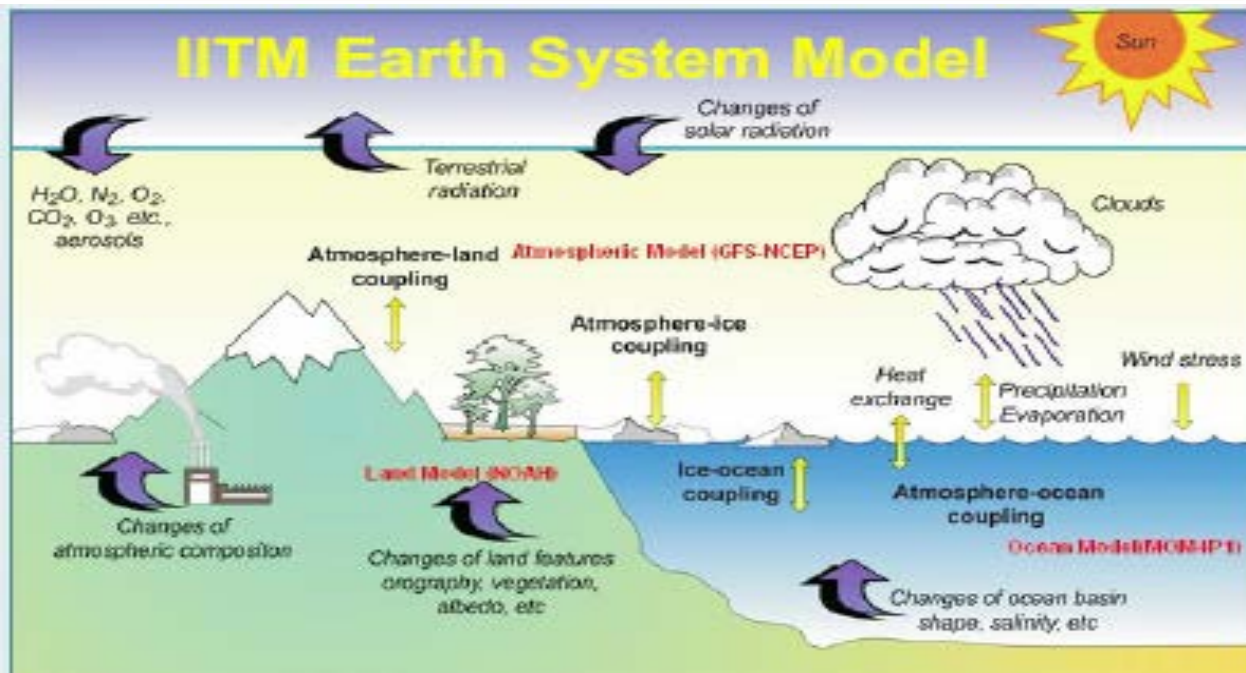
- The **NCEP CFS** Components
- Atmospheric **GFS (Global Forecast System) model**
  - – T126 ~ 110 km; vertical: 64 sigma – pressure hybrid levels
  - – Model top 0.2 mb
  - – Simplified Arakawa-Schubert convection (Pan)
  - – Non-local PBL (Pan & Hong)
  - – SW radiation (Chou, modifications by Y. Hou)
  - – Prognostic cloud water (Moorthi, Hou & Zhao)
  - – LW radiation (GFDL, AER in operational wx model)
  - – Land surface processes (Noah land model)
- Interactive Ocean: **GFDL MOM4** (Modular Ocean Model, ver.4)
  - – 0.5 deg poleward of 10°N and 10°S; and 0.25 deg near equator (10°S – 10°N)
  - – 40 levels
  - – Interactive sea-ice

# THE IITM EARTH SYSTEM MODEL

## Transformation of a Seasonal Prediction Model to a Long-Term Climate Model

BY P. SWAPNA, M. K. ROXY, K. APARNA, K. KULKARNI, A. G. PRAJEESH,  
K. ASHOK, R. KRISHNAN, S. MOORTHY, A. KUMAR, AND B. N. GOSWAMI

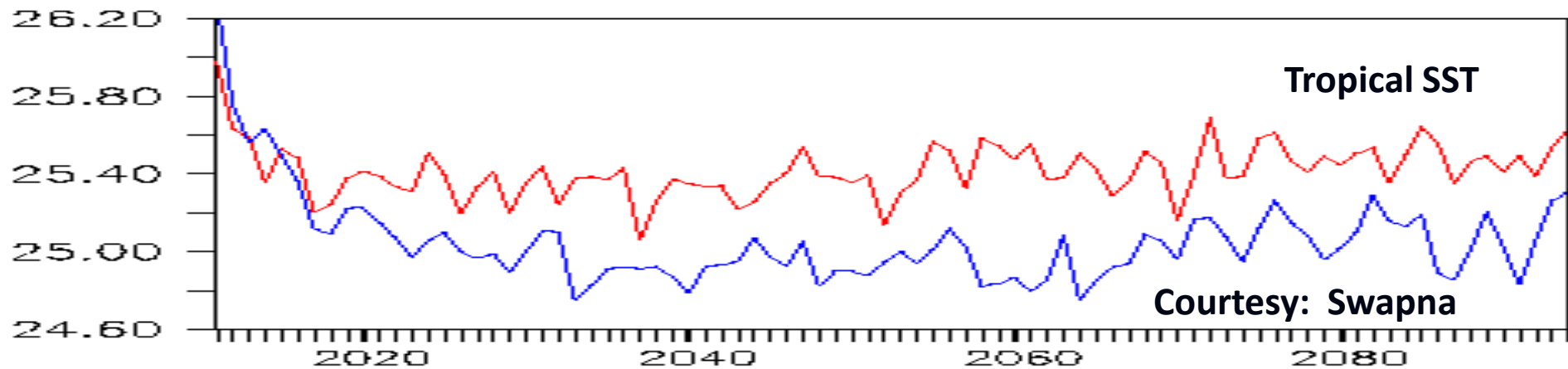
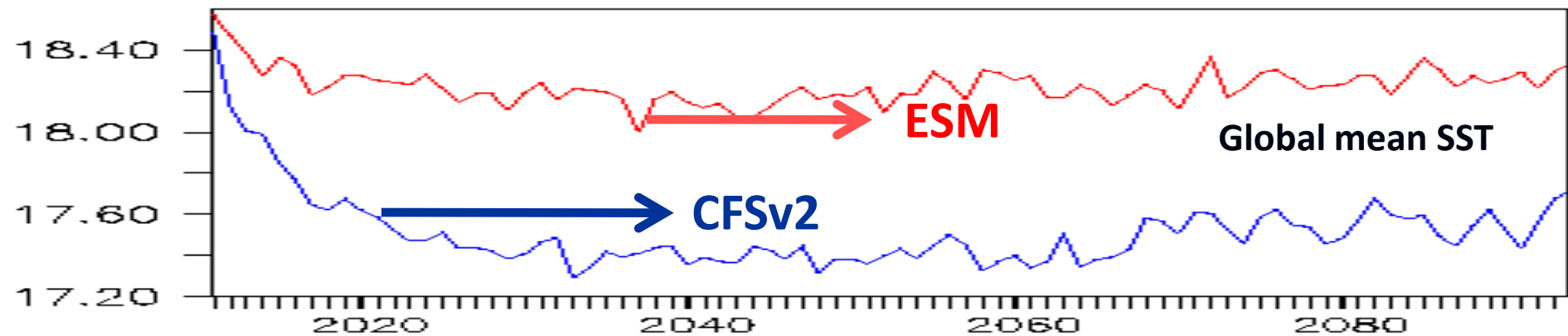
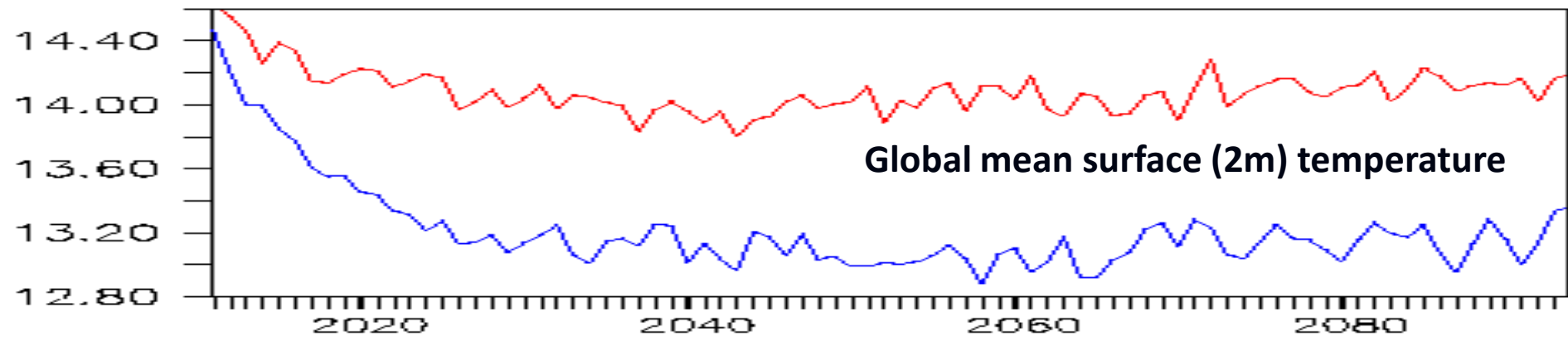
This work documents the fidelity of the newly developed Indian Institute of Tropical Meteorology climate model simulations and demonstrates its suitability to address the climate variability and change issues relevant to the South Asian monsoon.



Atmosphere: T126 spectral ( $\sim 190$  km), 64 vertical levels – ESMv1

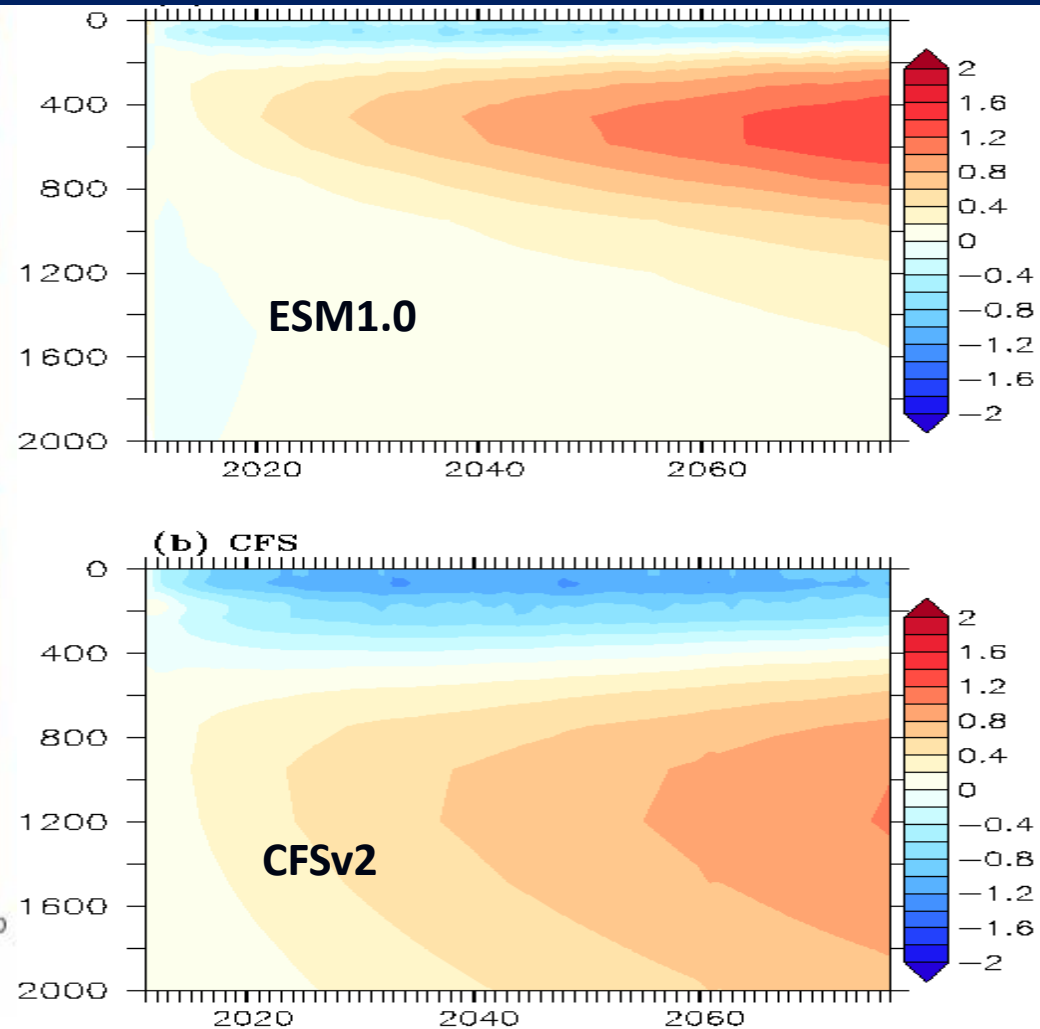
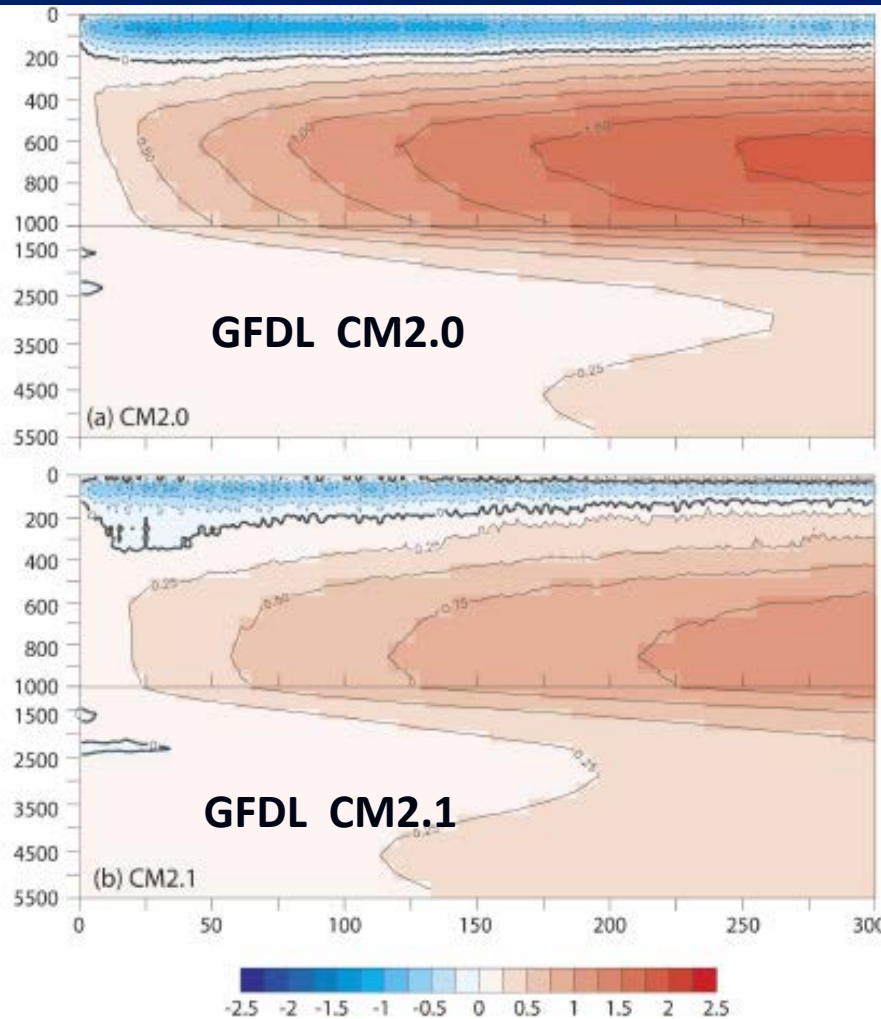
Ocean : 0.5 deg grid,  $\sim 0.25$  deg between 10N-10S, 40 vertical levels

# Annual mean temperature



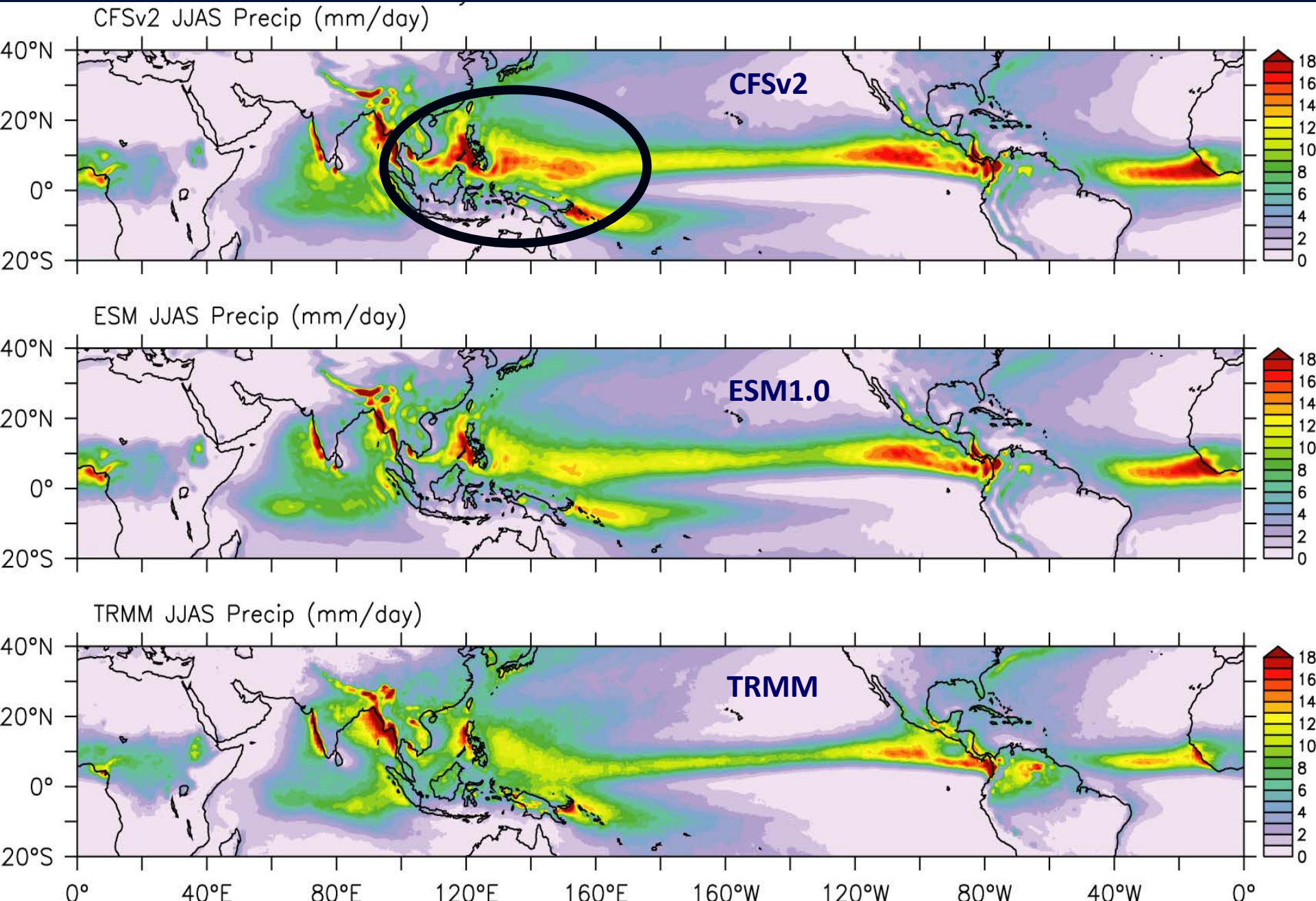


Coupled models drift towards a more equilibrated state. Initial rapid cooling of SST followed by warming trend. Significant subsurface drifts seen through multiple centuries of simulation. Vertical redistribution of heat with tendency of cooling in upper layers and warming in the subsurface – [Delworth et al. 2006](#)

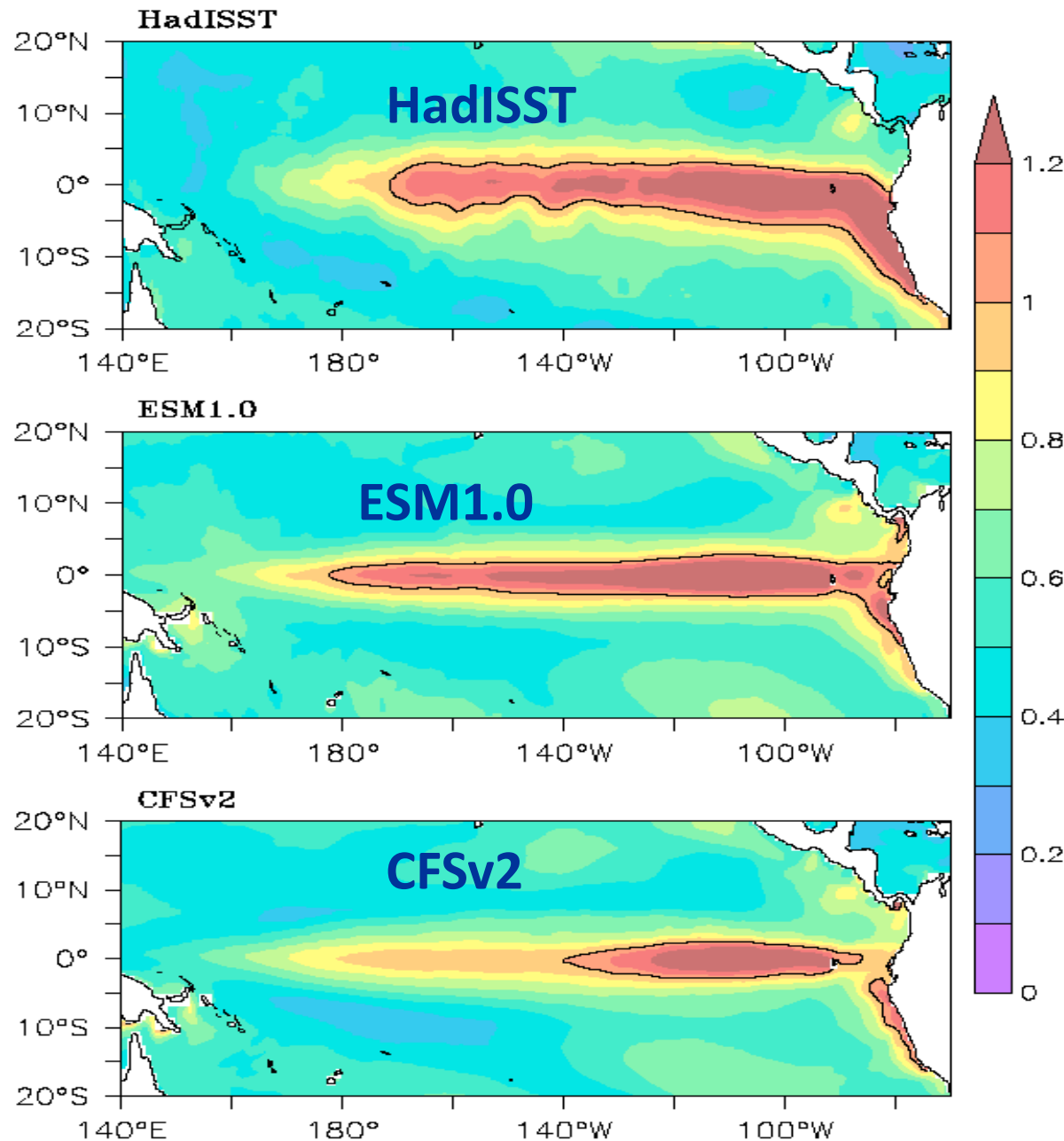


Differences between simulated and observed long-term global-mean ocean temperature as a function of depth and time.

