# Attempts in improving low clouds, deep clouds and mean states of CFSv2

## Partha Mukhopadhyay

Phani Murli Krishna, Medha Deshpande, M. Mahakur, Malay Ganai, Sahadat Sarkar, Snehlata Tirkey, Shilpa Malviya, Tanmoy Goswami, Kumar Roy

Student who worked: Abhik S. and Bidyut Goswami

Dr. V. S. Prasad and Dr. R. Ashrit (NCMRWF)

## Work presented here are from the following papers

Goswami et al. 2014

Bidyut B. Goswami, R. P. M. Phani, P. Mukhopadhyay, Marat Khairoutdinov and B. N. Goswami, "Simulation of the Indian Summer Monsoon in the Superparameterized Climate Forecast System version 2: Preliminary Results", **J. of Climate**, 2015

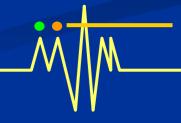
Abhik et al. 2015, Clim Dyn.

Ganai Et al. 2014, Clim. Dyn on Revised SAS

Ganai et al. 2016, JGR, diurnal

Abhik et al. 2017 (JAMES under revision)





# Outline

- Paradigm of Conventional Parameterization
- •Issues of CFSv2 biases related to convection
- Recent approaches in dealing convection parameterization in CFSv2
- •High resolution GFS/GEFS
- Future plans

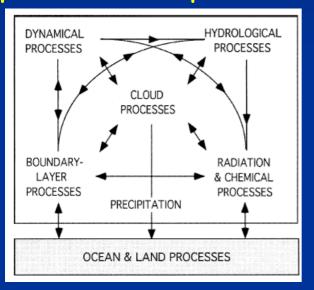


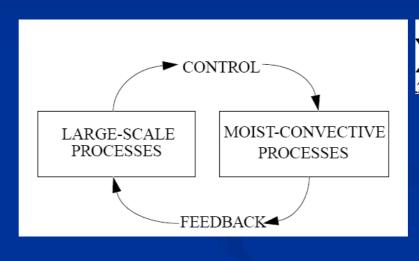


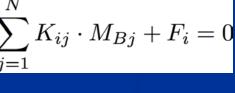
## Issues of cumulus Parameterization

The Cumulus Parameterization Problem: Past, Present, and Future By Akio Arakawa, JOC, 2004, Arakawa et al. 2011, Arakawa and Wu 2013, Wu and Arakawa 2014

• "Major practical and conceptual problems in the conventional approach of cumulus parameterization, includes inappropriate separations of processes and scales".







 $K_{ij}$  = effect of cloud j on cloud i.

 $F_i$  = environmental forcing for

cloud i