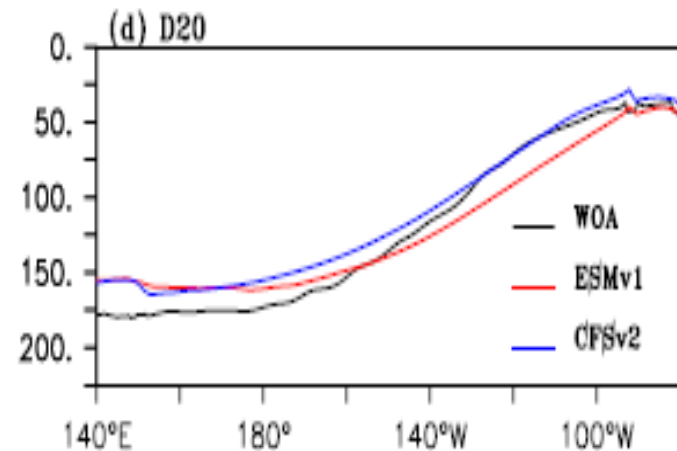
# Precipitation (mm day<sup>-1</sup>): JJAS mean



# Interannual variability: Standard deviation of SST

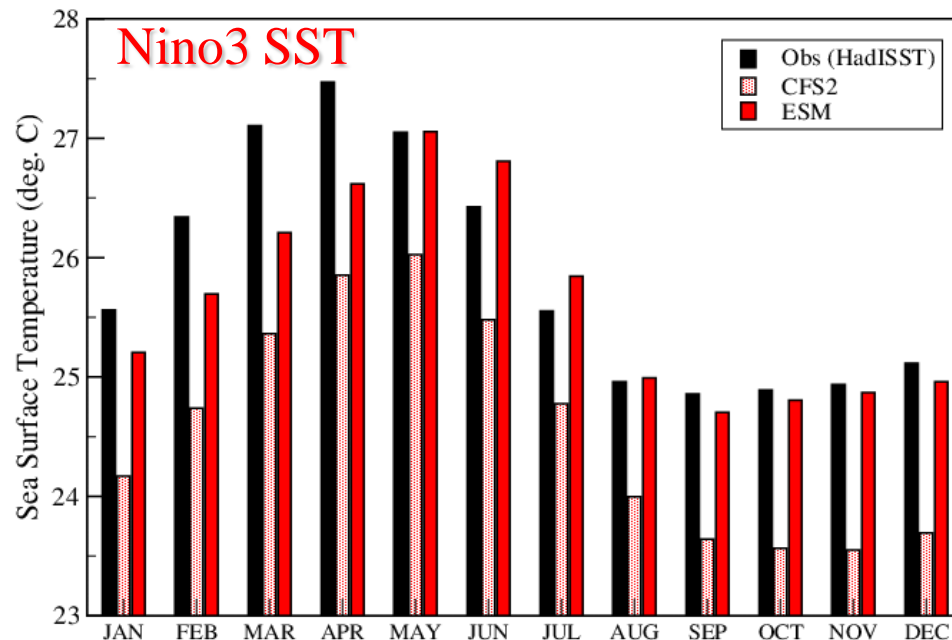
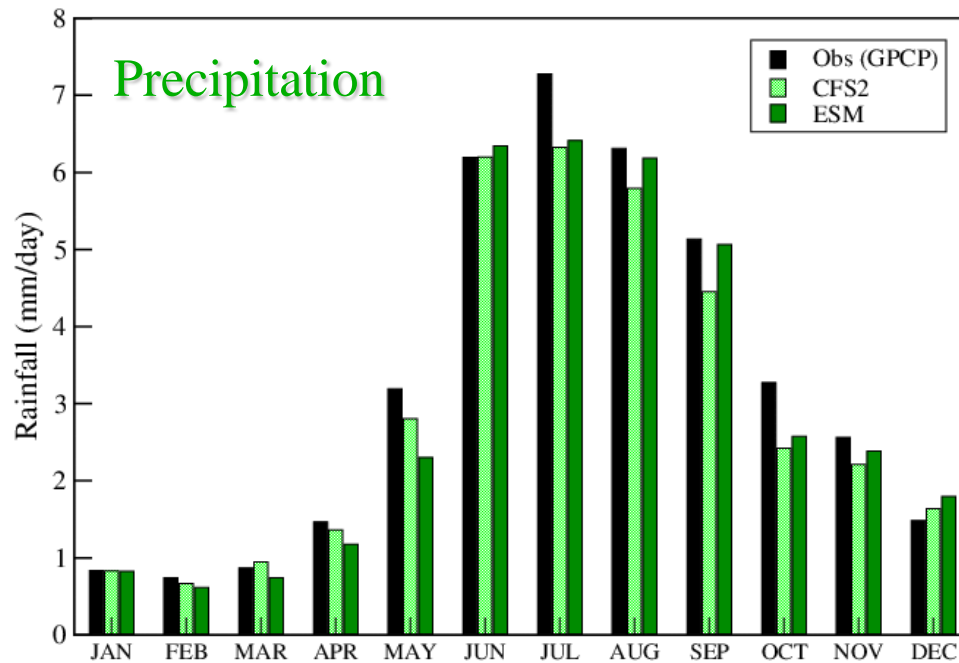


Interannual variability of Pacific SST in CFSv2 is mostly confined to the eastern equatorial Pacific; more realistic in ESM1.0



Courtesy: Swapna



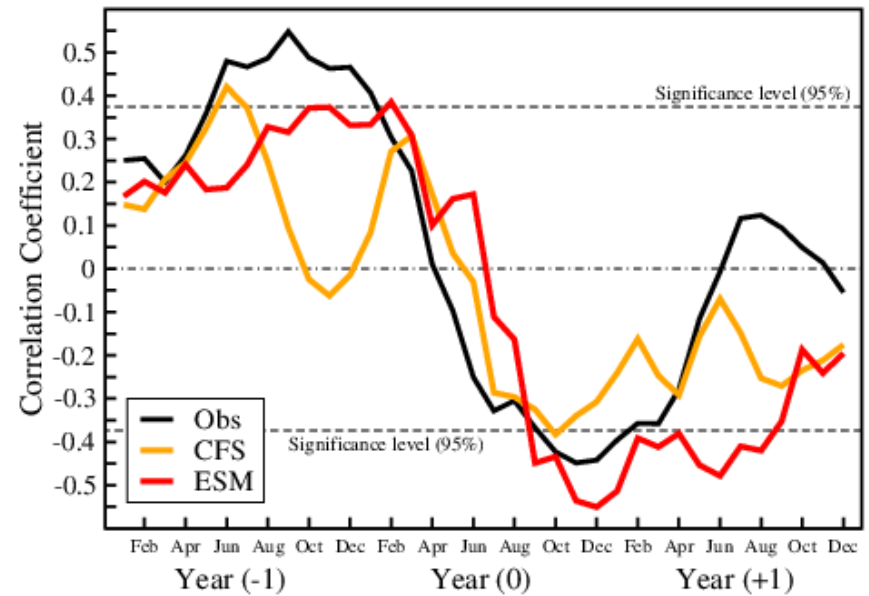


## Precipitation

(5N-35N; 65E-95E)

Indian (land + ocean)

## ENSO-Monsoon relationship



Lagged correlation between ISMR and Nino3 SST in the preceding/following months

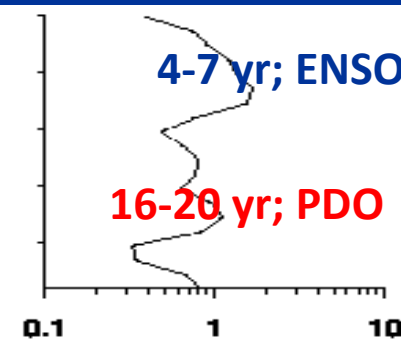
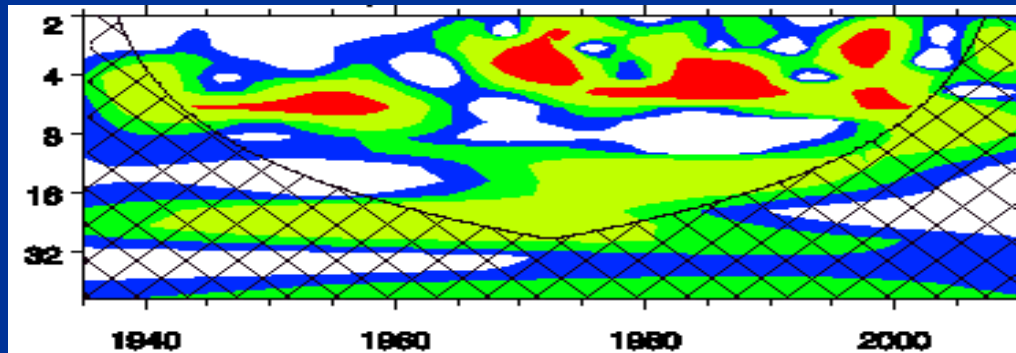
CFS2 : 30 years (yr17-yr46)

ESM : 30 years (yr17-yr46)



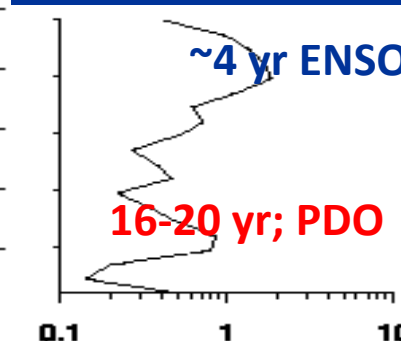
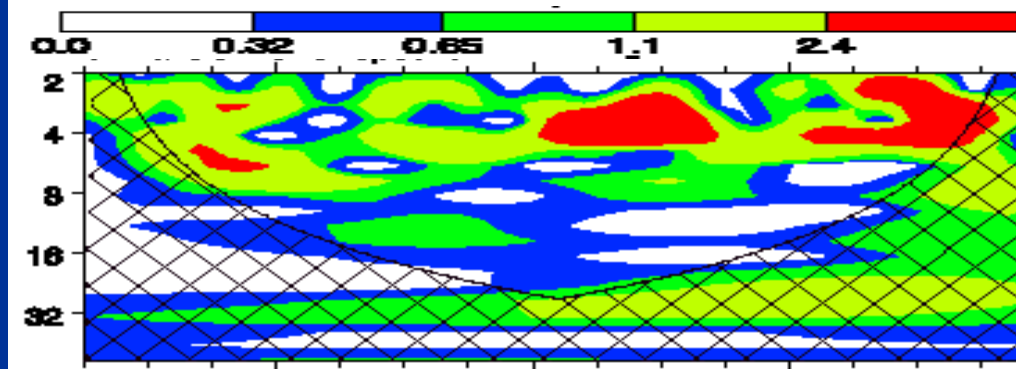
# Wavelet Power Spectrum of PC1 time-series. Power ( $C$ )<sup>2</sup> as a function of period and time

Period (year)



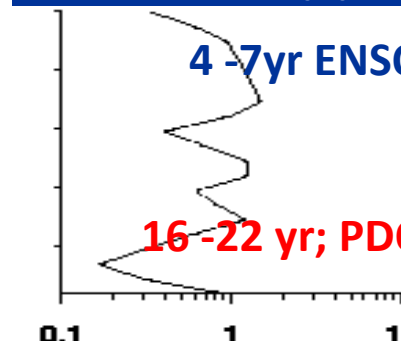
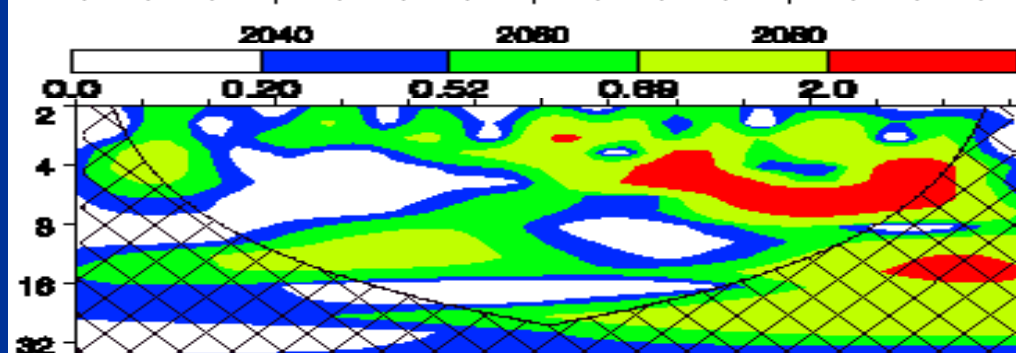
**HadISST**

Variance ( $C$ )<sup>2</sup>



**ESM1.0**

Variance ( $C$ )<sup>2</sup>



**CFSv2**

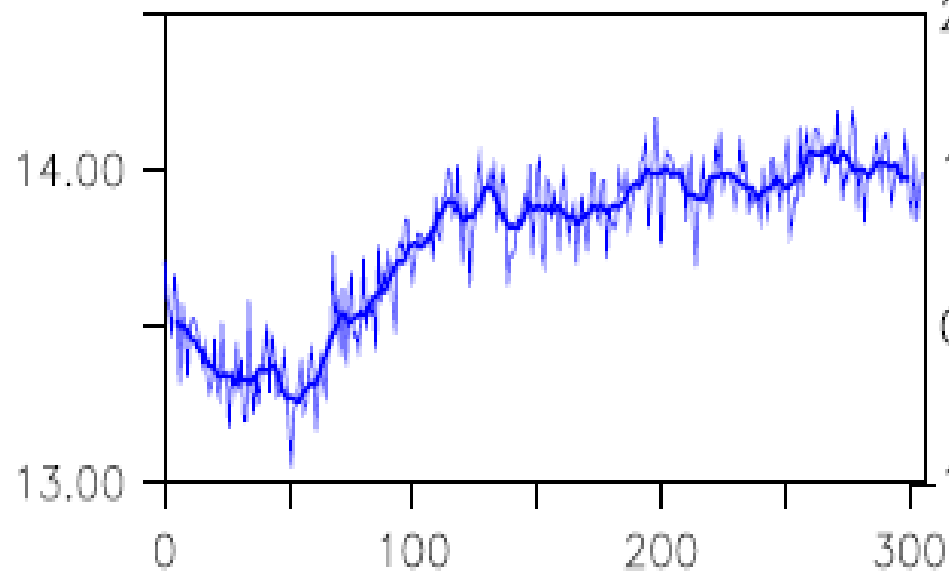
Variance ( $C$ )<sup>2</sup>

Time (year)

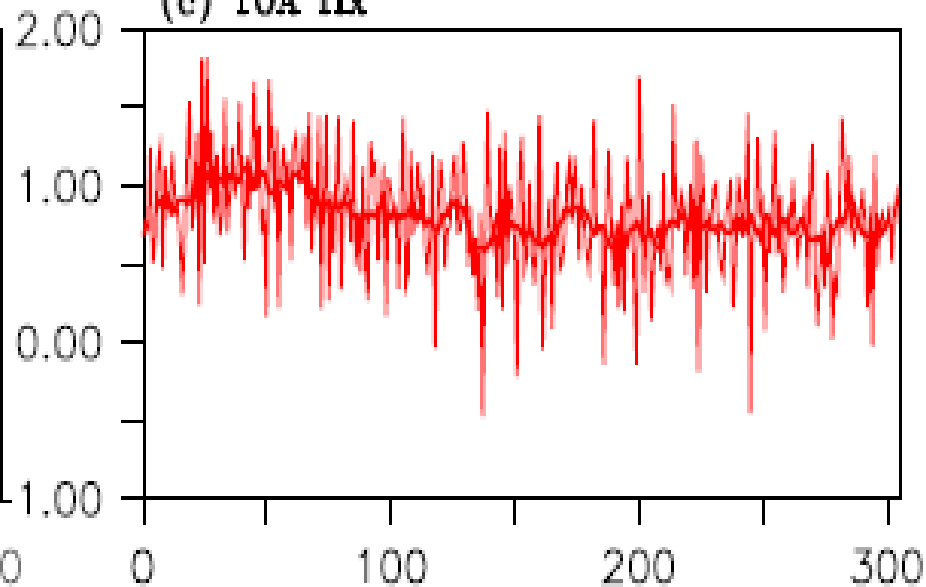
Courtesy: Swapna

## **Recent improvements in IITM ESM**

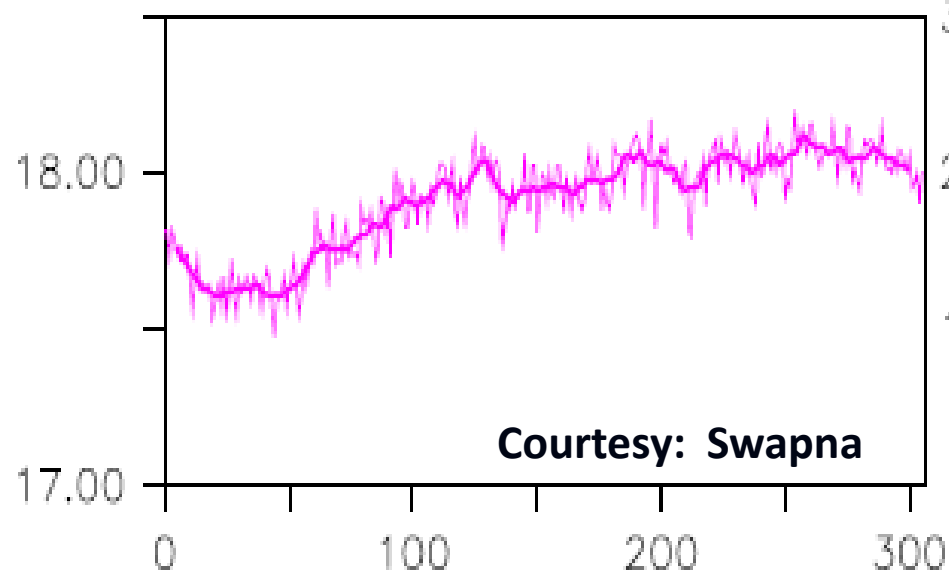
**(a) T<sub>air</sub>**



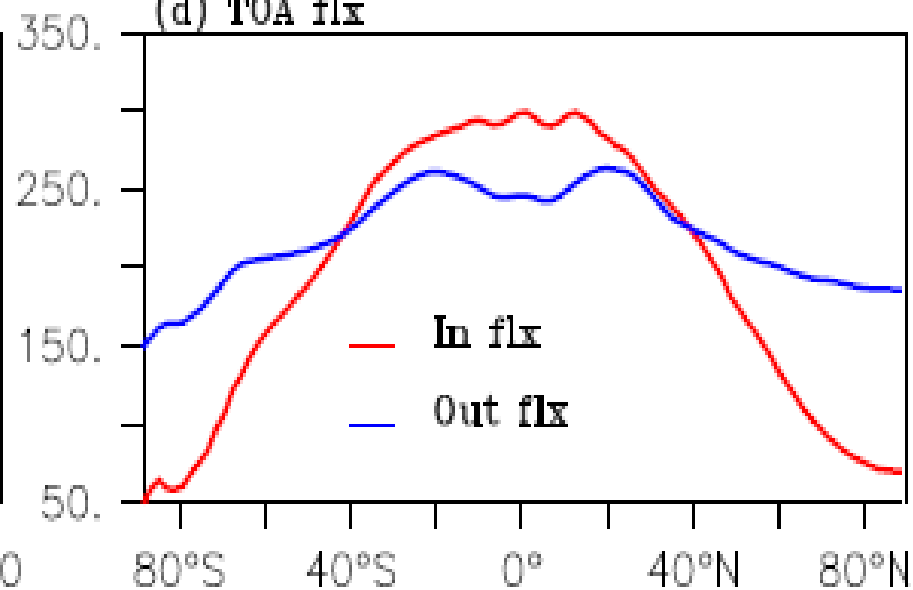
**(c) TOA flx**



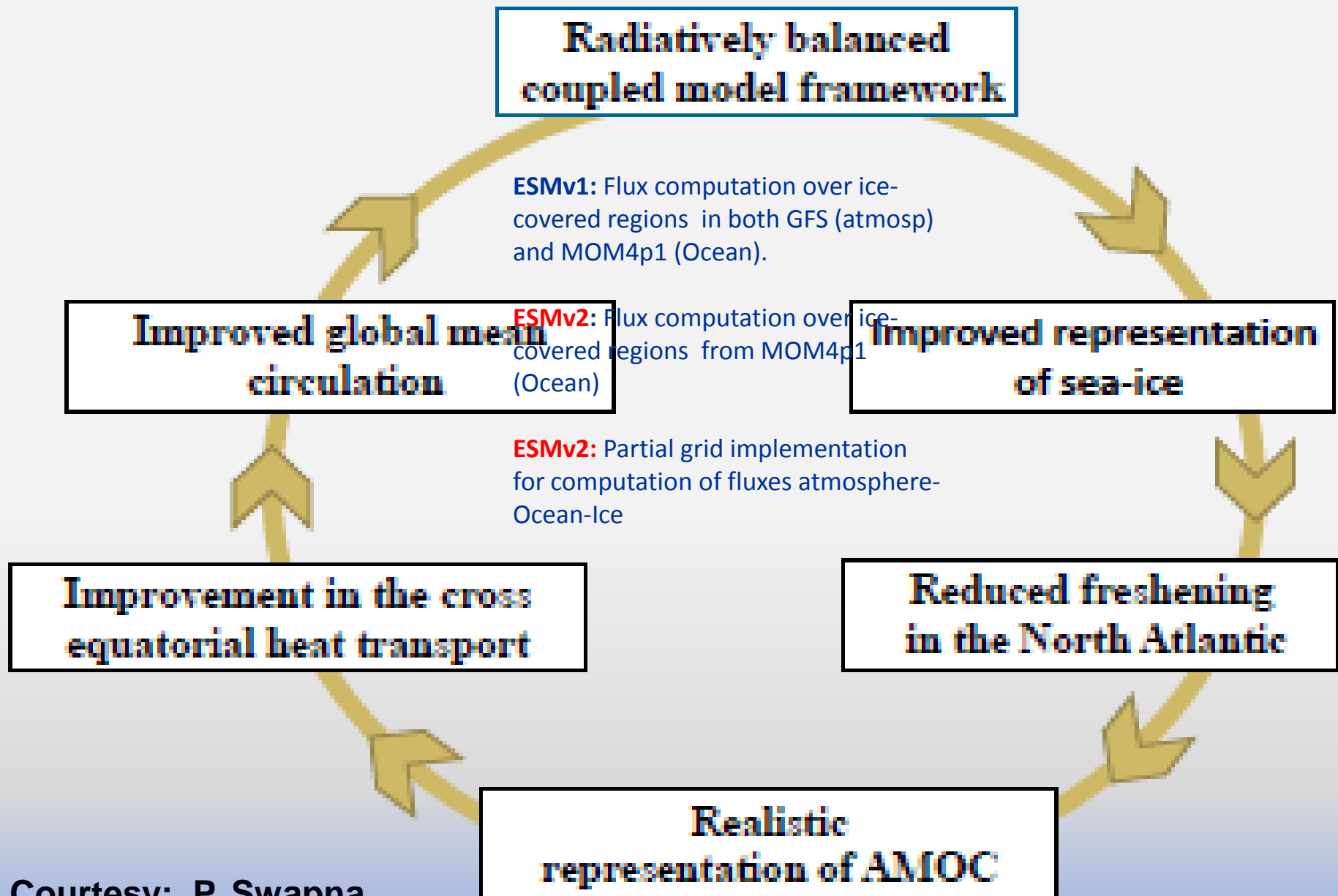
**(b) SST**



**(d) TOA flx**

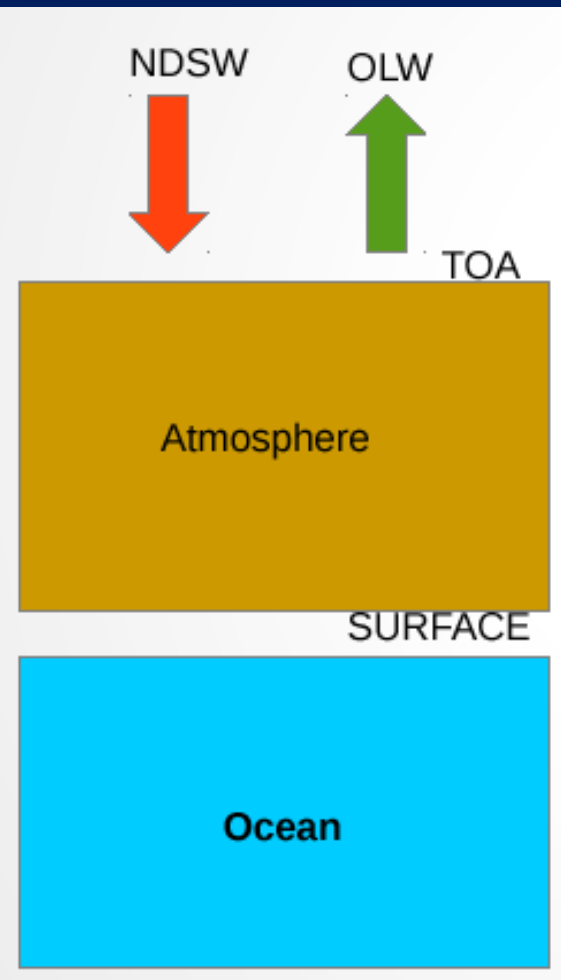


# Ocean-Atmosphere coupled feedbacks in IITM-ESM



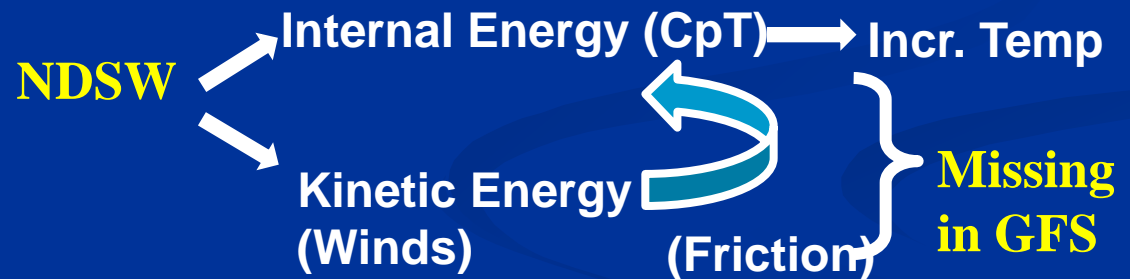


# TOA Energy Balance



**NDSW – Net downward Short wave flux at TOA**

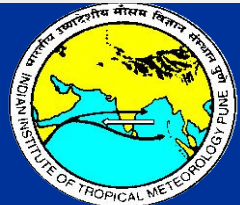
**OLW – Outgoing Longwave flux (depends on layer temperature according to Stefan Boltzman law)**



**TKE dissipation heating (Han)**

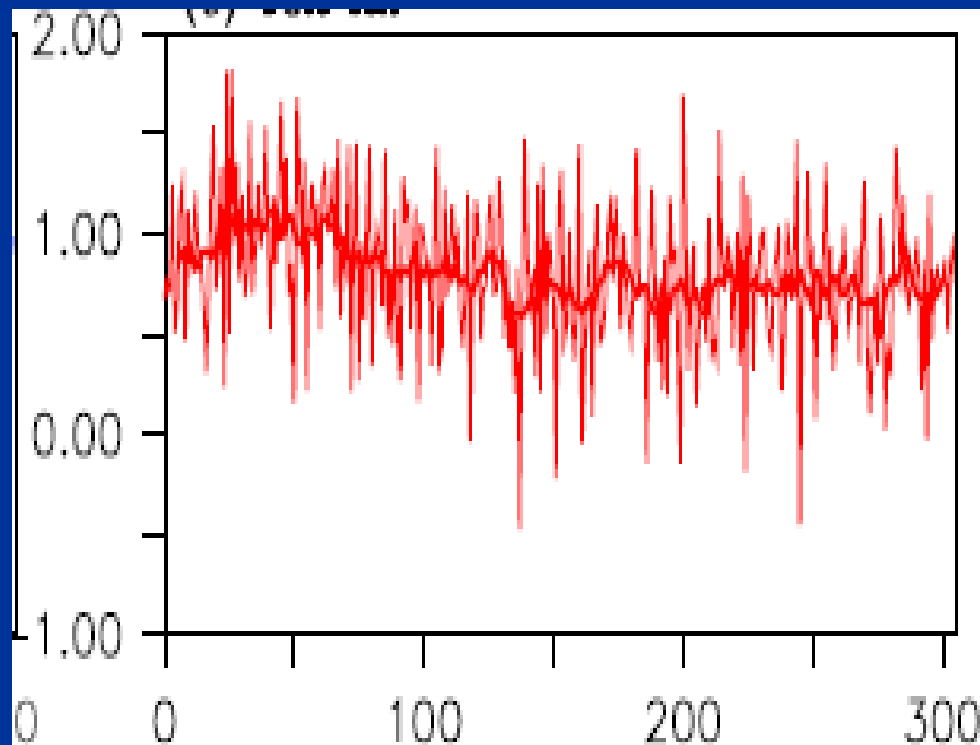
$$\varepsilon = \underbrace{-K_h \frac{g}{\theta_v} \frac{d\theta_v}{dz}}_{\text{buoyancy production}} + \underbrace{K_m \left| \frac{d\mathbf{u}}{dz} \right|^2}_{\text{shear production}}$$

Minimize atmospheric energy loss – Bretherton et al. 2012

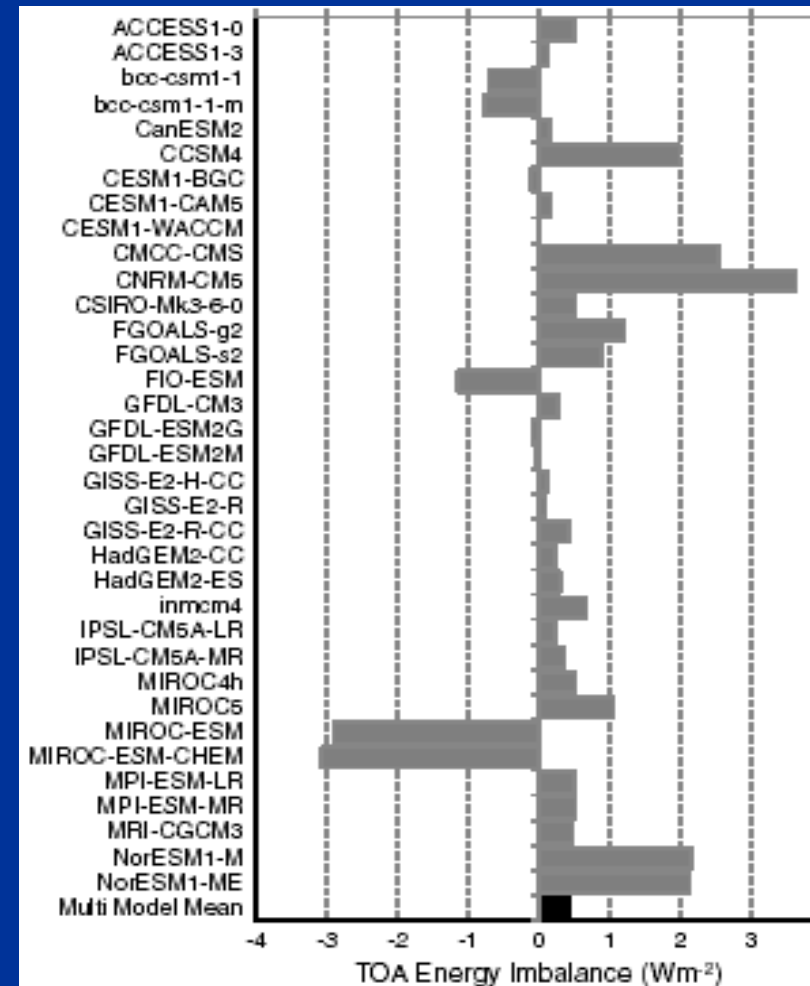


Courtesy: Prajeesh

## Energy Balance in IITM ESMv2

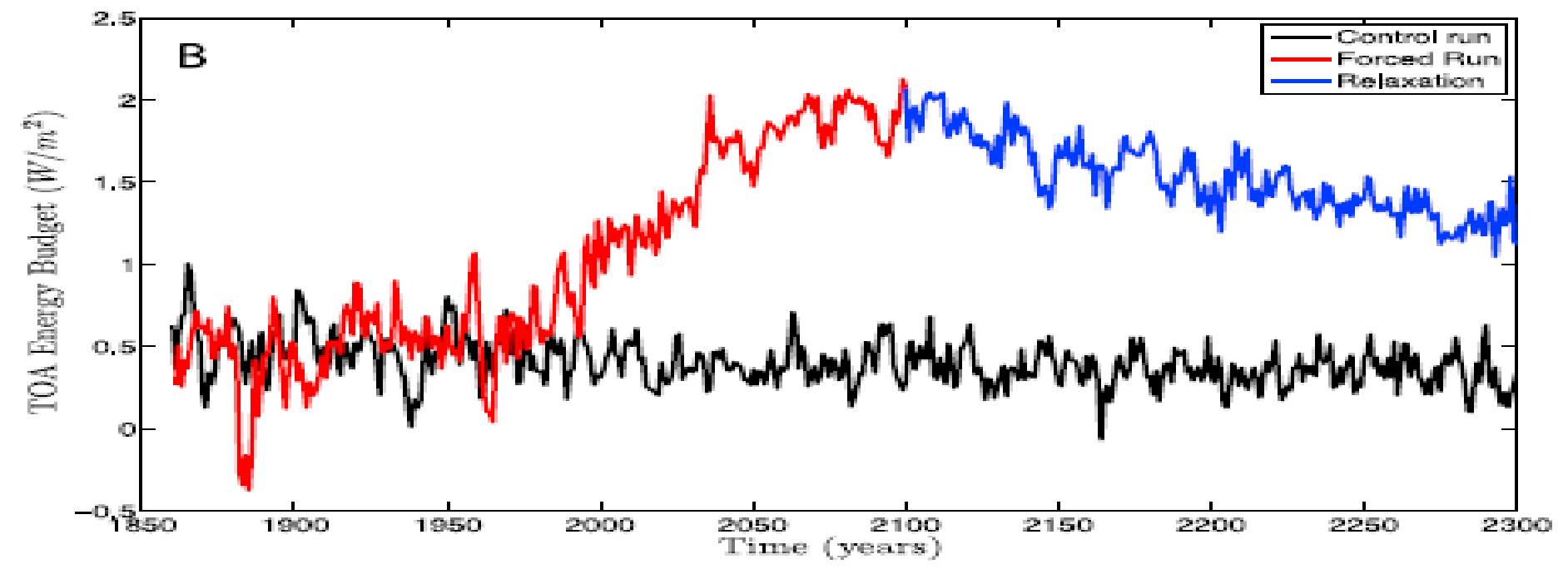


## TOA Energy Imbalance (CMIP5 Models)



Preindustrial TOA ( $\text{Wm}^{-2}$ )  
 Energy imbalance for CMIP5  
 Models (Forster et al., 2013)

Time-series of TOA energy budget (GFDL2.1 CM9) – V. Lucarini, F. Ragone, 2011, Rev. Geophy



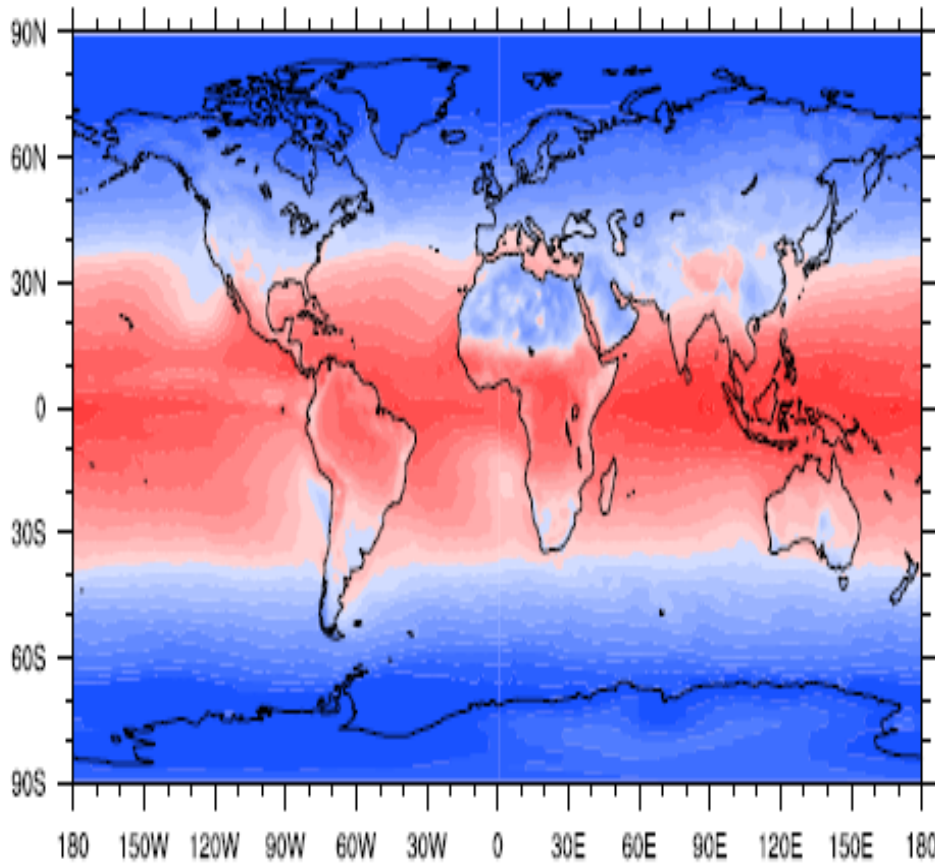
Black line is the preindustrial run. The red line shows the 20<sup>th</sup> century simulation and the 21<sup>st</sup> century portion of the SRES A1B simulation (started from the end of the 20<sup>th</sup> century simulation). The blue line shows the 22<sup>nd</sup> and 23<sup>rd</sup> century SRES A1B simulation

# Net Radiation ( $\text{W m}^{-2}$ ) at TOA

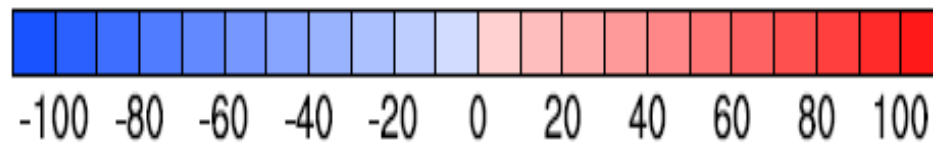
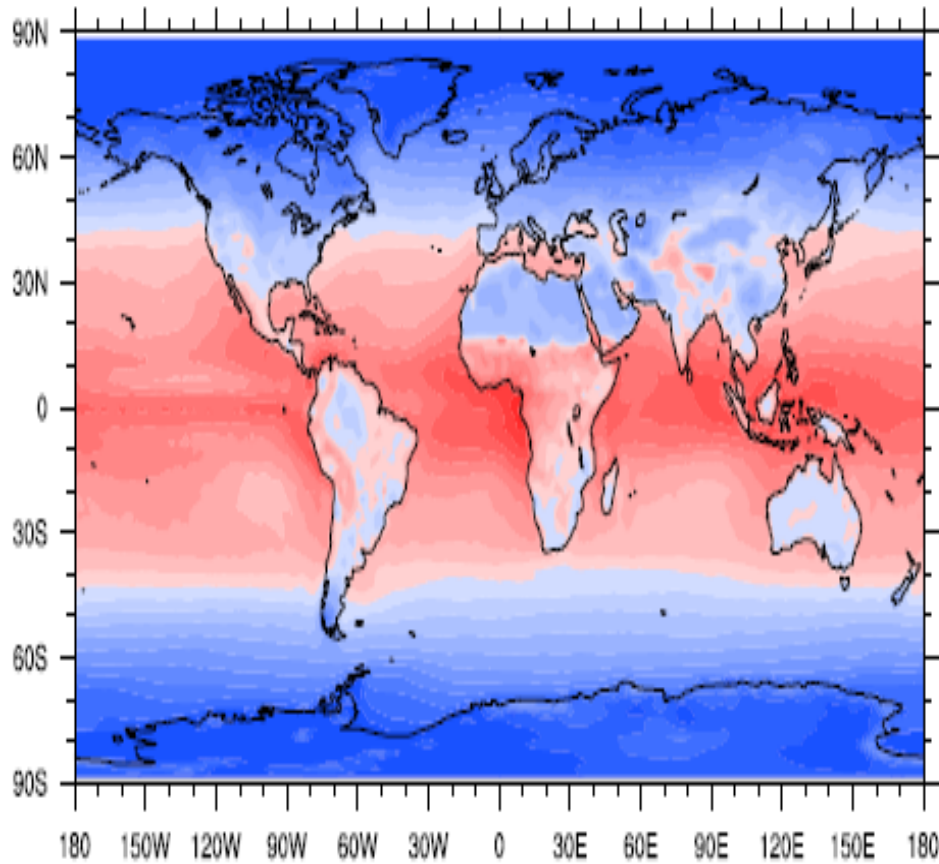
Obs (CERES)

IITM-ESMv2

a) Observation (CERES)



b) ESMv2





# Energy Balance in IITM ESM

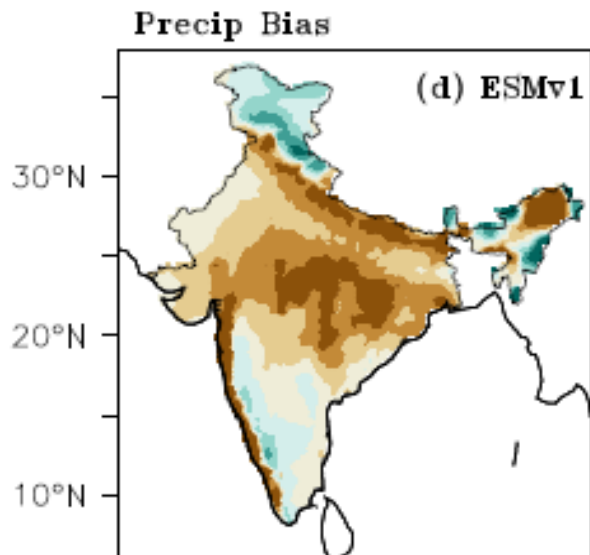
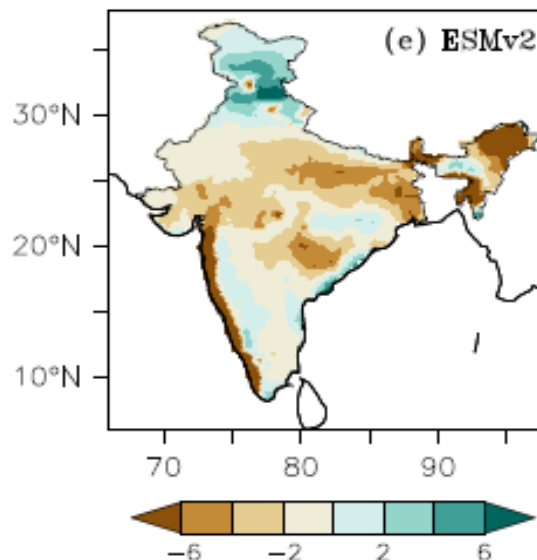
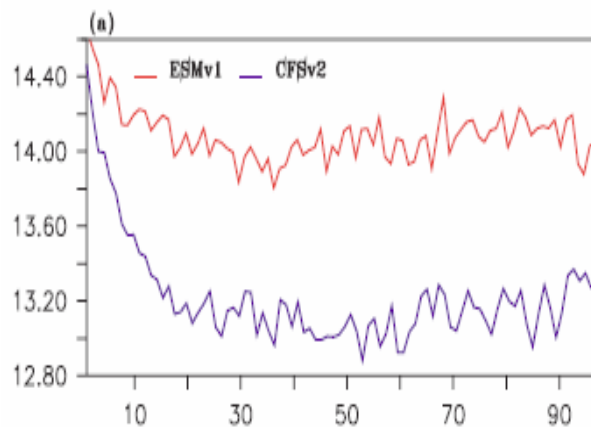
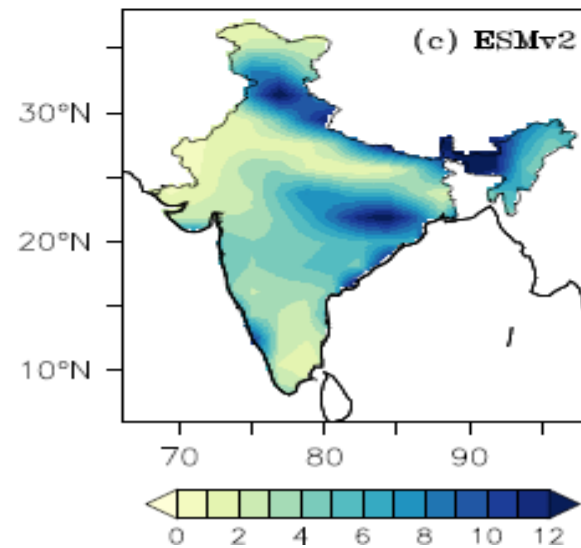
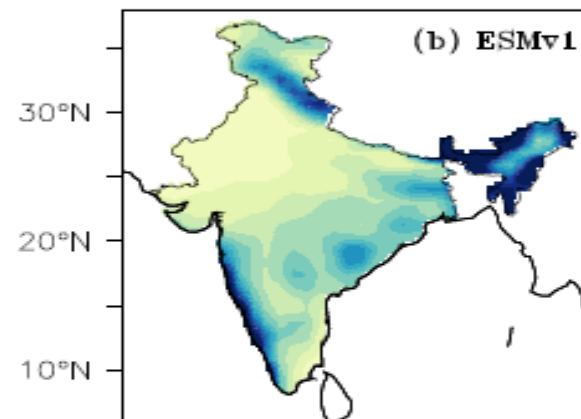
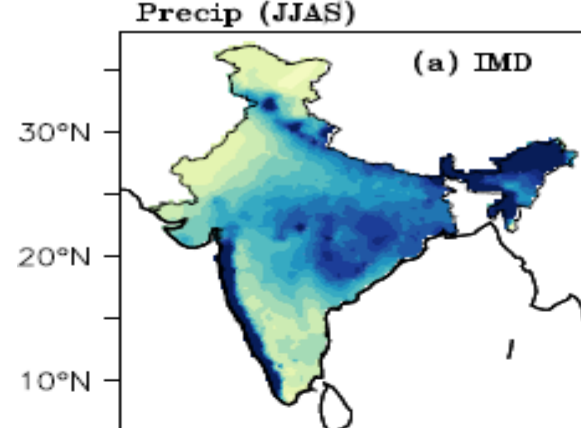
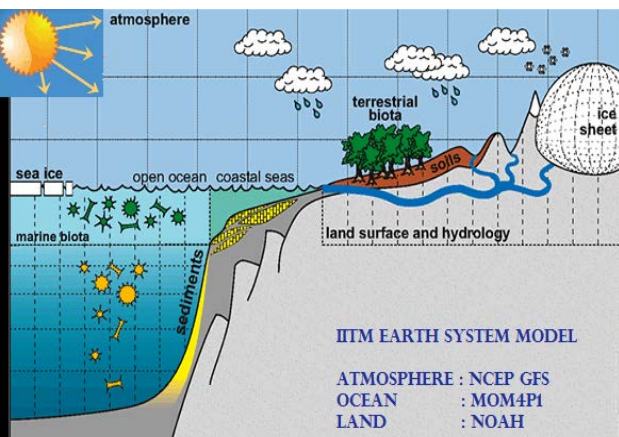
	<b>Net flux TOA (W m<sup>-2</sup>)</b>	<b>Net Flux Surface (W m<sup>-2</sup>)</b>	<b>Difference (W m<sup>-2</sup>)</b>
<b>ESMv1 (T126)</b>	<b>6.6</b>	<b>1.2</b>	<b>5.4</b>
<b>ESMv2 (T62)</b>	<b>0.80</b>	<b>0.75</b>	<b>0.05</b>

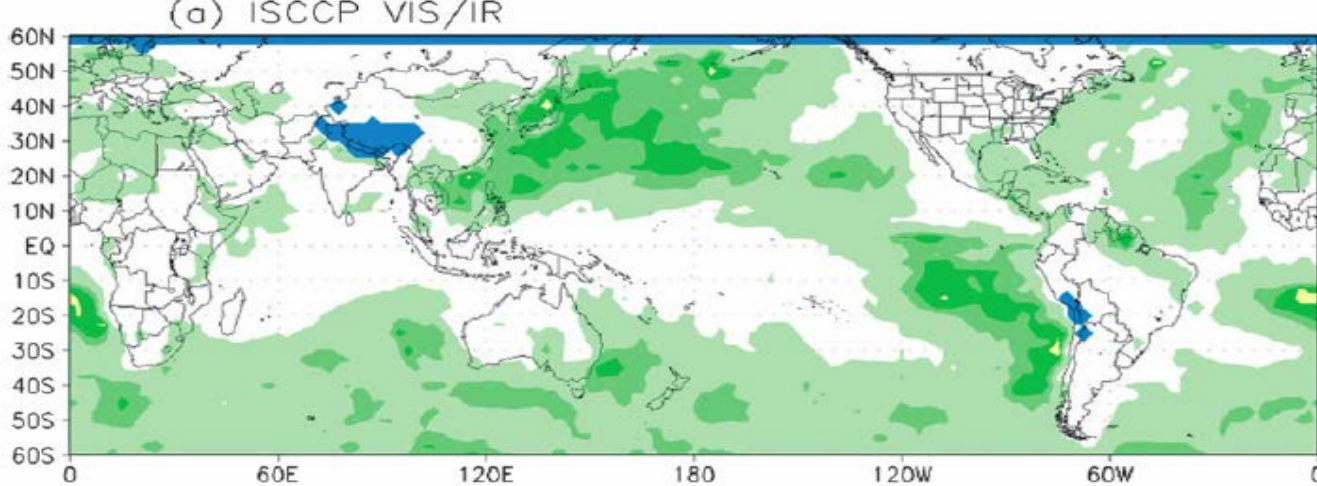
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## Transformation of a Seasonal Prediction Model to a Long-Term Climate Model

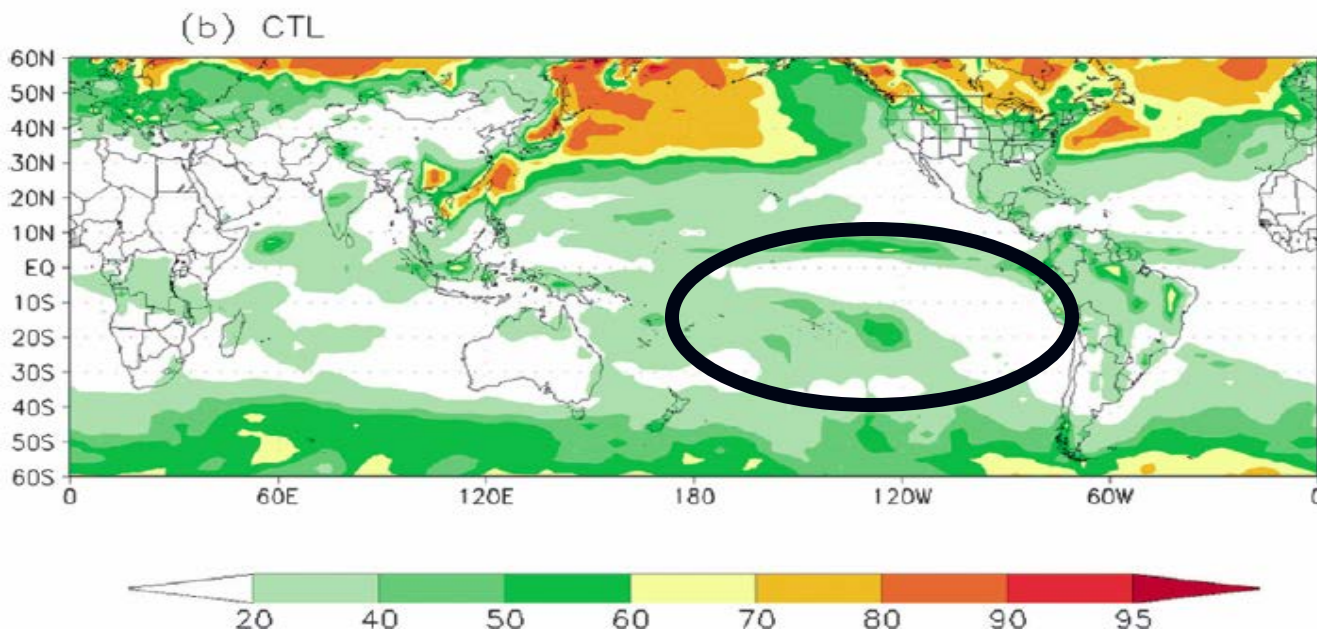
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This work documents the fidelity of the newly developed Indian Institute of Tropical Meteorology climate model simulations and demonstrates its suitability to address the climate variability and change issues relevant to the South Asian monsoon.





Monthly mean low cloud cover (%) for January 2003 from ISCCP (Rossow and Schiffer, 1991) VIS/IR satellite observations (blue color indicates 'no data' available).



Control simulation using the old shallow convection Scheme of NCEP GFS

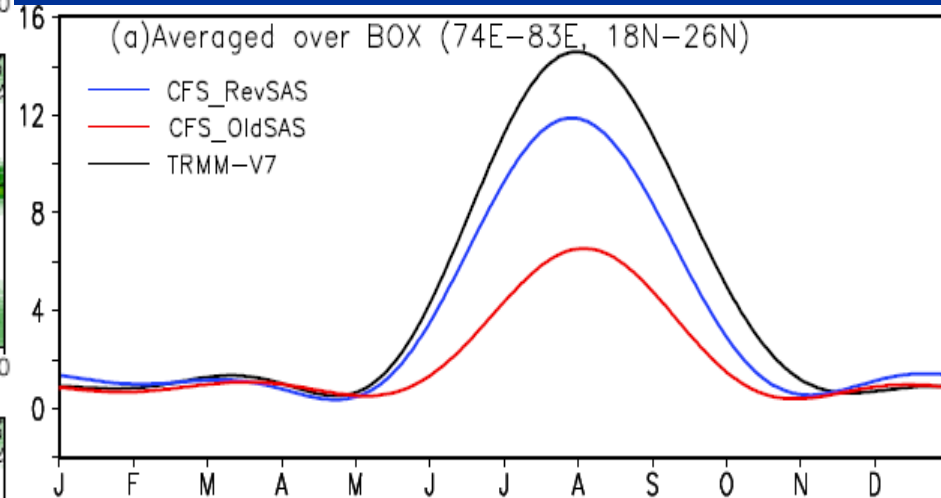
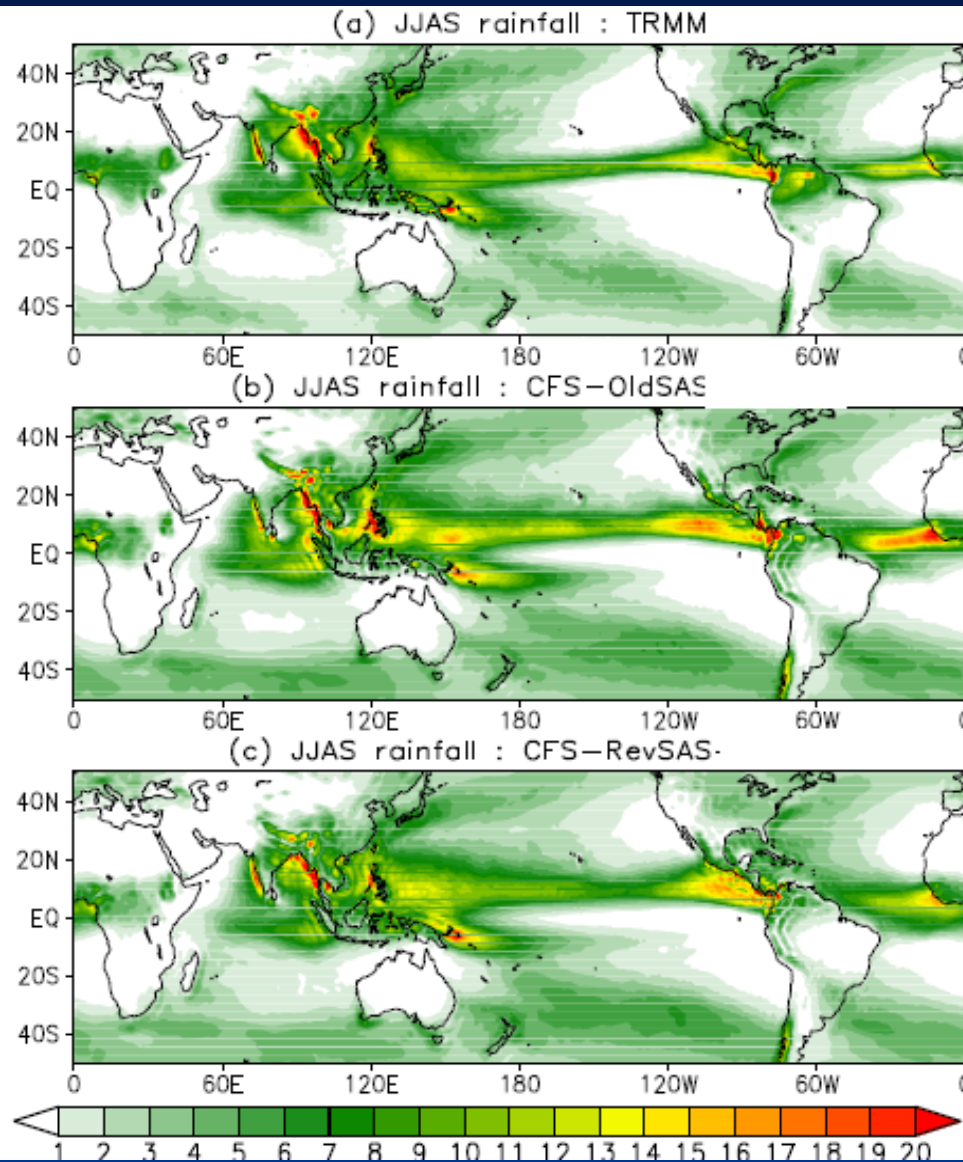
**Han and Pan, 2011**

Long-standing problems in NCEP GFS: Systematic underestimation of stratocumulus clouds in the eastern Pacific and Atlantic Oceans; and the frequent occurrence of unrealistic excessive heavy precipitation, the so-called grid-point storms

# Impact of Revised SAS (Simplified Arakawa Schubert) convective parameterization on monsoon rainfall simulation in CFSv2 - Malay, G, Phani, R.M, P. Mukhopadhyay

Climate Dynamics (2014)

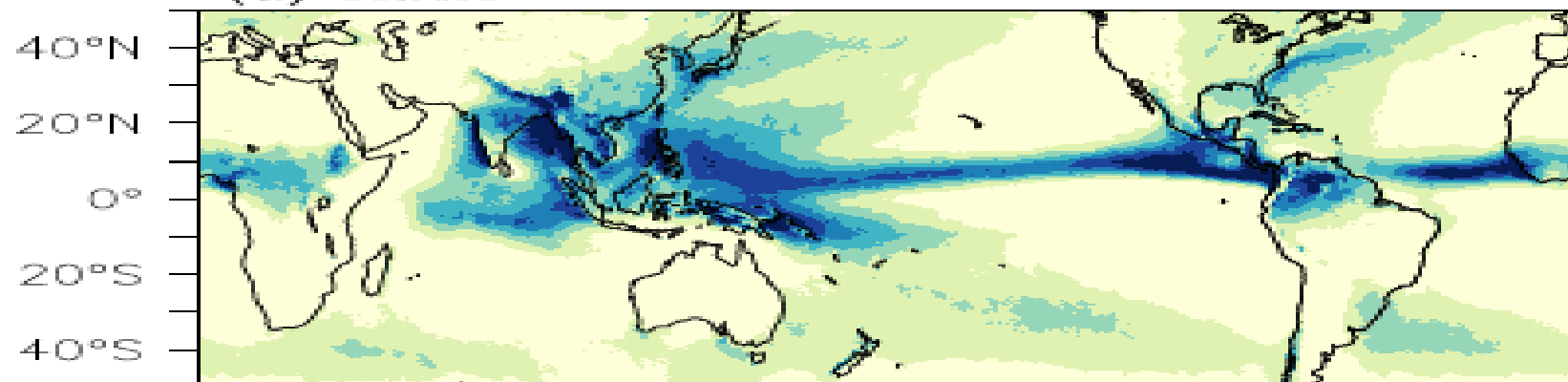
## Annual cycle of rainfall over Indian region



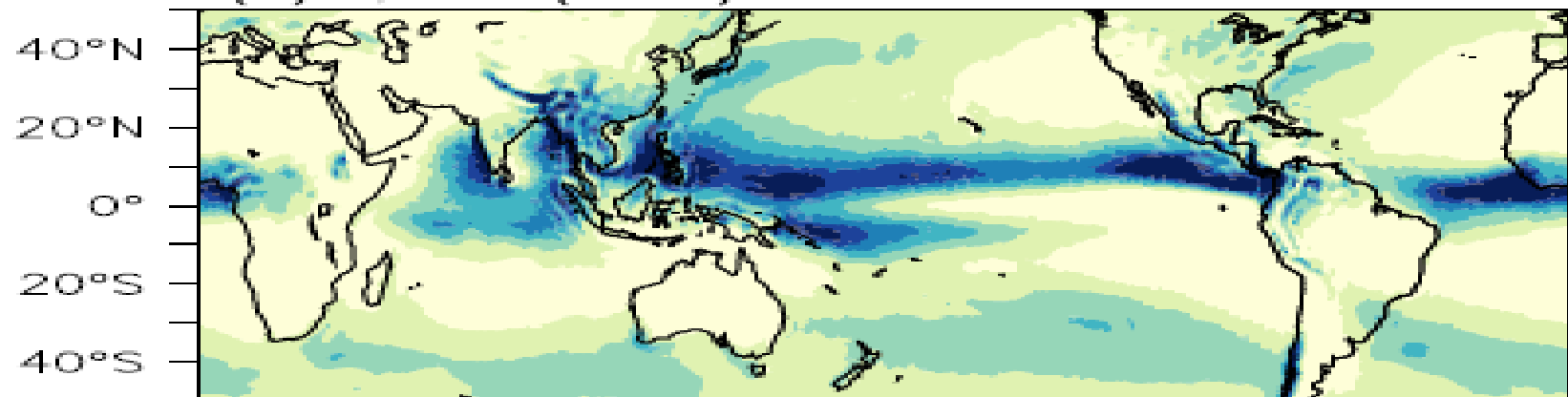
CFSv2 T126 free run: 15 years - Courtesy: P. Mukhopadhyay, IITM



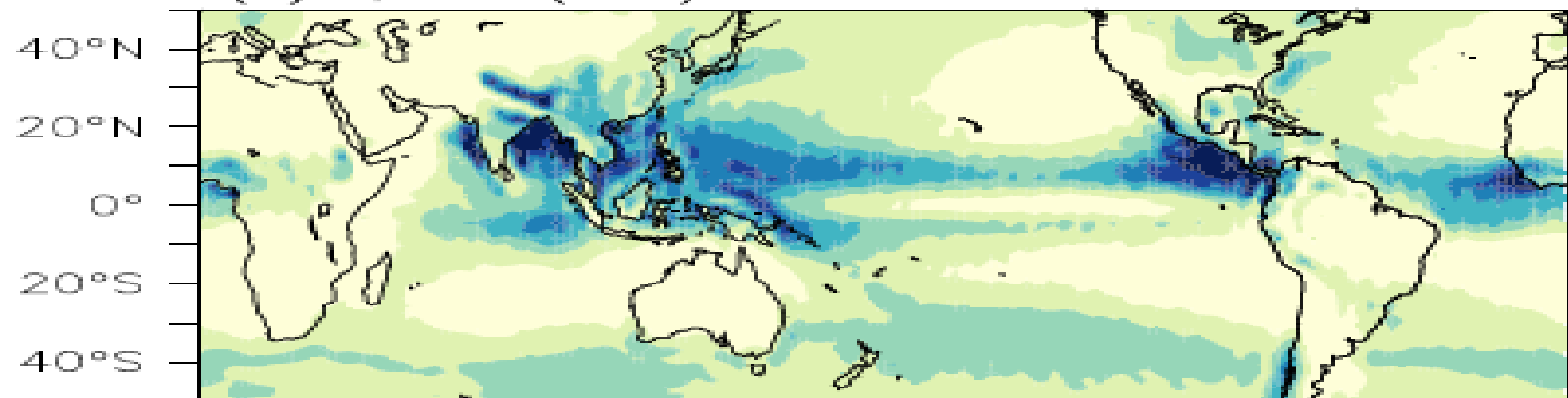
**(a) TRMM**



**(b) ESMv1 (T126)**



**(c) ESMv2 (T62)**



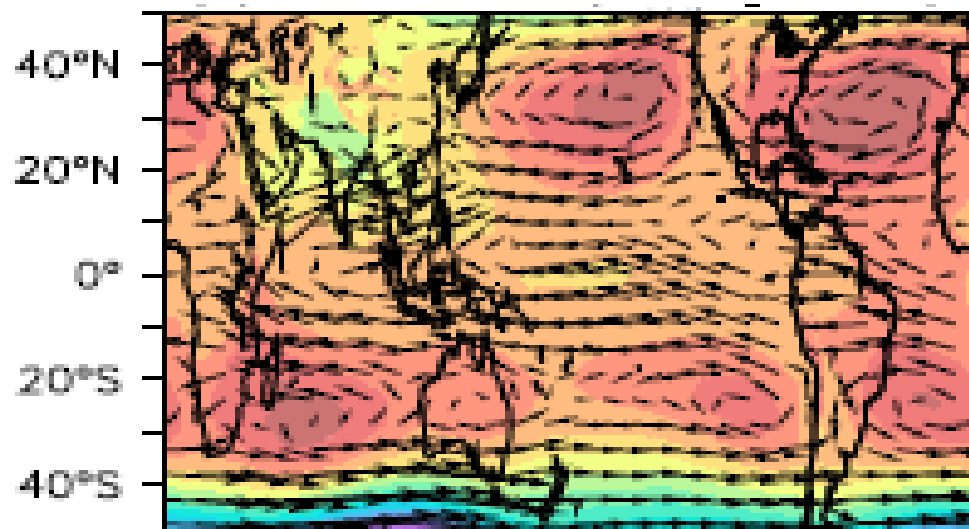
0° 80°E 160°E 120°W 40°W

12

8

4

0



Winds & Geopotential Height: 850 hPa

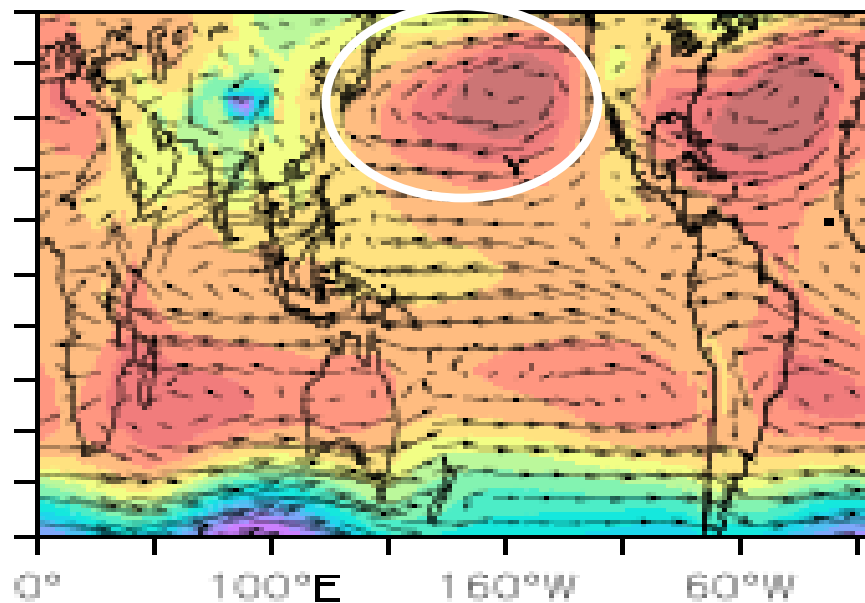
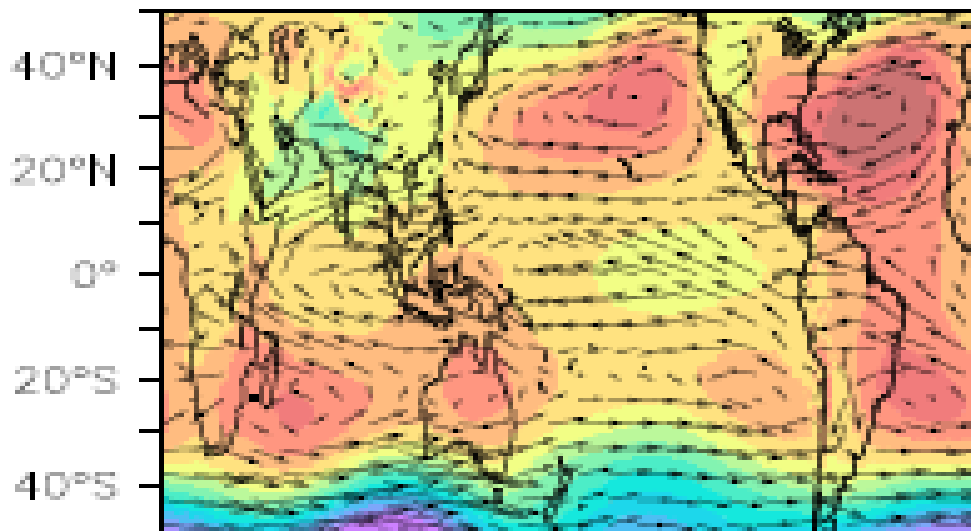
**JJAS**

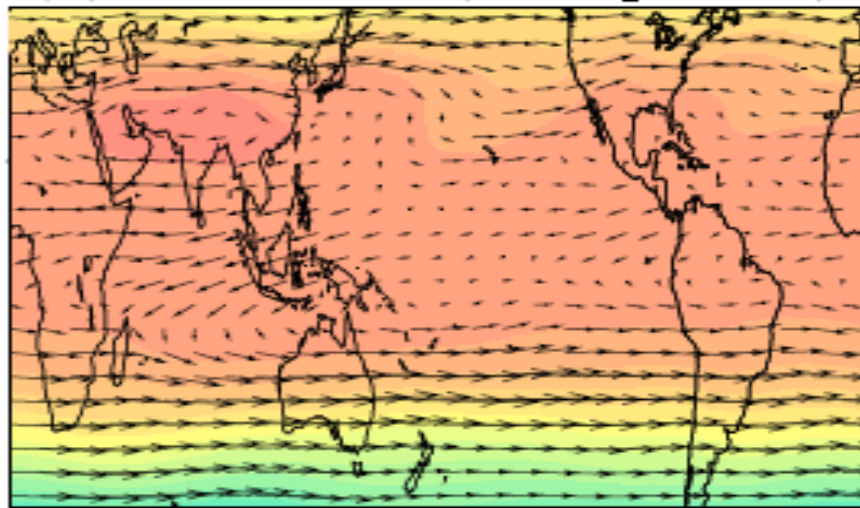
ERA-Interim

- Pacific sub-tropical anticyclone
- Easterly trade winds over Pacific

ESM-v1

ESM-v2

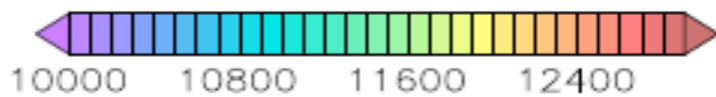




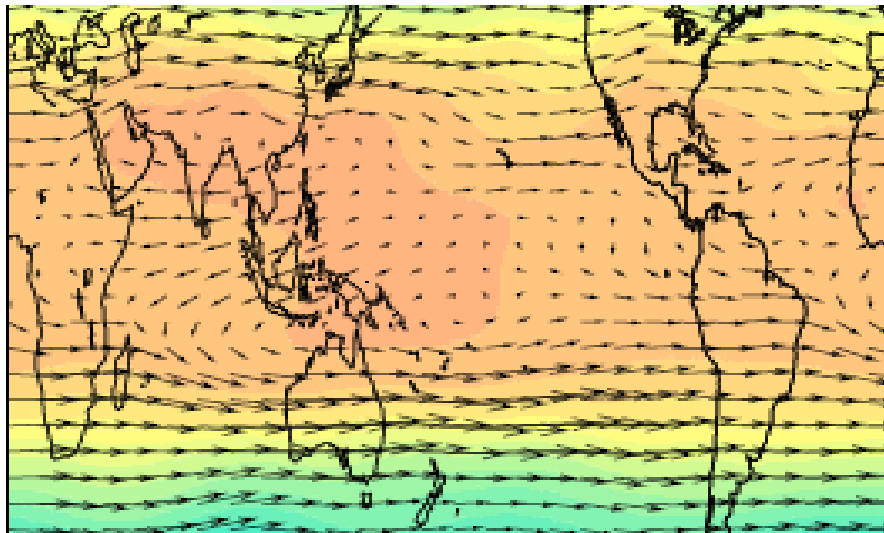
Winds & Geopotential Height: 200 hPa

**JJAS**

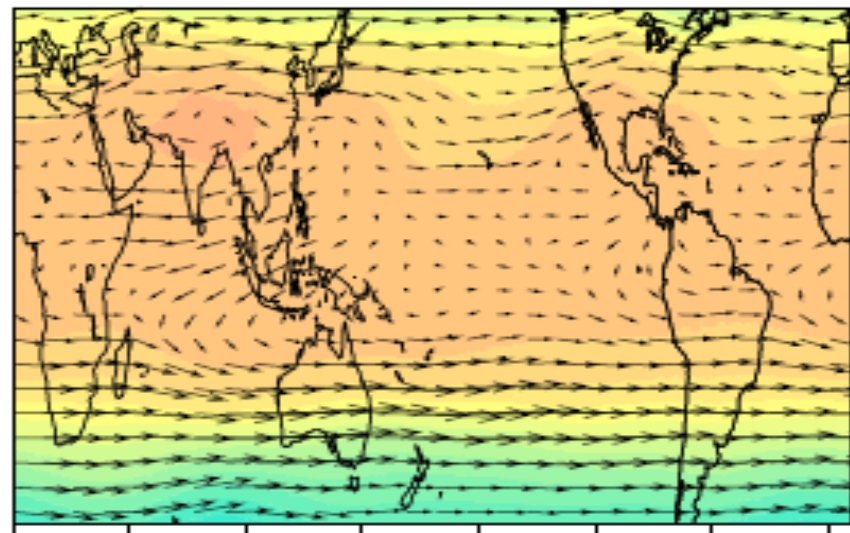
ERA-Interim



**ESM-v1**



**ESM-v2**

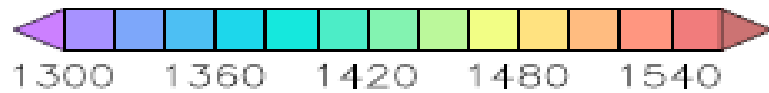
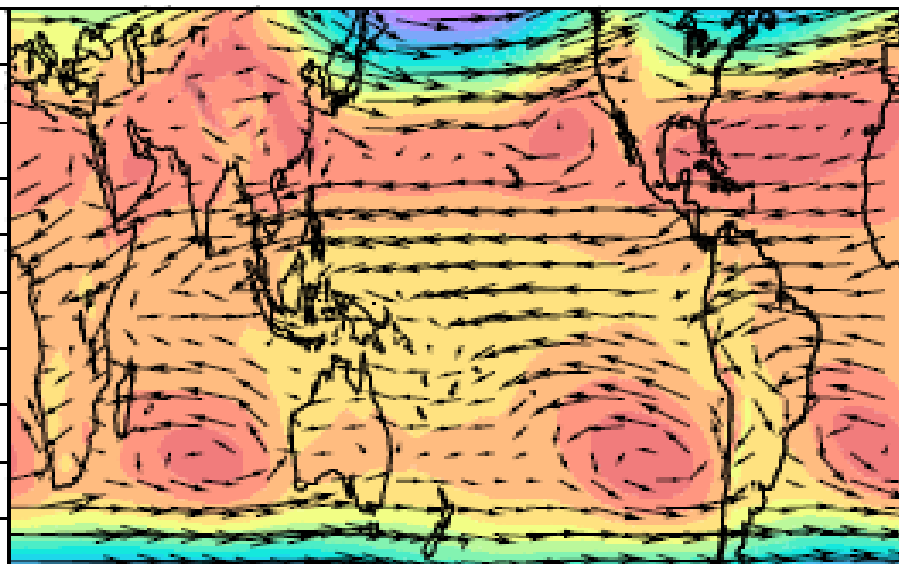




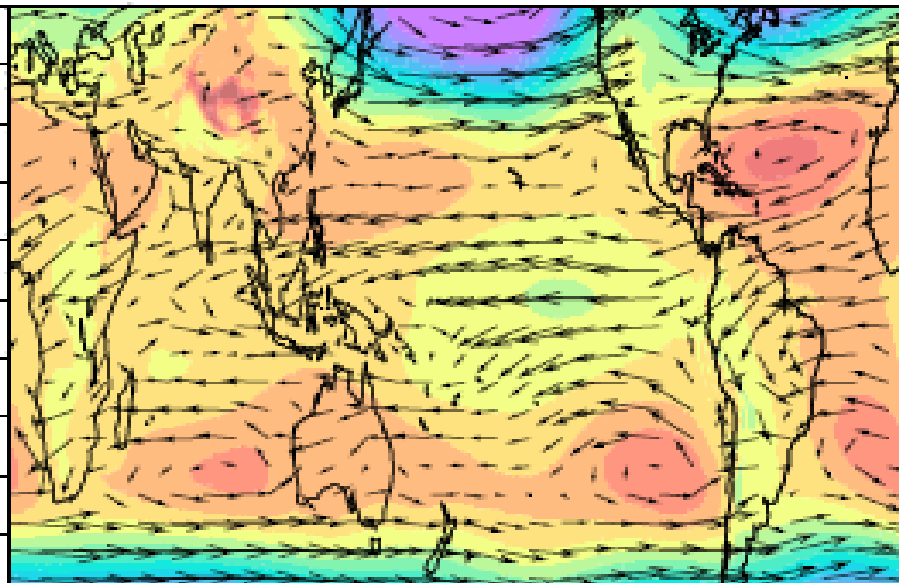
Winds & Geopotential Height: 850 hPa

**DJF**

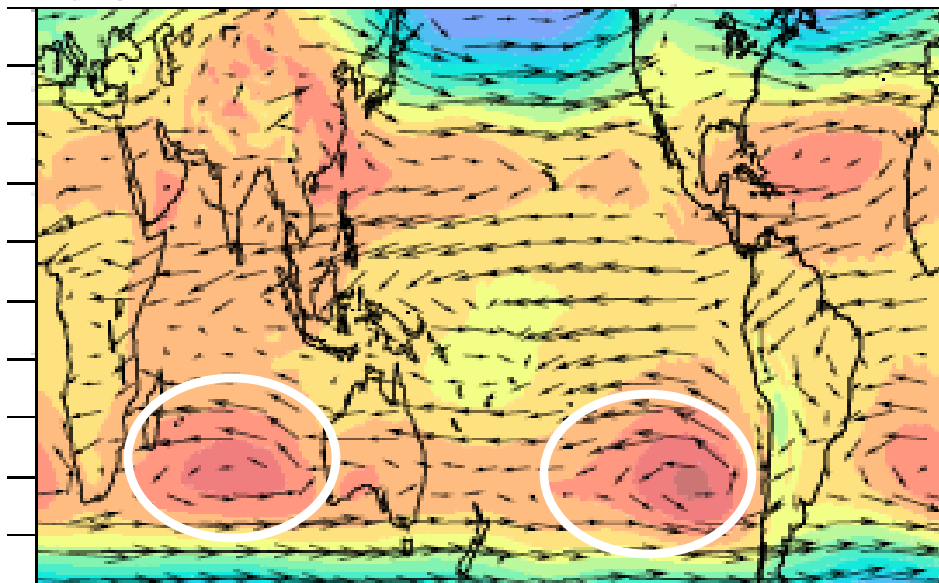
ERA-Interim



**ESM-v1**



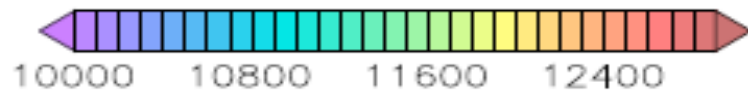
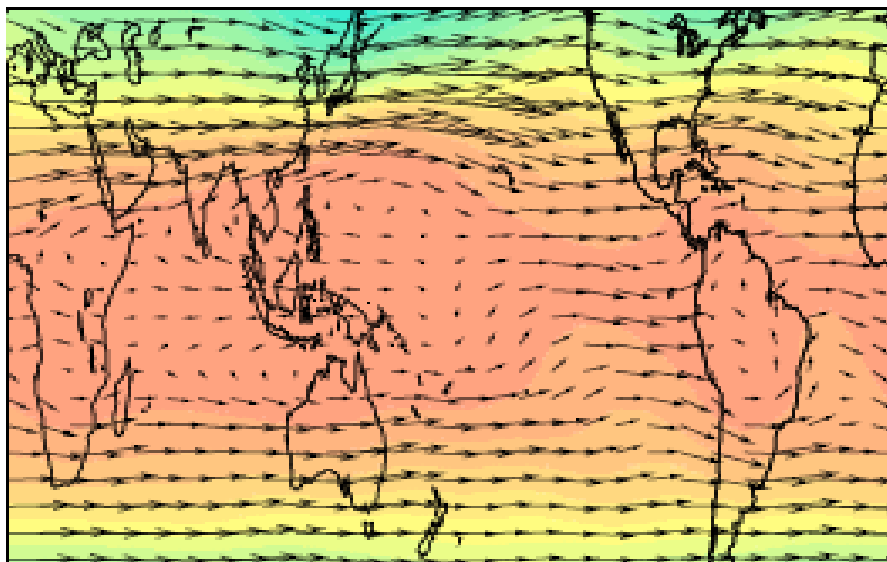
**ESM-v2**



Winds & Geopotential Height: 200 hPa

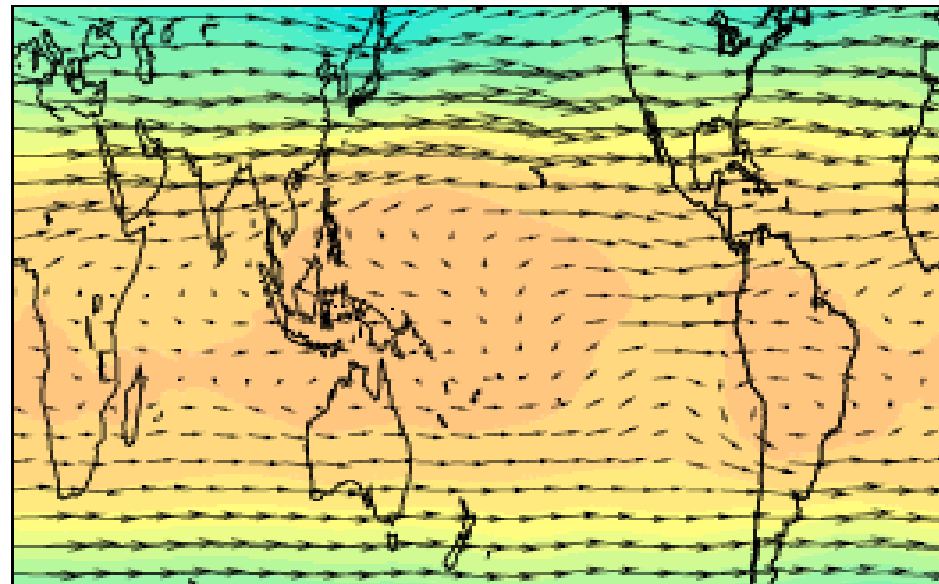
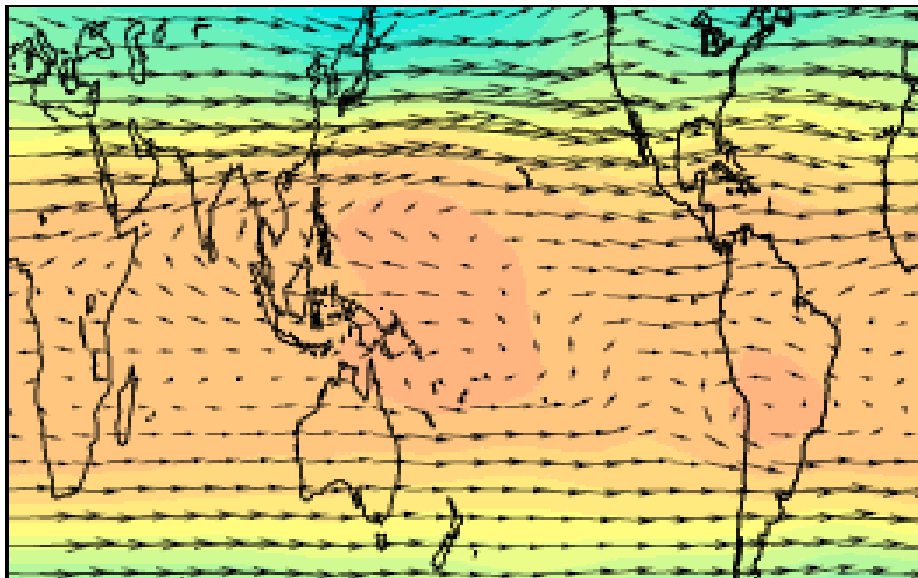
**DJF**

ERA-Interim



ESM-v1

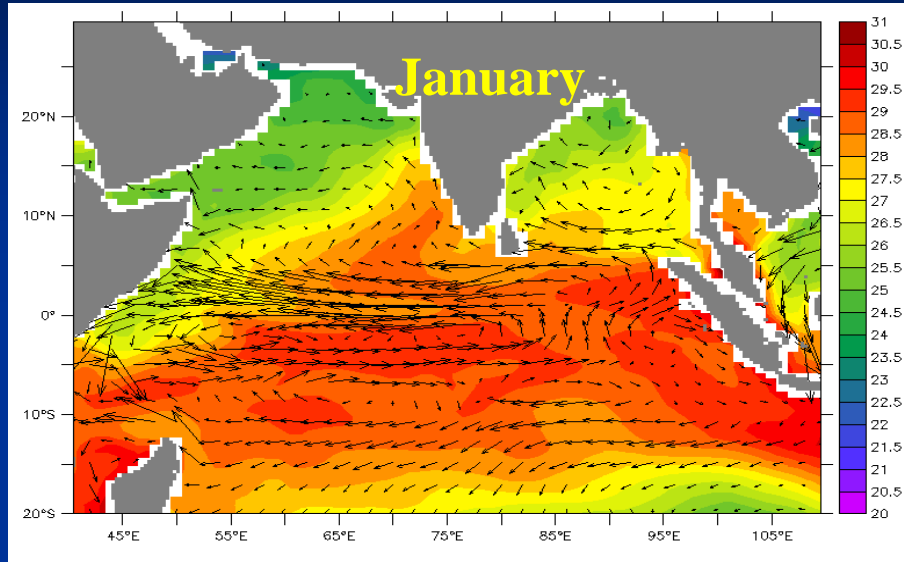
ESM-v2



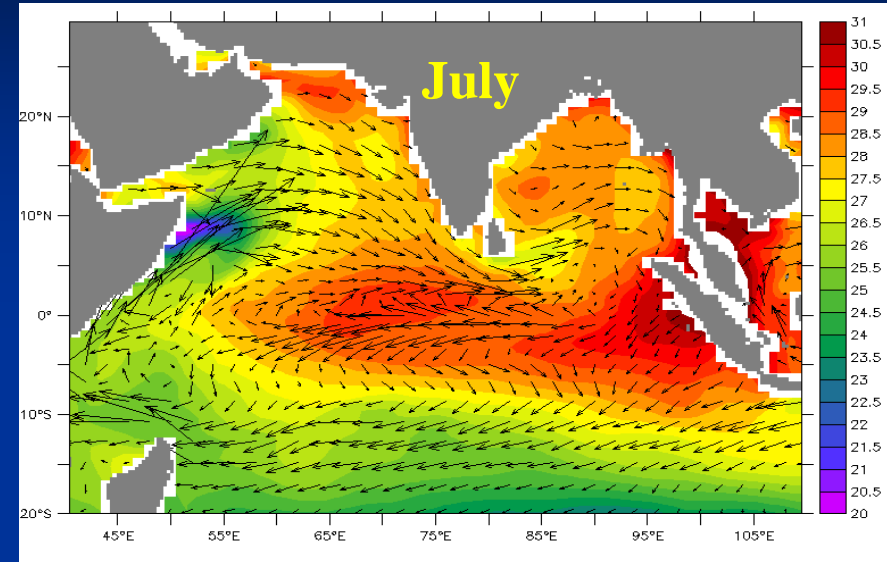
# MOM4p1 forced ocean simulation – 130 year spin up

## Physical and Biogeochemical Parameters for Tropical Indian Ocean

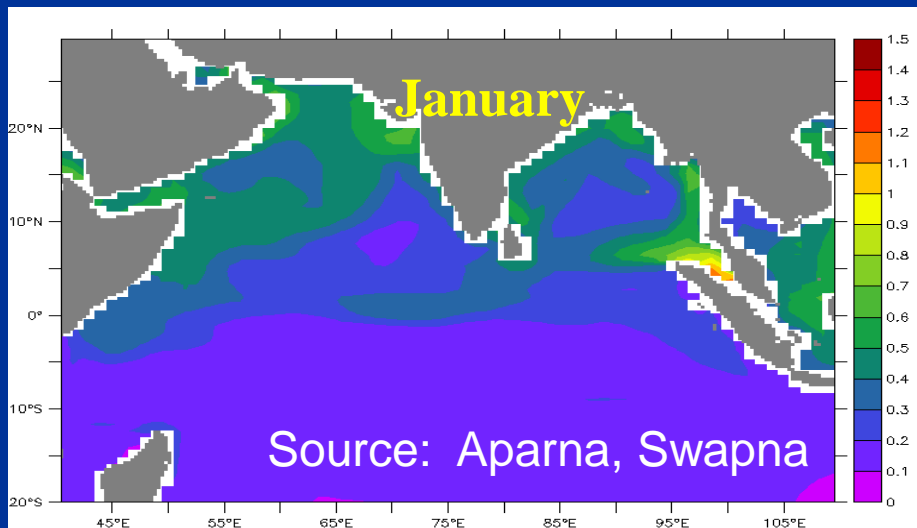
SST and currents)



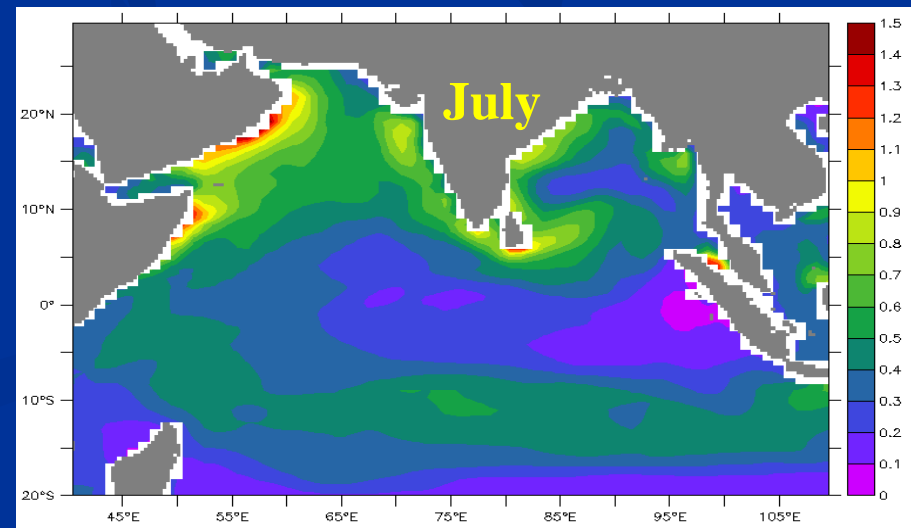
SST and currents)



Chlorophyll



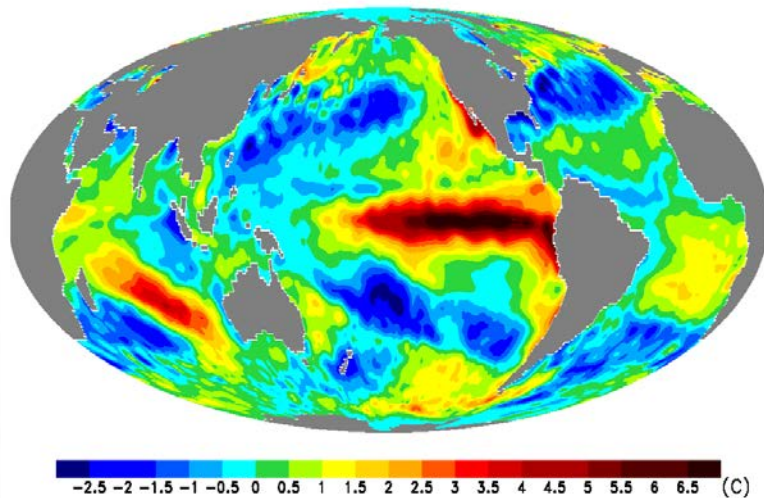
Chlorophyll



1997-98: Strongest El Niño ever recorded!

Sea Surface Temperature

Dec 1997 minus Dec 1998



In January 1998 (top right) the 1997-1998 El Nino event was at its height. Because of the weakness of the trade winds at this time, the upwelling of nutrient-rich water was suppressed in the equatorial Pacific. The absence of a green band along the equator in this image is indicative of relatively low chlorophyll concentrations there.

By July 1998 (bottom right) the trade winds had strengthened and equatorial upwelling had resumed giving rise to widespread phytoplankton blooms in the equatorial belt

(Ref: Wallace and Hobbs, 2006)

SeaWiFS Captures El Niño - La Niña Transitions in the Equatorial Pacific

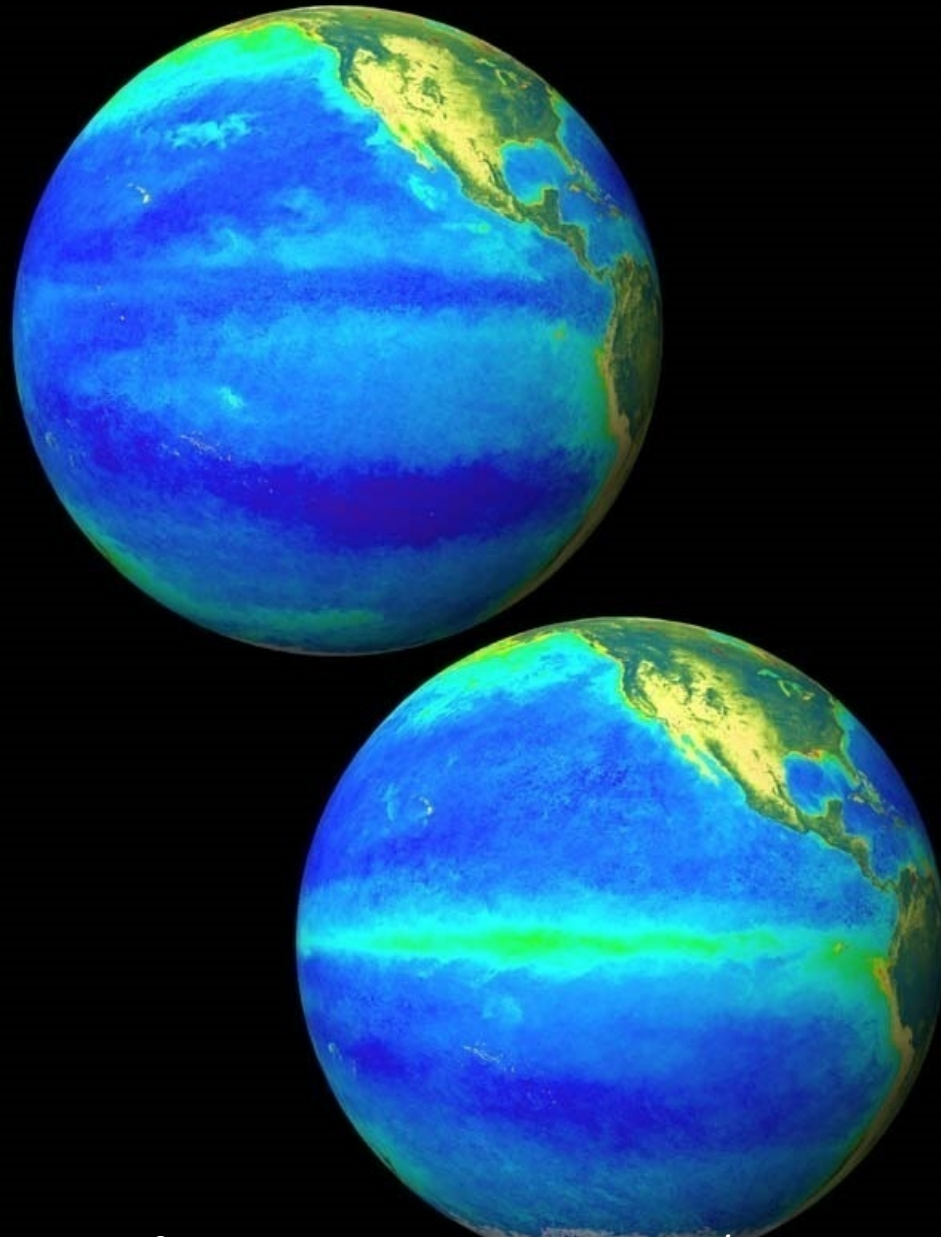
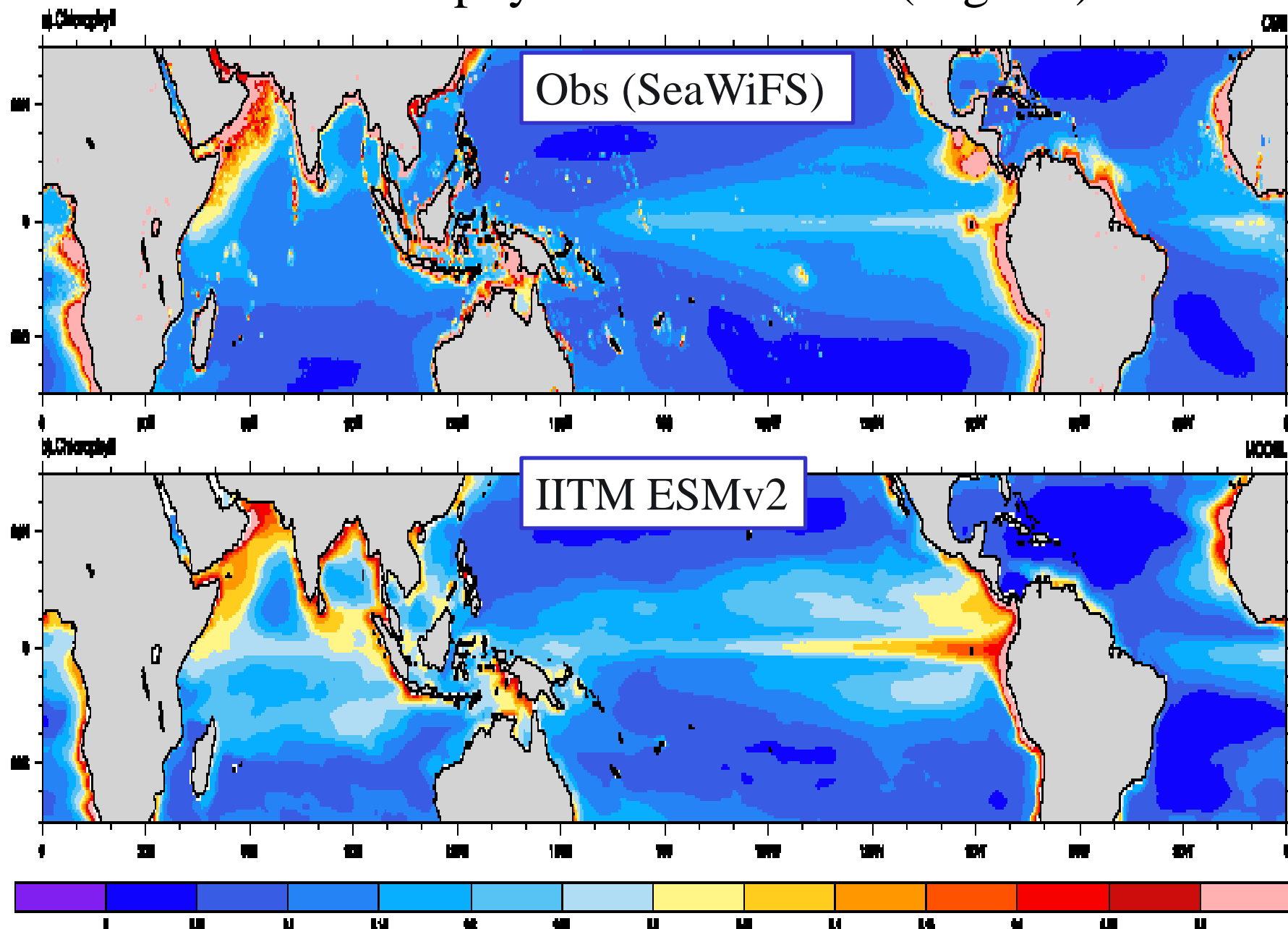


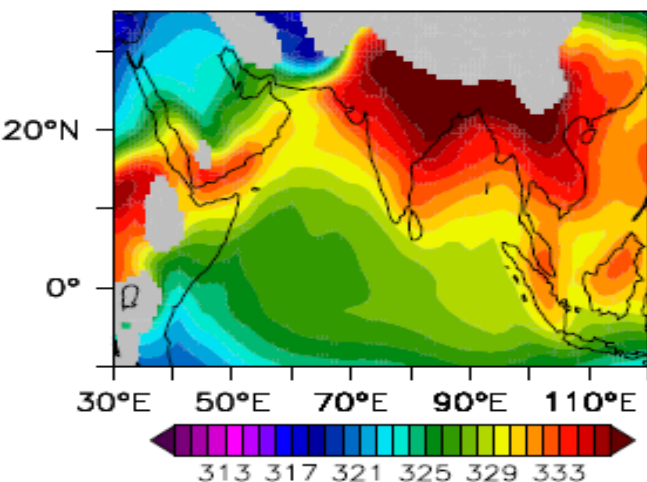
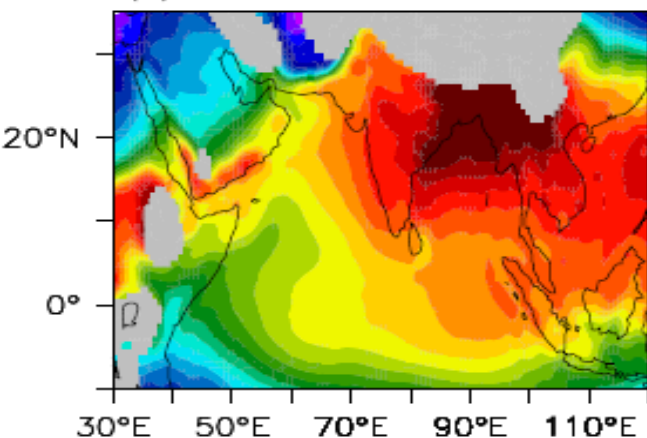
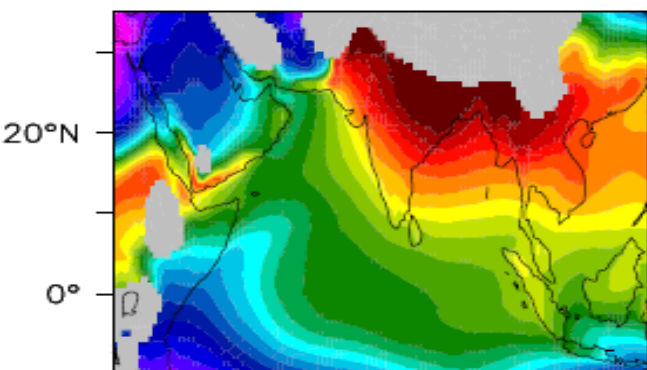
Image from SeaWIFS Project, NASA / GSFC

# Chlorophyll Concentration ( $\text{Mg m}^{-3}$ )



Courtesy: Sandeep, CCCR

## Moist Static Energy

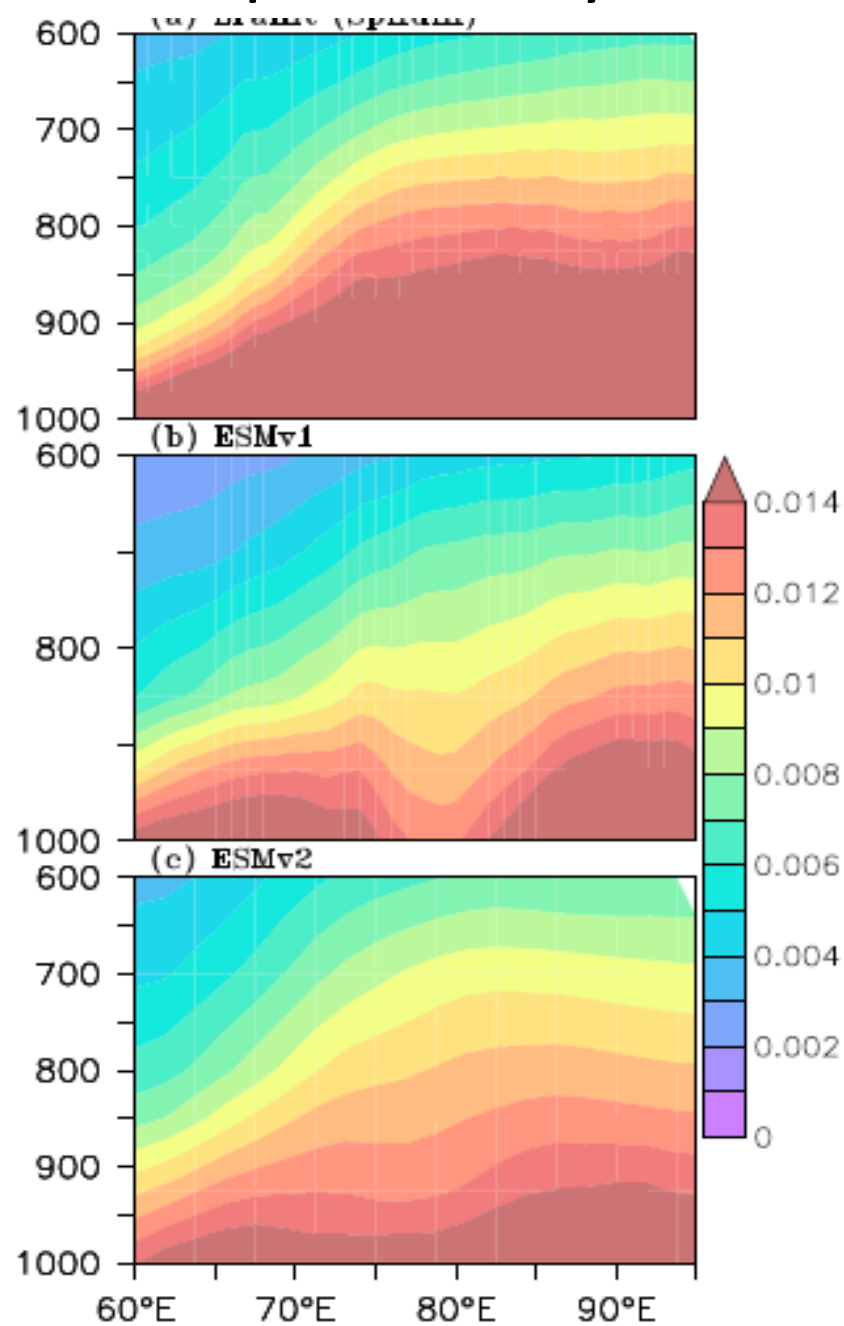


ERA  
Interim

ESMv1

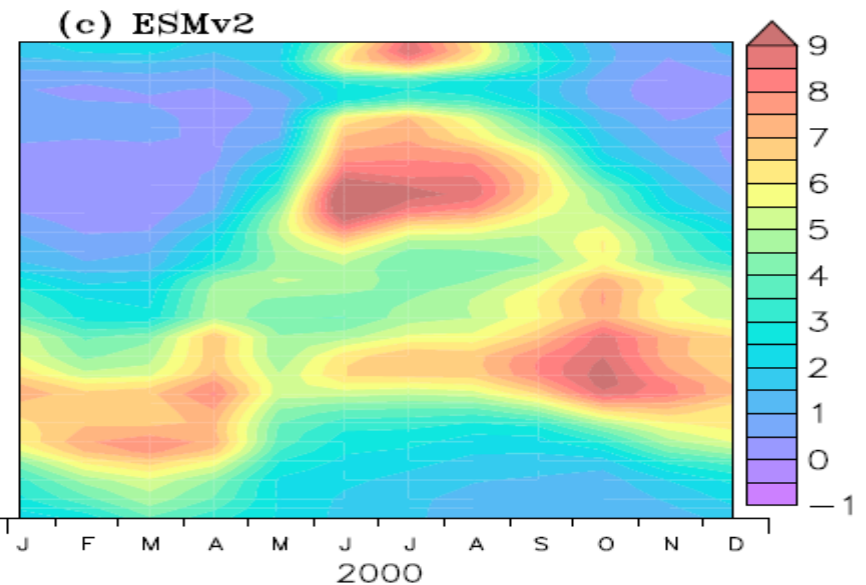
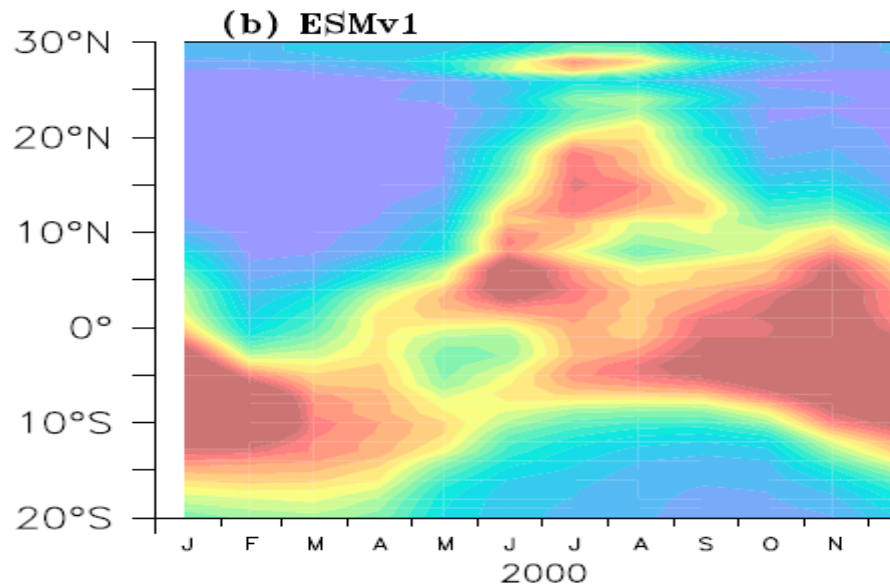
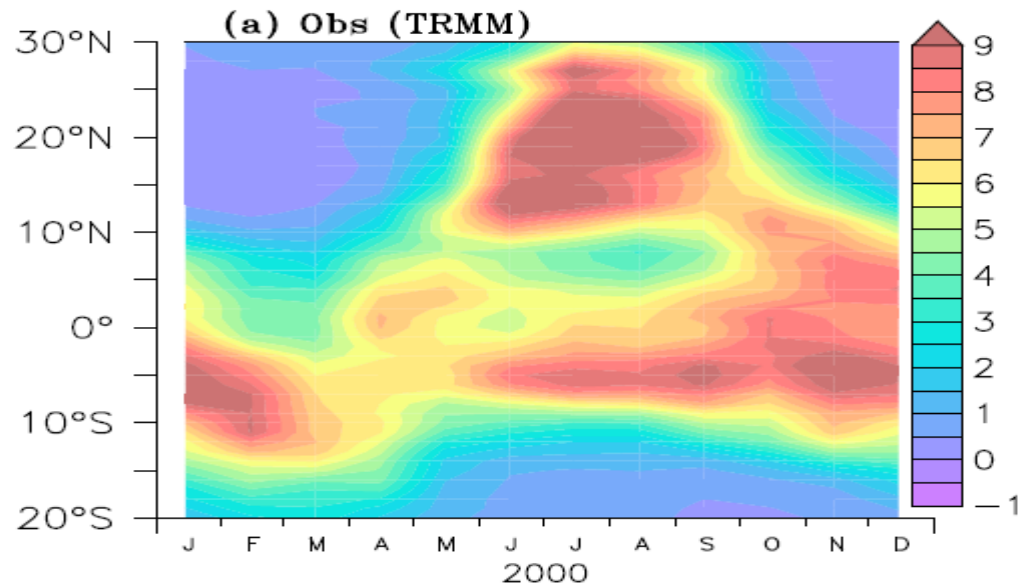
ESMv2

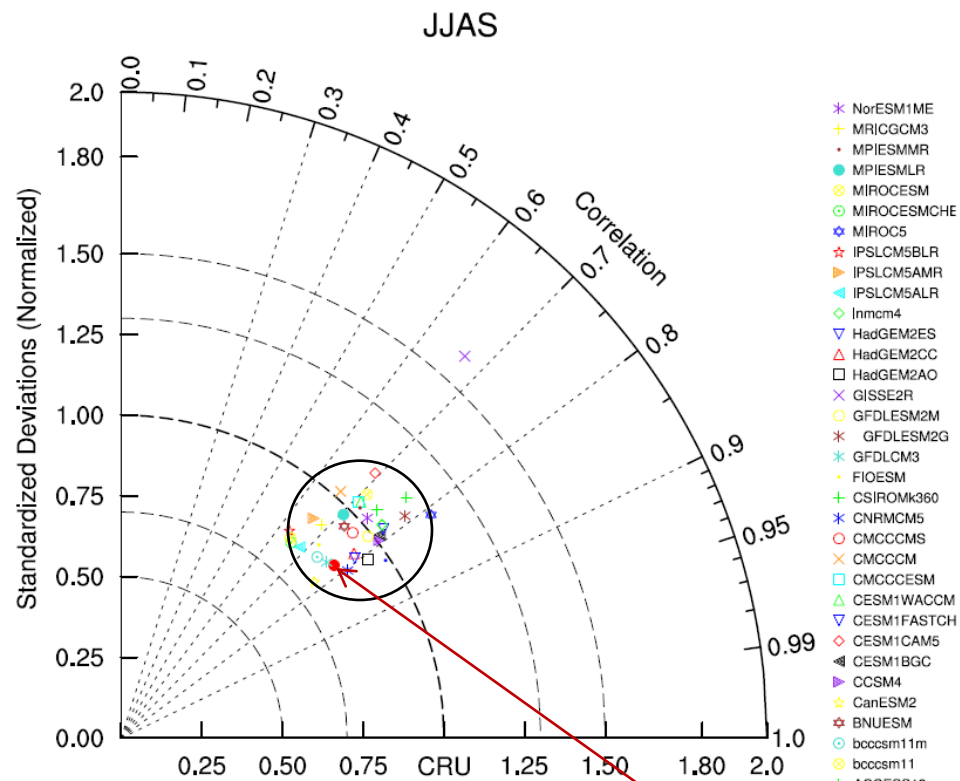
## Specific Humidity





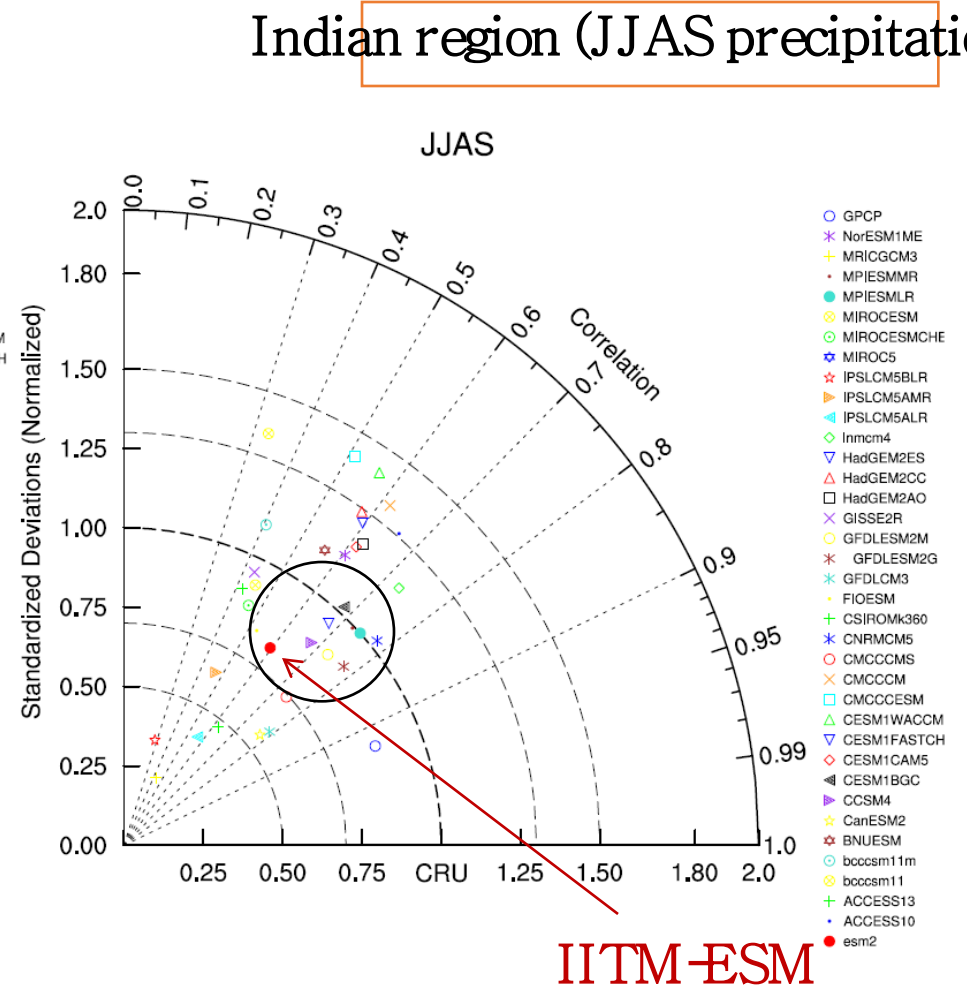
# Precipitation Seasonal Cycle (70E-90E)



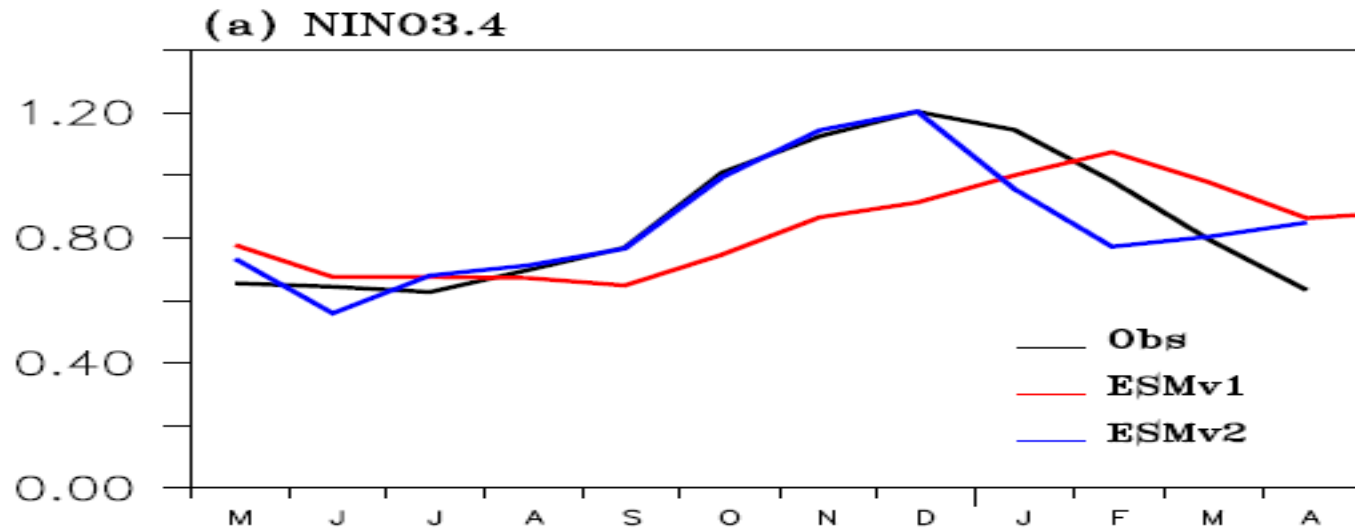


Asian region (JJAS precipitation)

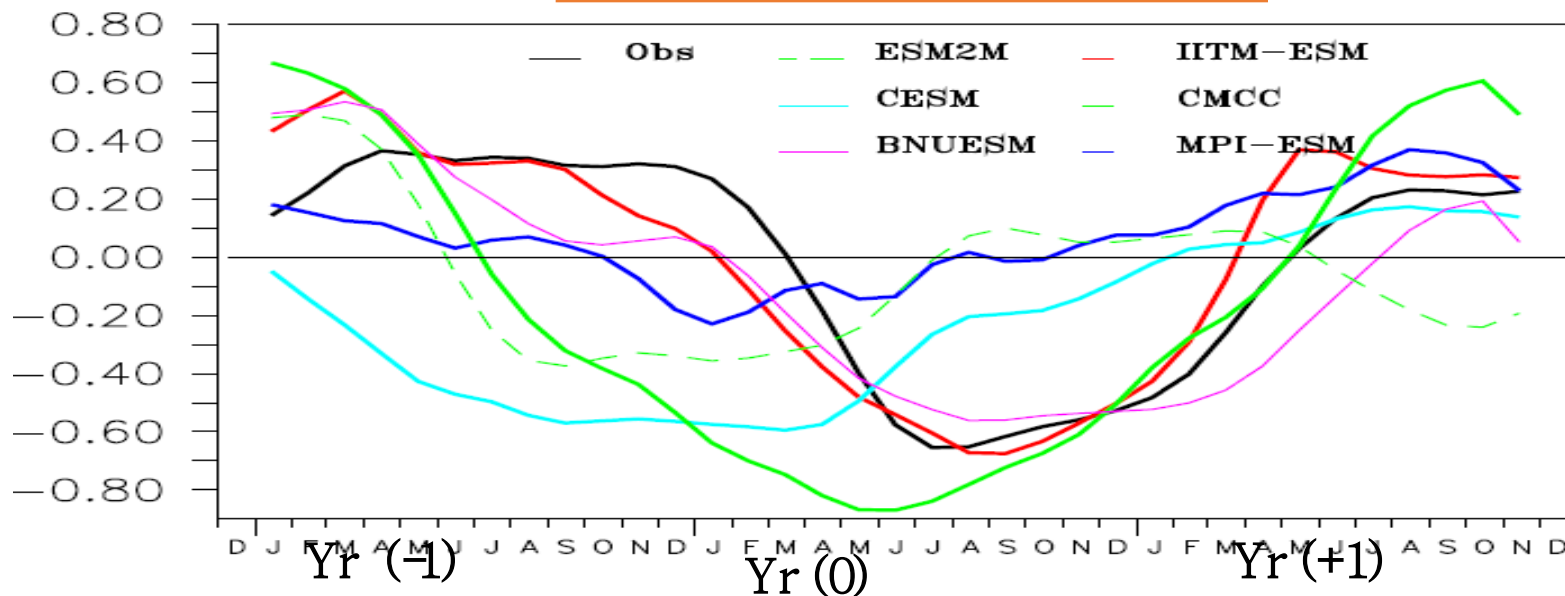
IITM-ESM



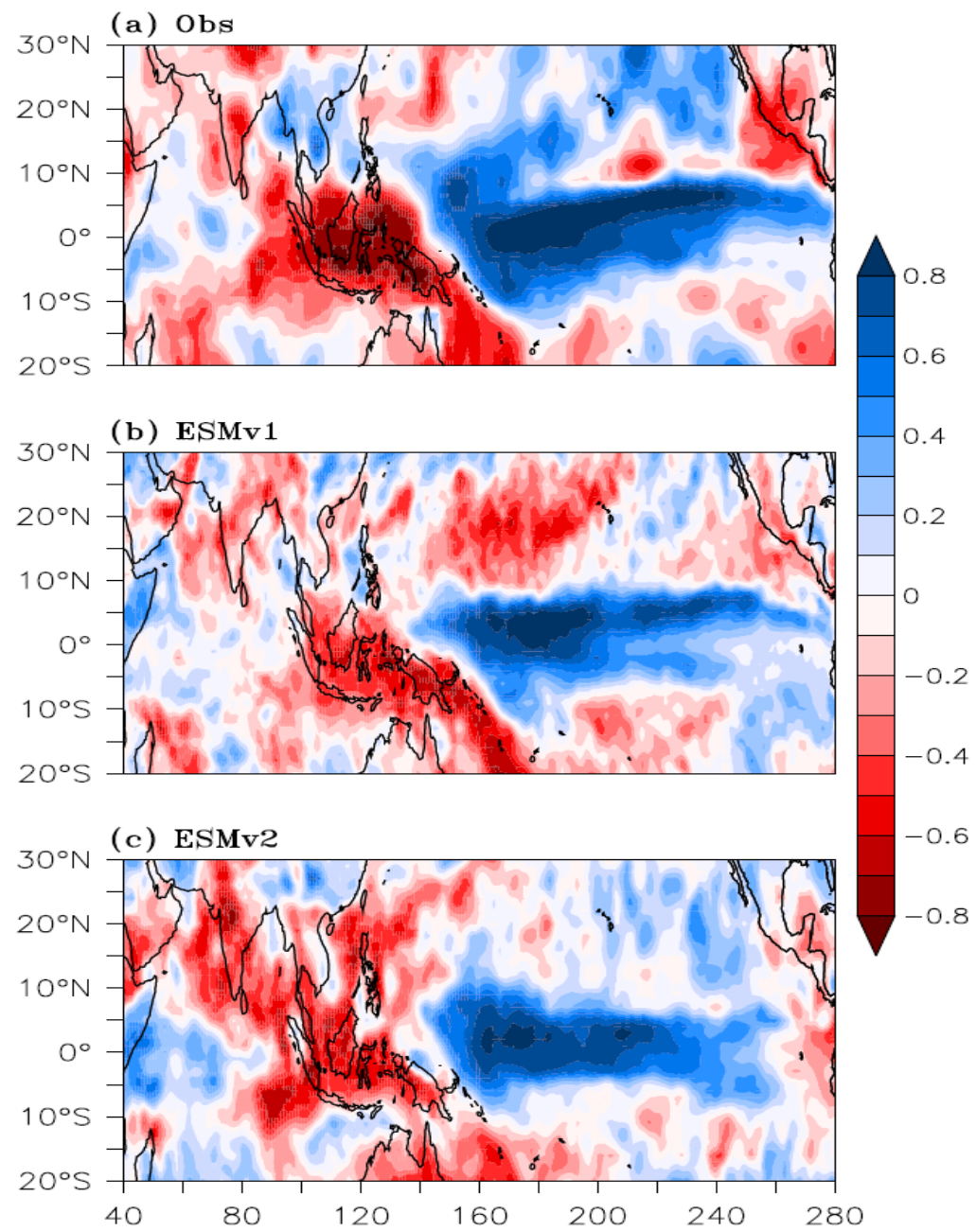
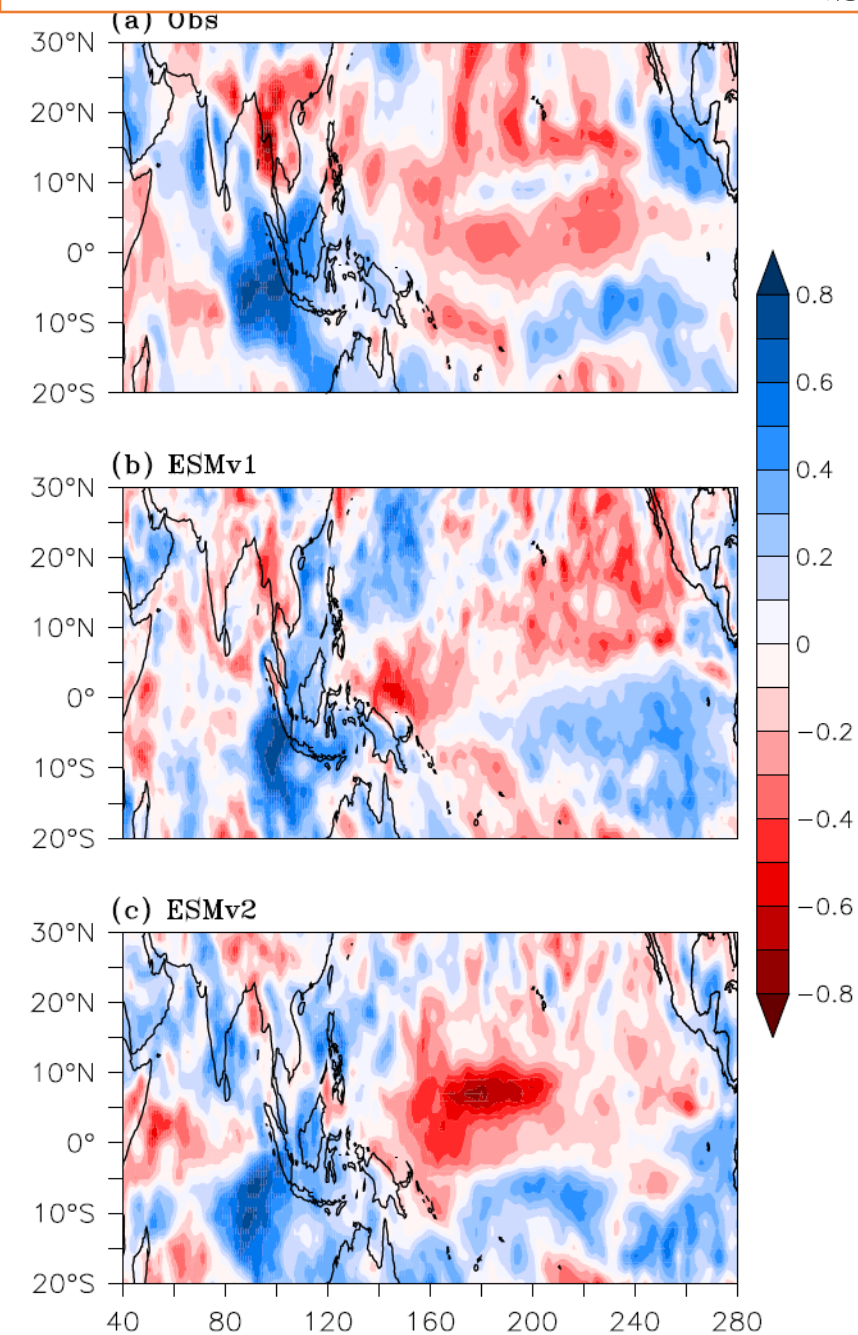
# Seasonal variability of NINO3.4



## ENSO-Monsoon teleconnection

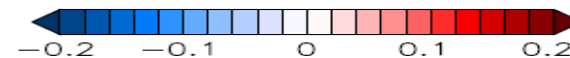
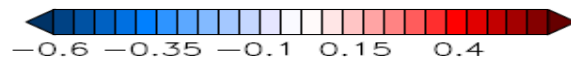
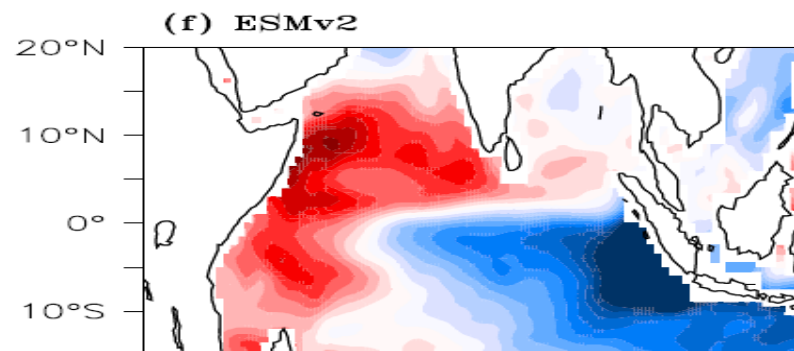
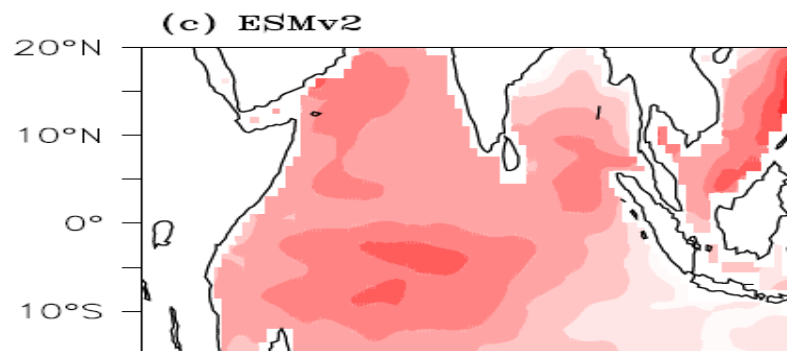
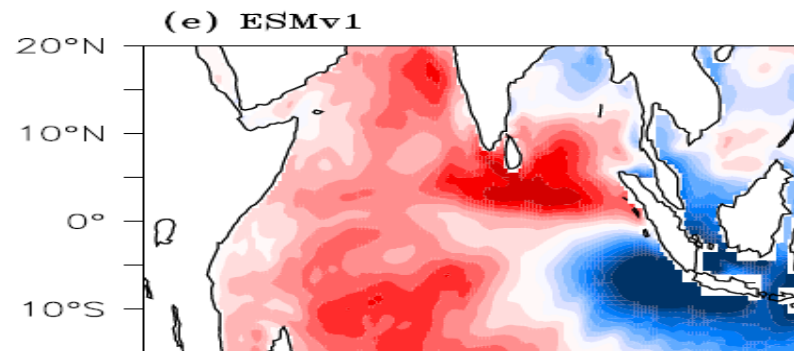
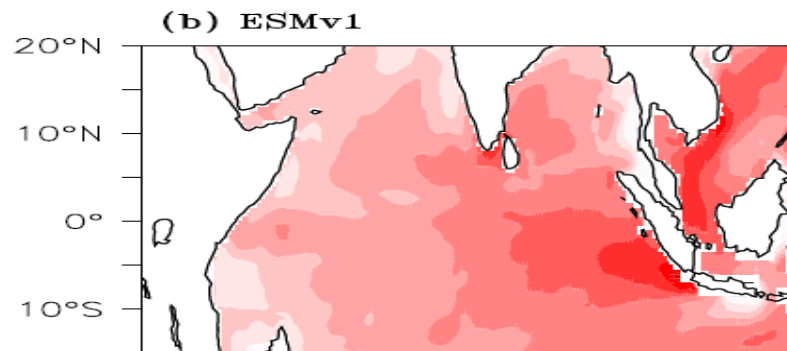
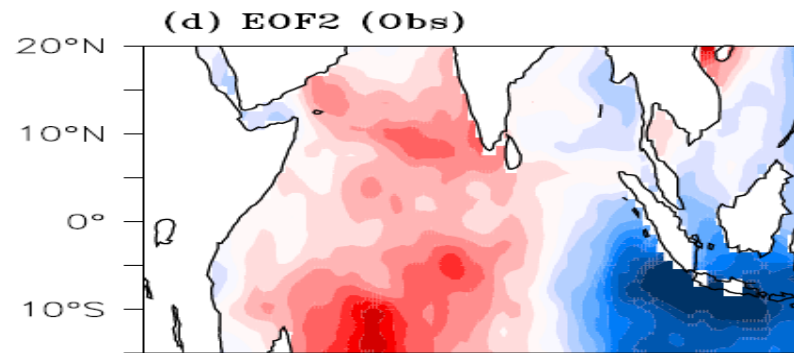
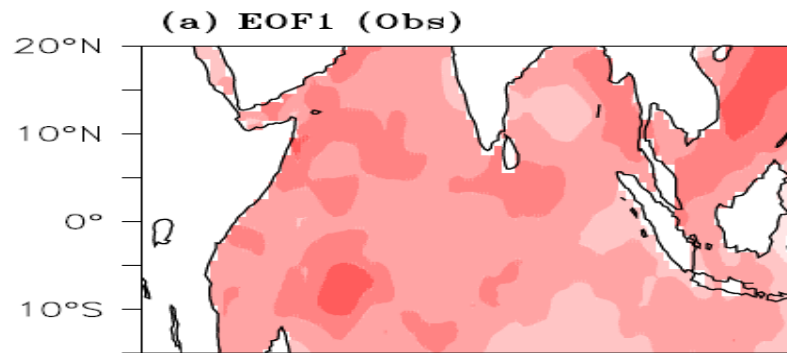


# Monsoon teleconnection (SST EIO vs JJAS precipitation) ENSO Monsoon teleconnection (NINO3.4 vs JJAS precipitation)





# Tropical Indian Ocean Variability (IOBM & IOD)



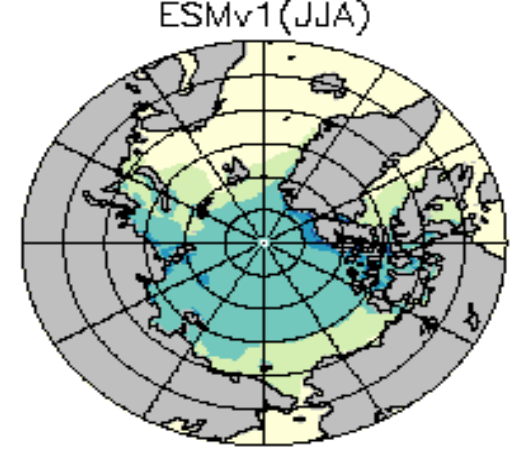
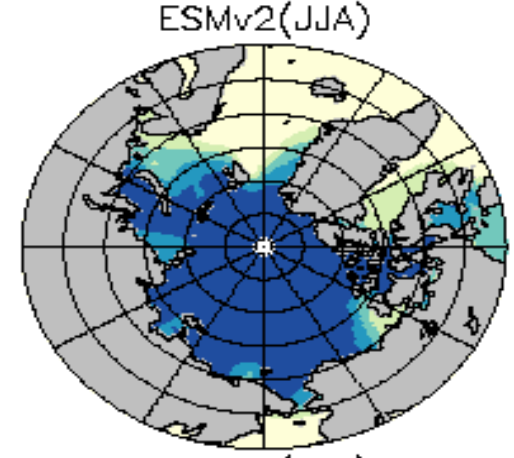
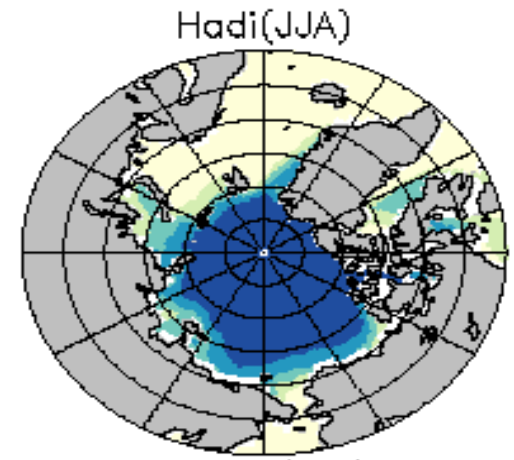
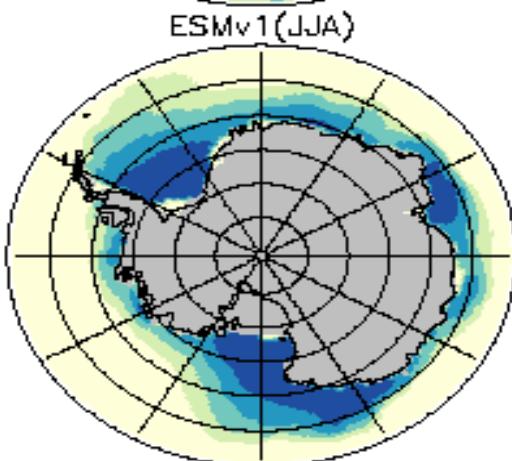
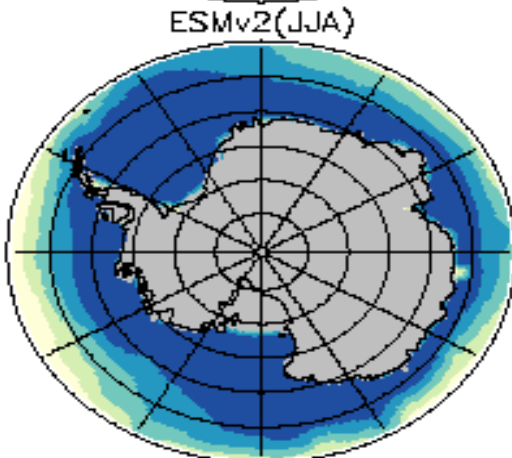
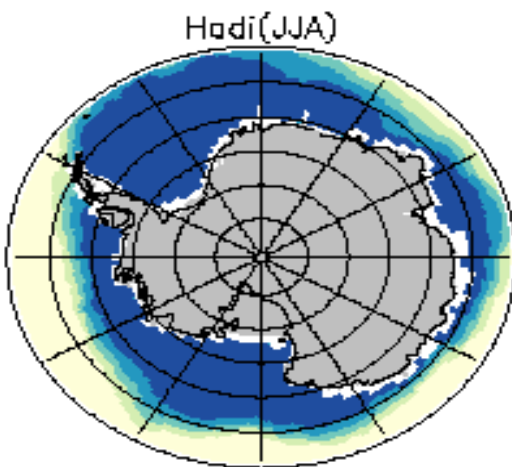
# Sea-Ice concentration

Obs

ESMv2

Improved simulation of  
NH sea-ice during JJA

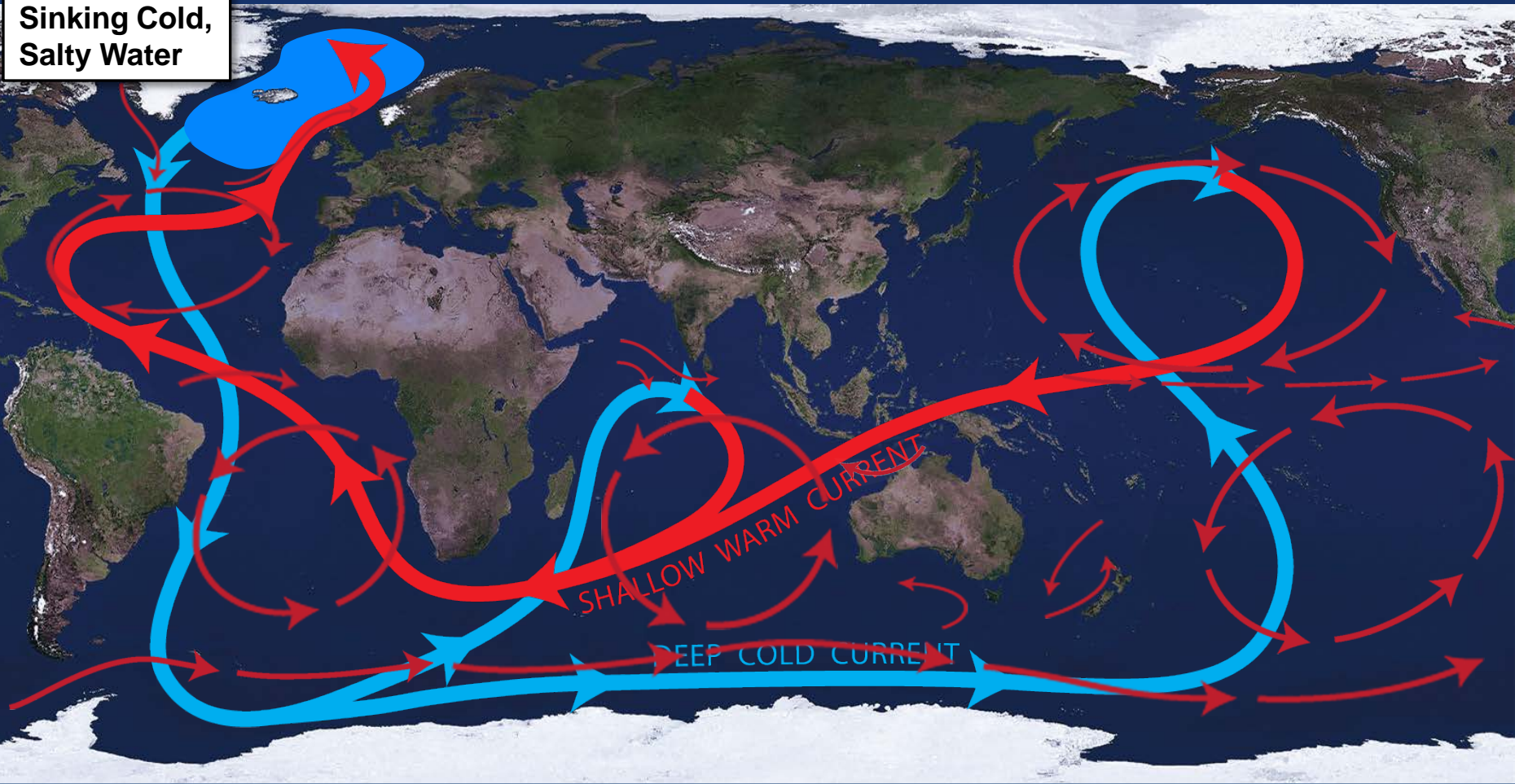
ESMv1

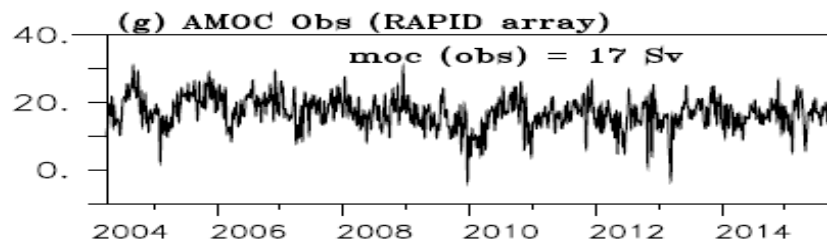
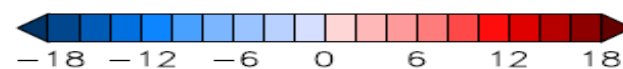
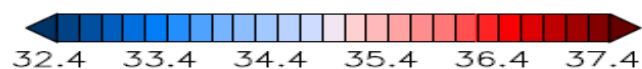
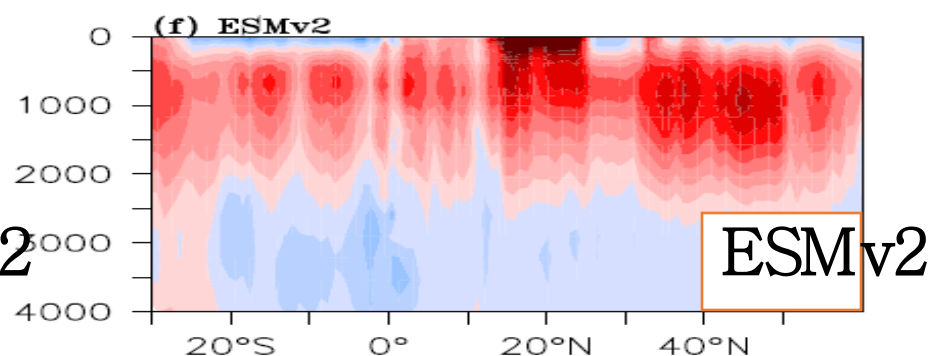
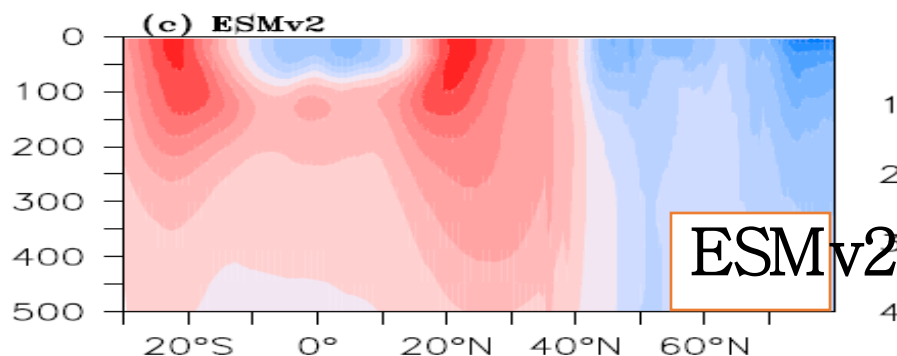
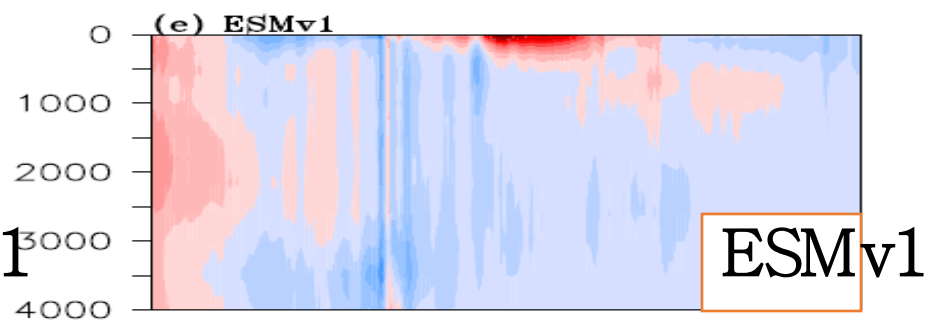
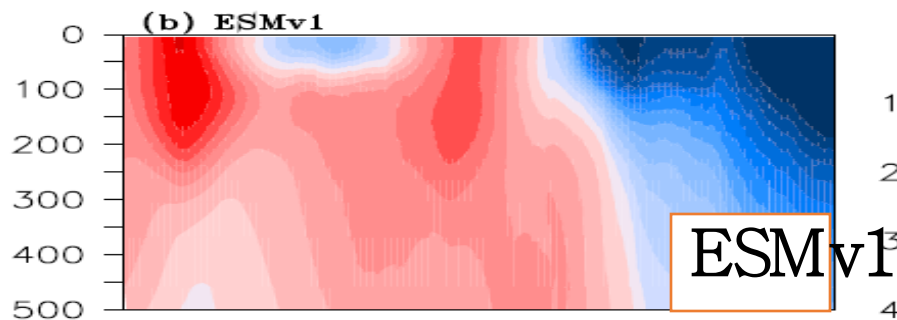
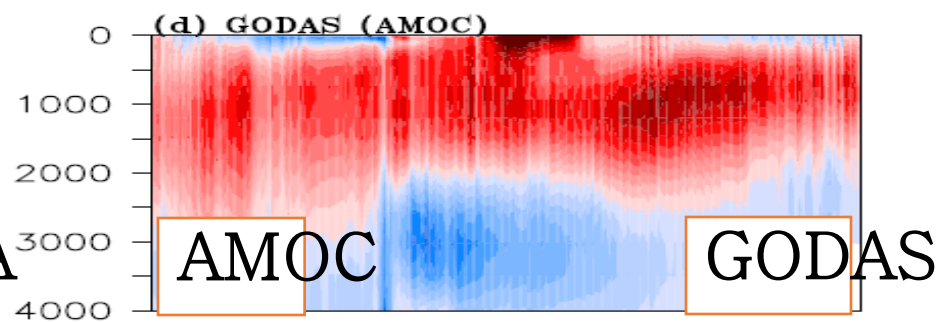
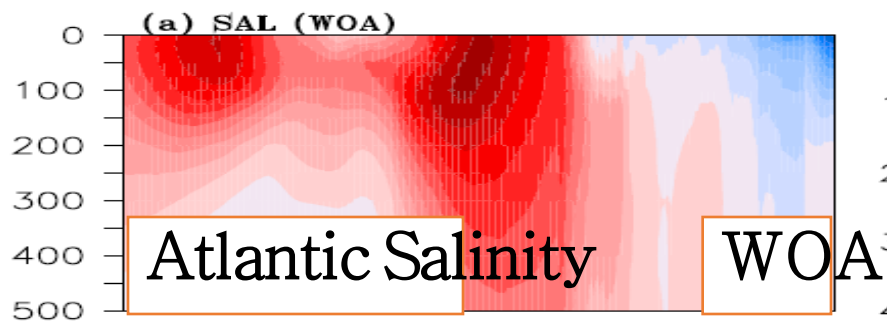




# Thermohaline Circulation (THC) Global Conveyor Belt

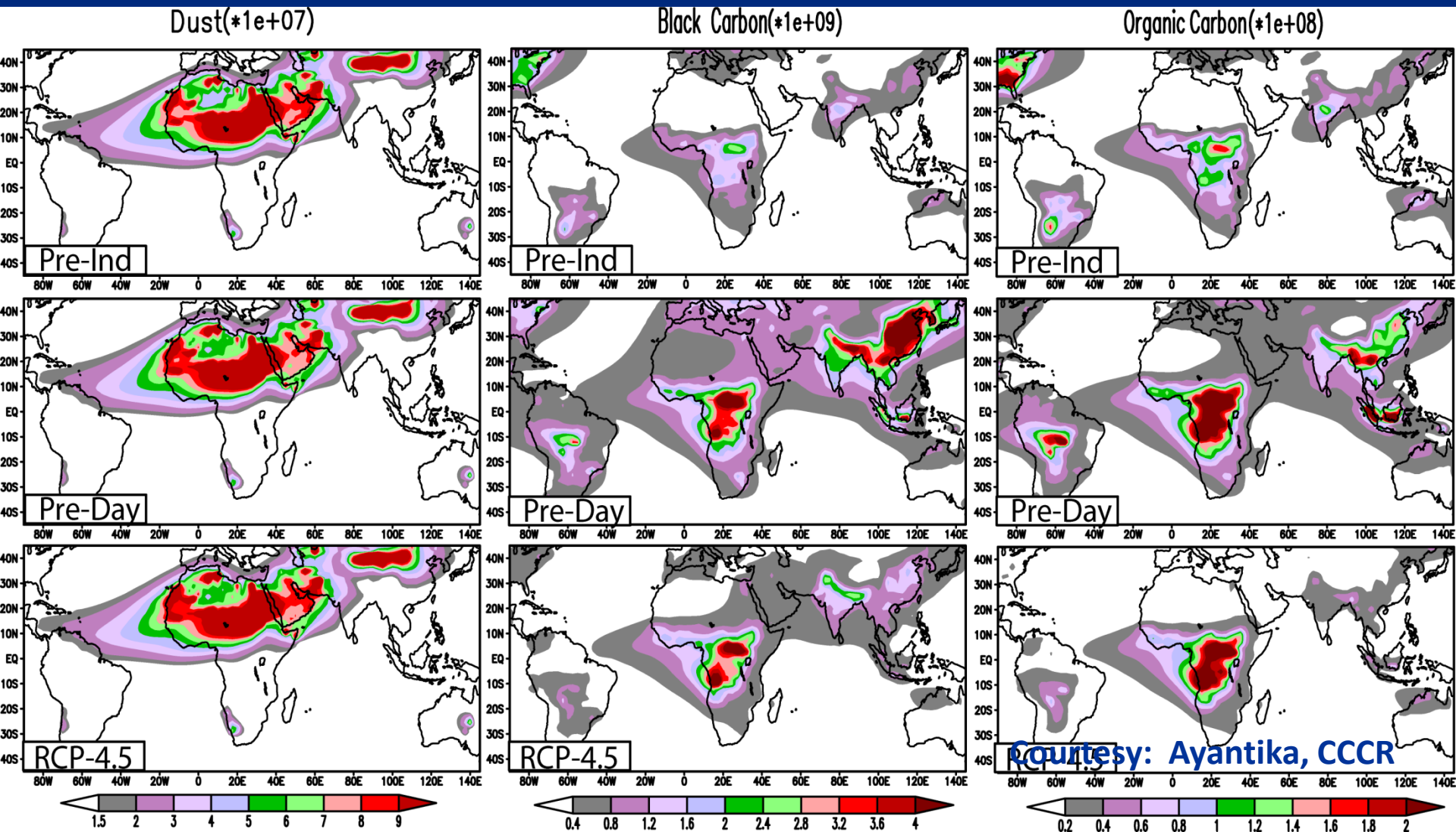
Sinking Cold,  
Salty Water





# Prescribed time-varying aerosol distributions in IITM-ESM from CMIP

Total column aerosol content provided by CMIP5 for **Pre-industrial period (1850 -1879)**, **Present day (1980 – 2009)** and **RCP4.5 (2089 – 2109)**. The units of the aerosol fields (Dust, BC and OC) are kg/kg. Information about other aerosol fields (eg. Sulphate, Sea Salt and Secondary Organic Carbon) is also available from CMIP)





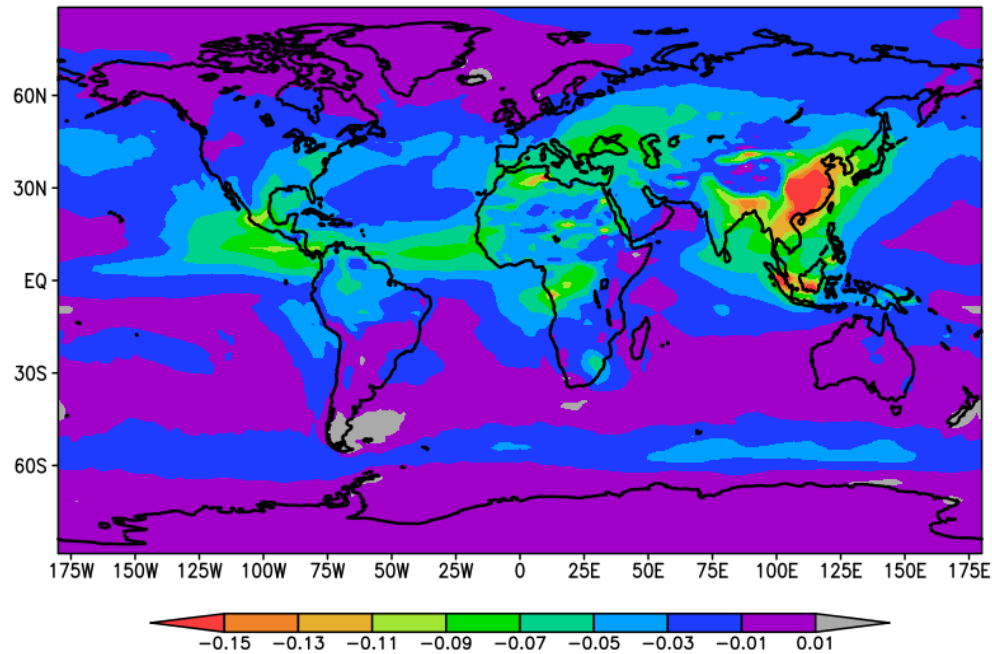
# CMIP Aerosols

- Aerosol concentration for the following species: SO<sub>4</sub>, black carbon, organic carbon, secondary organic aerosols, dust and sea-salt
- Wavelength resolved complex refractive indices and estimates of the aerosol size distributions (geometric mean, geometric std.dev) for different relative humidity are supplied to a Mie code [Mischenko et al., 1999, 2002] for optical property calculations
- Mie parameters averaged over size distributions are pre-tabulated as a function of RH, and then used to calculate aerosol optical properties e.g. AOD for a given time and grid cell

$$AOD = \sum_{i=1}^N \frac{3}{4} \frac{\bar{Q}_{ext,i}(RH)}{\rho_i r_{eff,i}(RH)} M_i \quad (\text{Curci et al 2012})$$

- Aerosol Optical Depth, Single Scattering Albedo, Asymmetry Parameter calculated for ESM SW and LW bands
- The aerosol optical properties are used as input in ESM RRT radiation calculation

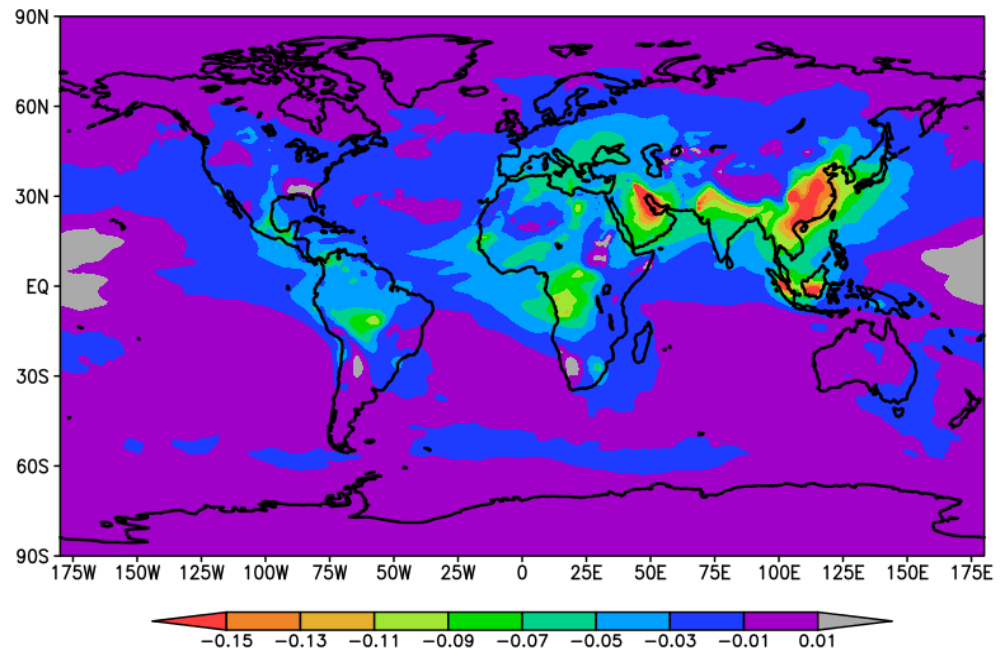
Calculated for IITM ESM



AOD difference map

Preindustrial (1870)- Present (2005)

CESM (CAM5)



Courtesy: Ayantika, CCCR

# Time-varying aerosol distributions in IITM ESM from CMIP

(Courtesy: Ayantika Dey Choudhury; Data source: Stefan Kinne, Bjorn Stevens, Max Planck)

To understand the impact of natural tropospheric aerosols on climate: as observed in IITM-ESM

Sensitivity experiments: 50yrs model integration

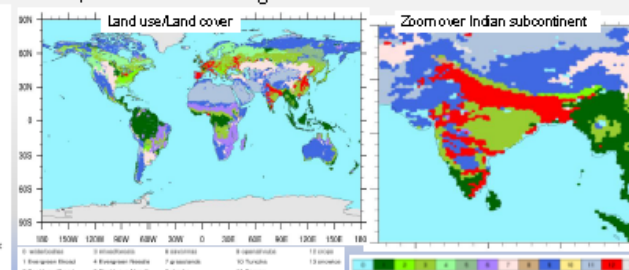
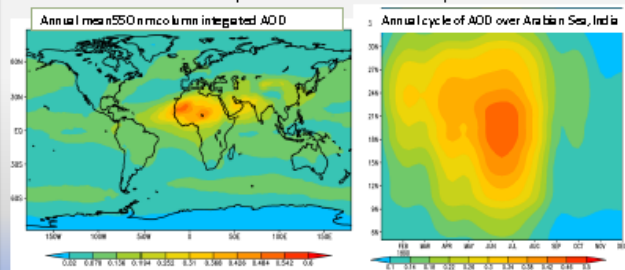
1) Pre-Industrial aerosols 2) No aerosols (Clean Environment)

## PI Aerosols

- MAC-V2 aerosol data ( $1^\circ \times 1^\circ$  AOD, SSA, ASY and 20 layer EXT) (Max Planck Institute Aerosol Climatology V2, update of MAC-v1, Kinne et al 2013)
- The 3D (T62L64) AOD, SSA, ASY for IITM ESM-V2 is constructed by weighting the aerosol properties with the vertical profile of aerosol extinction data.
- Natural aerosols for 1850 are representative of Pre-Industrial era and implemented on monthly timescales

## PI Land use/ Land cover

- UNH transient historical crop and pasture dataset ( $.25^\circ \times .25^\circ$ ) (Hurtt et al. (2015) for the period 1870-2015)
- Crop area in IITMESM-v2 set to UNH data. Reduction of anthropogenic area implies proportional increase in all natural vegetation types for the cell. The most abundant vegetation type is set as dominant for that cell
- This method is used to derive the Pre-Industrial land use/land cover at T62 grid



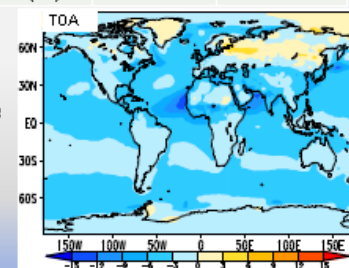
Aerosol Direct Radiative Forcing

## Aerosol Radiative Forcing ( $\Delta F = F_{aer} - F_{clean}$ )

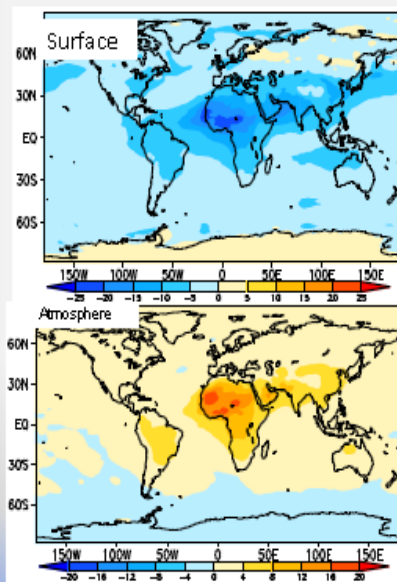
$F_{aer}$  is the net flux in the presence of natural aerosols  
 $F_{clean}$  is the net flux reference atmospheric condition (clean)

ARF		Clear Sky ( $W/m^2$ )	Total Sky ( $W/m^2$ )
TOA	Net Total Radiation ( $S_{NH} - U_{NH}$ )	-2.63	-0.45
	Net Solar Radiation ( $S_{NH}$ )	-3.49	-0.89
ATM	Net Total Radiation ( $S_{NH} - U_{NH}$ )	2.47	2.39
	Net Solar Radiation ( $S_{NH}$ )	1.65	1.52
SRC	Net Total Radiation ( $S_{NH} - U_{NH}$ )	-5.11	-2.84
	Net Solar Radiation ( $S_{NH}$ )	-5.15	-2.39

Solar Radiative forcing for clear sky



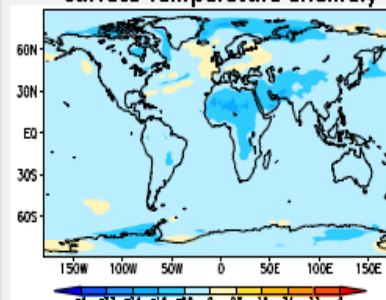
## Solar Radiative forcing for clear sky



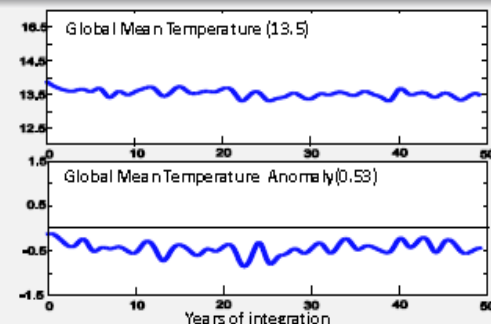
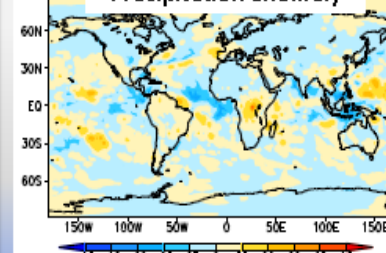
- ✓ TOA direct solar radiative flux for clear sky conditions indicate negative forcing over tropics
- ✓ The Surface radiative forcing show considerable decrease over regions with high natural aerosol loading
- ✓ The atmospheric radiative forcing show an increase in absorption of solar energy over high aerosol regions

Surface Temperature & Precipitation Change

## Surface Temperature anomaly



## Precipitation anomaly



Global Tropics	
Temperature Change $\sim -0.48^\circ C$	Precipitation Change 0.06 mm/day

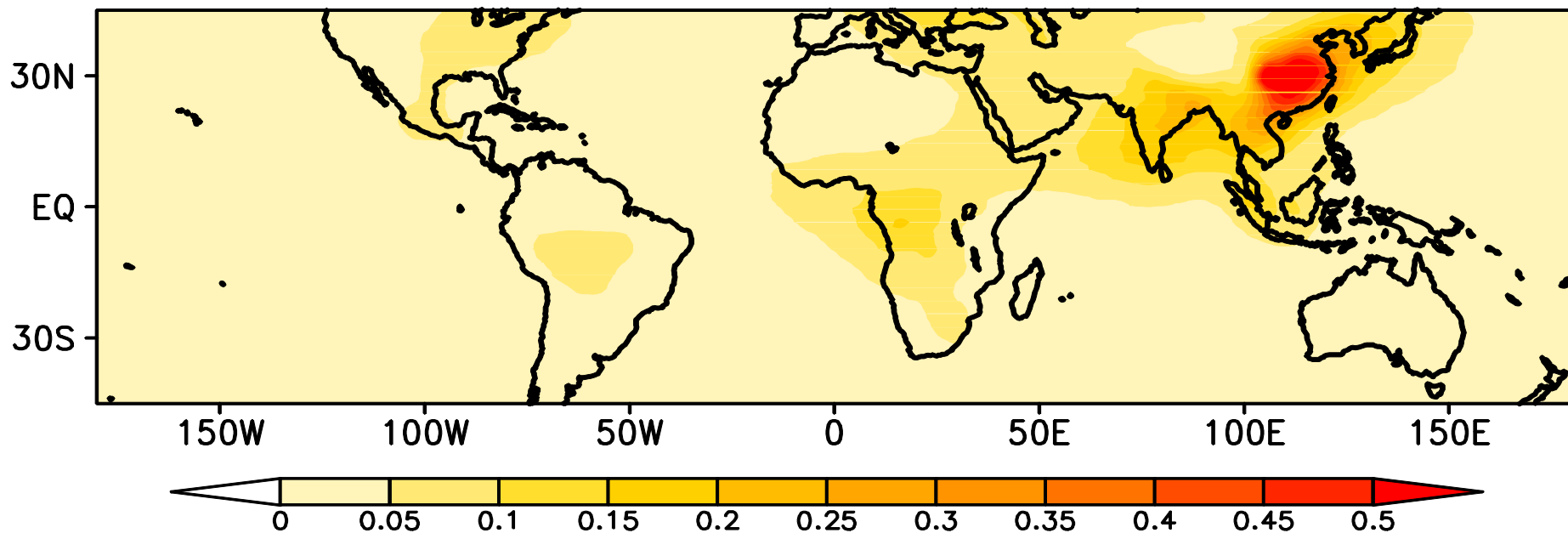
- ✓ Surface temperature response indicates cooling over almost entire globe, specially over northern hemispheric continents
- ✓ Precipitation anomaly shows a reduction over various tropical regions like South Asia and equatorial Atlantic

Pre-Industrial model setup

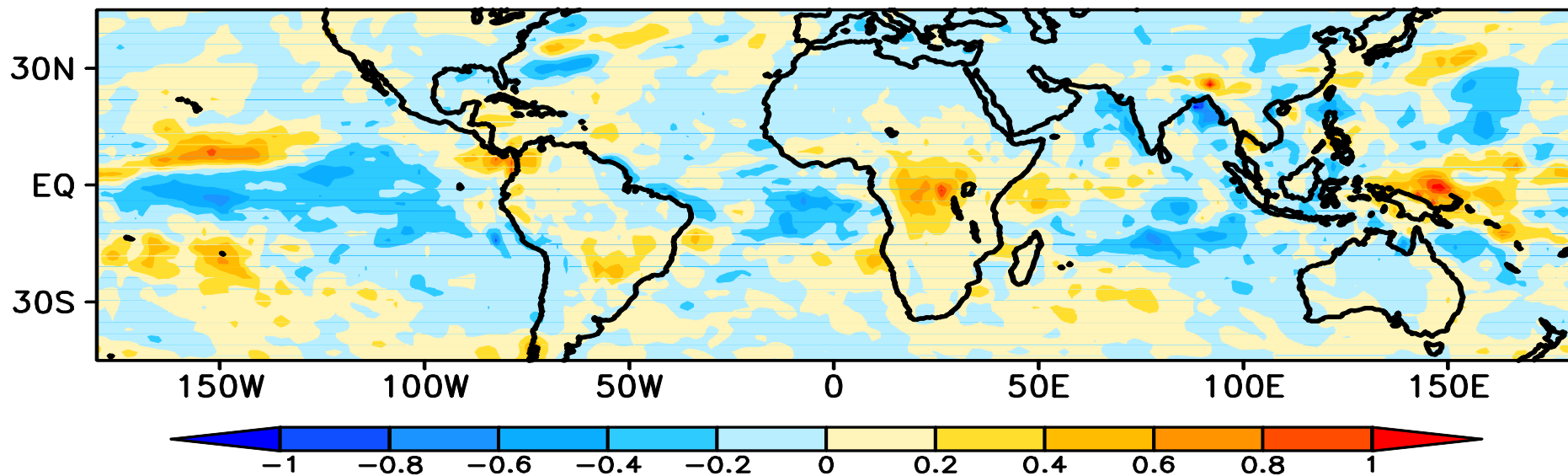
Aerosol Direct Radiative Forcing



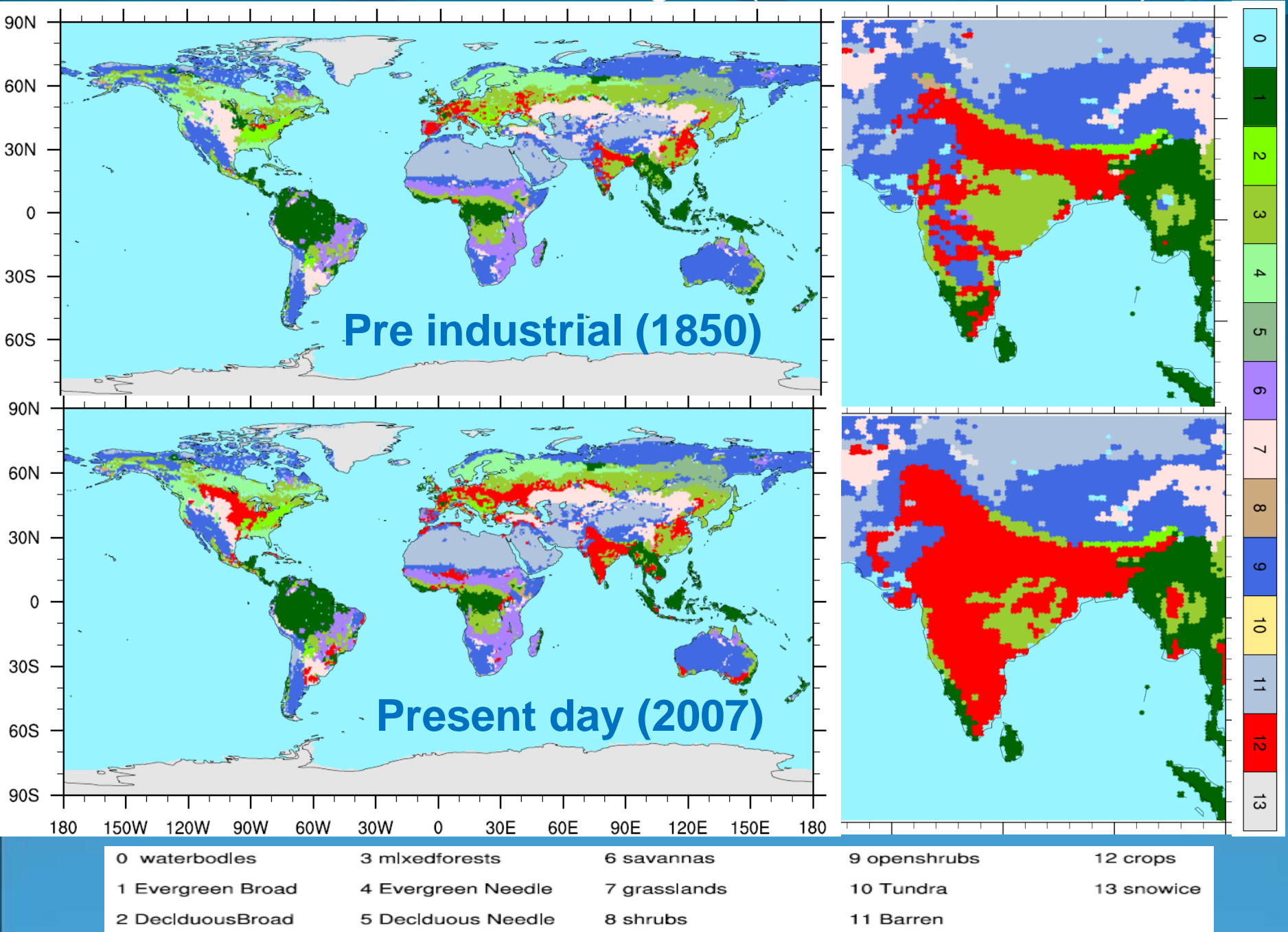
AOD-550nm(2005-1850)



Precipitation(2005-1850)



# Land use/land cover changes (Hurtt et al., 2015)



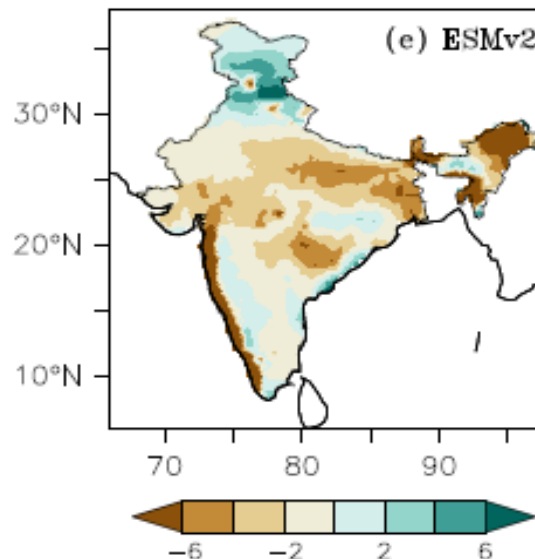
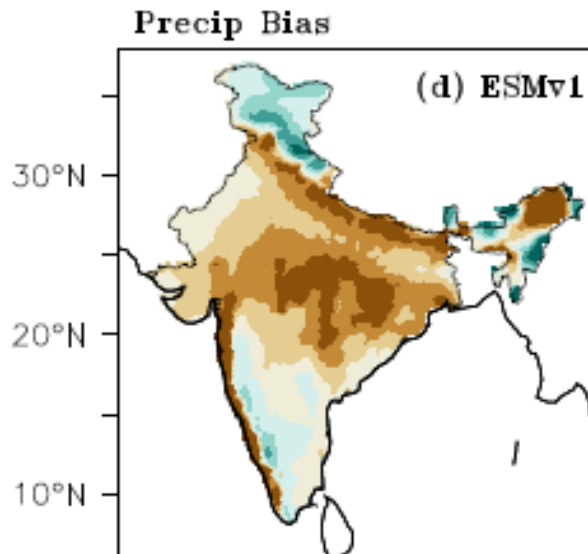
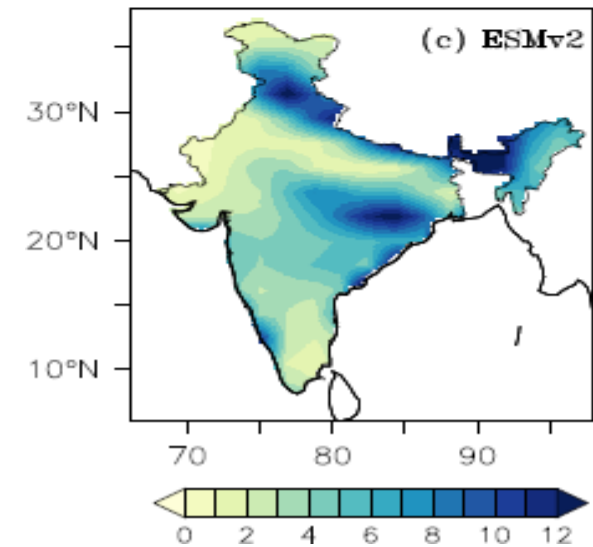
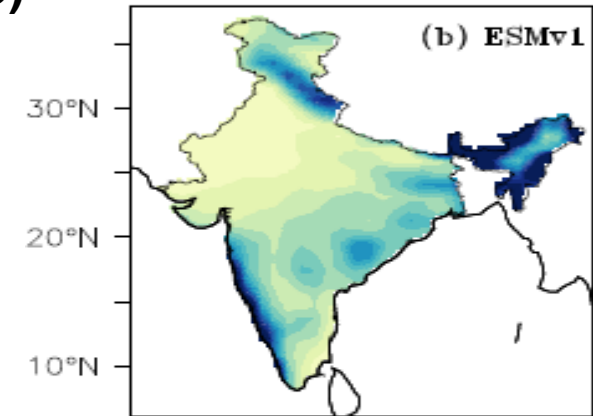
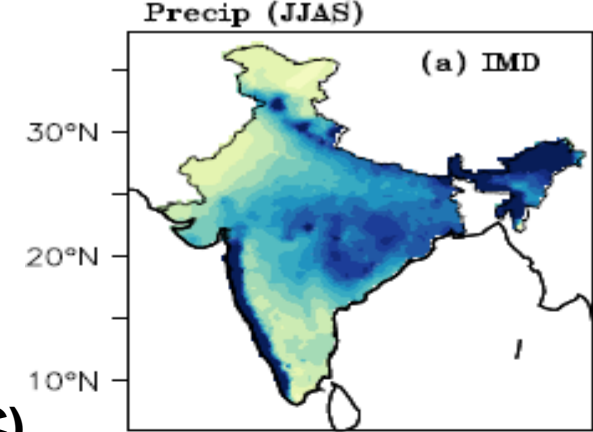
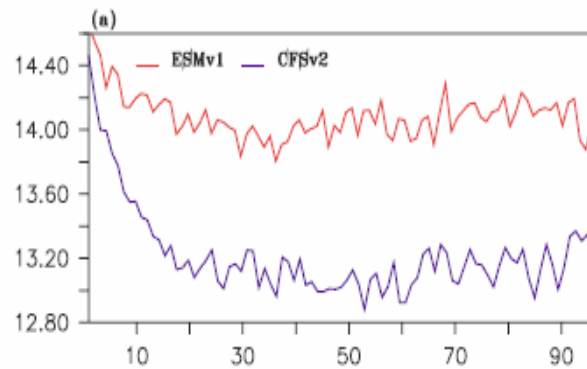
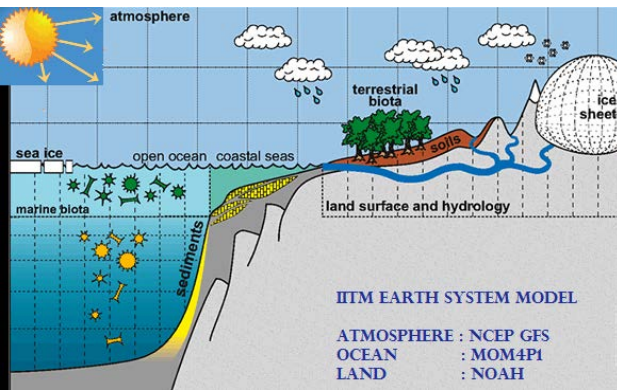
# THE IITM EARTH SYSTEM MODEL

Transformation of a Seasonal  
Prediction Model to a Long-Term  
Climate Model

BY P. SWAPNA, M. K. ROXY, K. APARNA, K. KULKARNI, A. G. PRAJEESH,  
K. ASHOK, R. KRISHNAN, S. MOORTHY, A. KUMAR, AND B. N. GOSWAMI

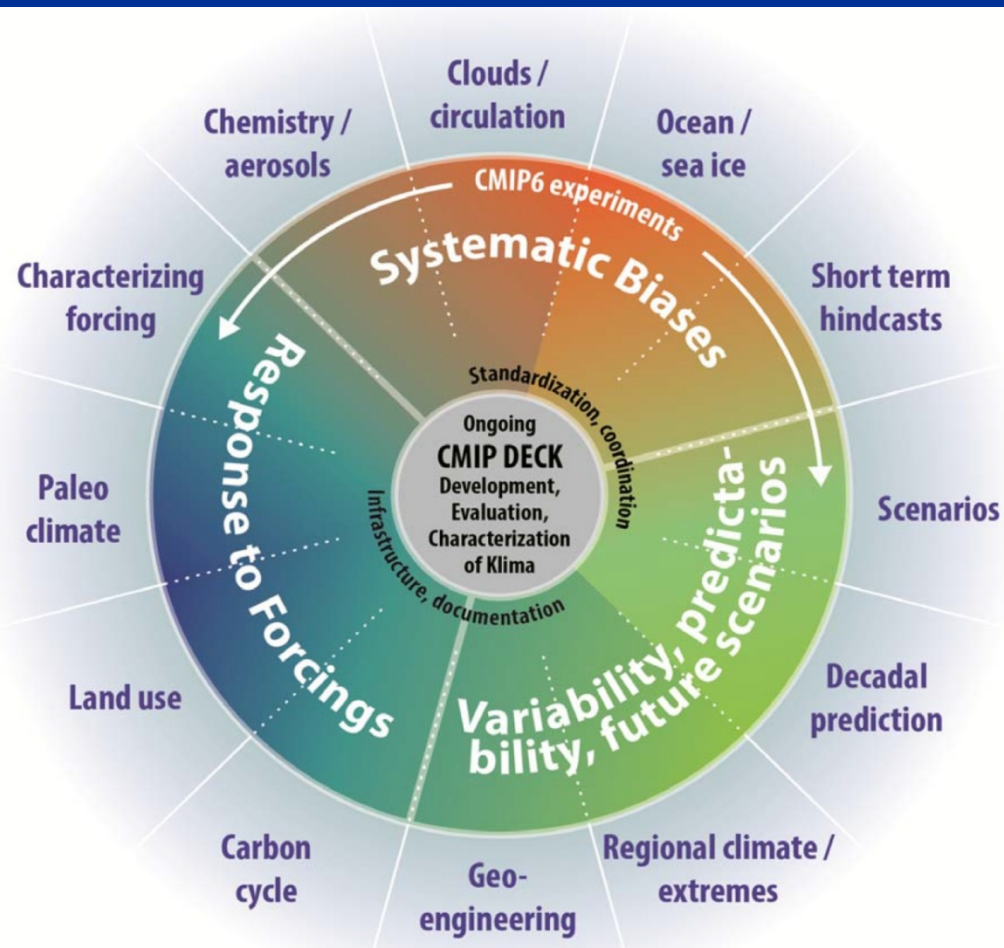
Centre for Climate Change Research, IITM, Pune

The first climate model from India to contribute to the  
next Intergovernmental Panel on Climate Change (IPCC)



# CMIP6 Schematic: Participation in the 6<sup>th</sup> Intergovernmental Panel for Climate Change (IPCC)

Initial proposal for the CMIP6 experimental design has been released



CMIP6 Concept: A Distributed Organization under the oversight of the CMIP Panel

IITM ESM will participate in the climate modeling CMIP6 experiments for the IPCC 6<sup>th</sup> Assessment Report



# The DECK experiments

## The DECK experiments:

- provide continuity across past and future phases of CMIP
- evolve only slowly with time
- are already common practice in many modelling centres
- are to be done by all participating coupled models

## Specifically:

1. an AMIP simulation (~1979-2010);
2. a multi-hundred year pre-industrial control simulation;
3. a 1%/yr CO<sub>2</sub> increase simulation to quadrupling to derive the transient climate response;
4. an instantaneous 4xCO<sub>2</sub> run to derive the equilibrium climate sensitivity;
5. a simulation starting in the 19th century and running through the 21<sup>st</sup> century using an existing scenario (RCP8.5).

# Summary

• **IITM ESMv1**: First version of IITM ESM has been successfully developed at CCCR, IITM by incorporating MOM4P1 (with ocean biogeochemistry) component in CFSv2. Major improvements are seen in IITM ESMv1 vis-à-vis CFSv2 :

- Significant reduction of cold bias of global mean SST by  $\sim 0.8^{\circ}\text{C}$
- ENSO & PDO are robust and spatially more coherent in ESM1.0
- ENSO and monsoon links are well-captured
- The IITM Earth System Model: Transformation of a seasonal prediction model to a long term climate model - *Swapna et al. 2015 (Bulletin of American Meteorological Society)*

• **IITM ESMv2**: Further improvements are incorporated in IITM ESMv1

- Reduced TOA radiation imbalance significantly
- Improved mean monsoon precipitation over South Asia
- Improved sea-ice distribution in the Arctic and Antarctic
- Improved Atlantic Meridional Overturning Circulation (AMOC)
- Interactive ocean biogeochemistry
- Included time-varying aerosol properties (3D fields) for the CMIP experiments
- Improved hydrological balance through discharge of runoff from land to ocean

• **IITM ESM** to participate in the upcoming CMIP6 activity & IPCC AR6 assessment



# Future Roadmap

- **Basic research:** Scientific questions on detection, attribution and future projections of global and regional climate change, including the South Asian monsoon, in addition to contribution to CMIP6 and IPCC AR6
- Development of High Resolution Global Model (~grid size 27 km) Atmospheric version of IITM-ESM for dynamical downscaling. Generation of high resolution global climate and monsoon projections.  
**Timeline: 2018-2021**
- High-resolution IITM-ESM coupled model (atmosphere grid size: 27 km, ocean grid: 0.5 deg x 0.5 deg and 0.25 deg x 0.25 deg near equator) for long-term climate. **Timeline: 2020-2025+**
- Development of next-generation IITM-ESM coupled model, to include new components (eg., interactive aerosols, chemistry, carbon cycle).  
**Timeline: 2020-2025+**

**Thank you**

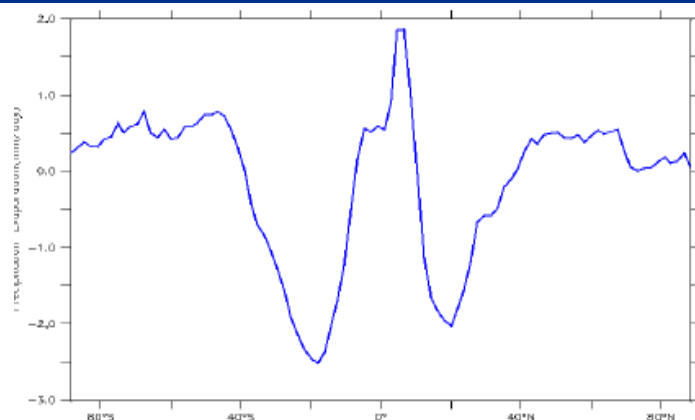
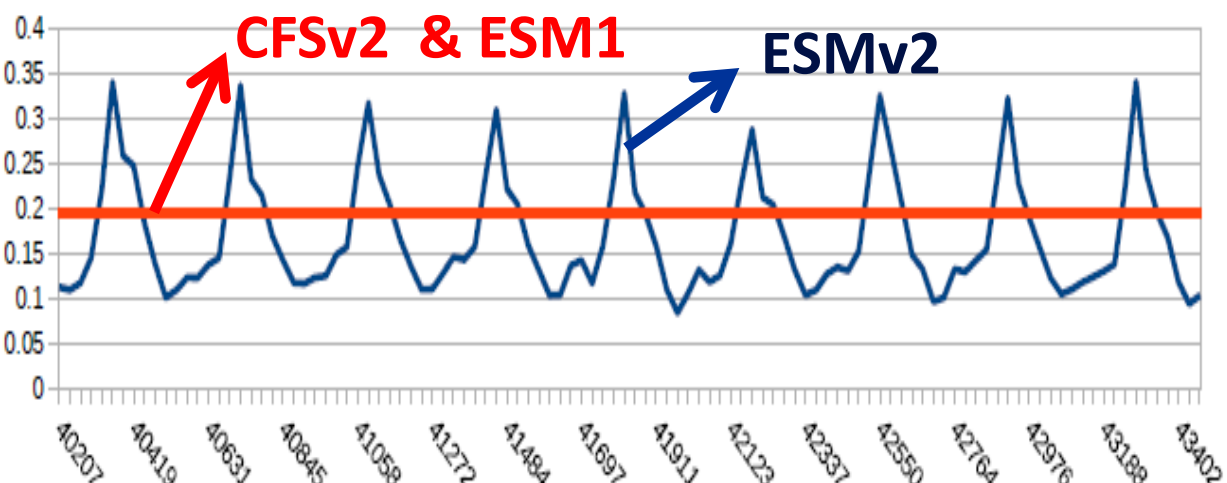
# Water balance in ESMv2

CFSv2 and ESMv1: Constant value of runoff was used in the Ice Model

ESMv2: Runoff calculated from Land Model & discharged into the nearest ocean point

Runoff ( $\text{kg m}^{-2} \text{s}^{-1}$ )

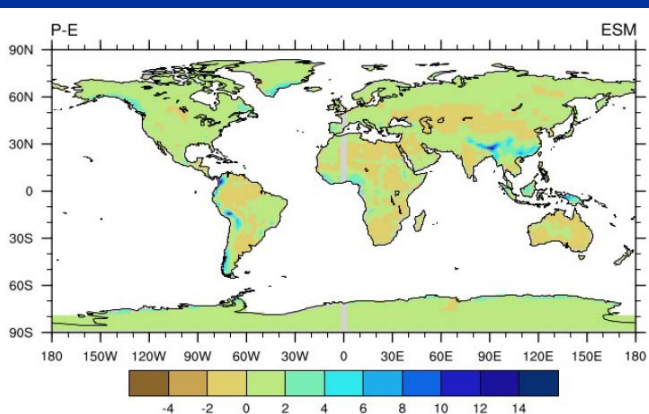
Zonal mean (P-E) in  $\text{mm day}^{-1}$



Latitude

Precipitation minus Evaporation

Runoff from Land Model



Hydrology statistics

Total Runoff from  
Land =  $1.06 \times 10^9 \text{ kg s}^{-1}$

Total Water Discharge into  
Ocean =  $1.06 \times 10^9 \text{ kg s}^{-1}$

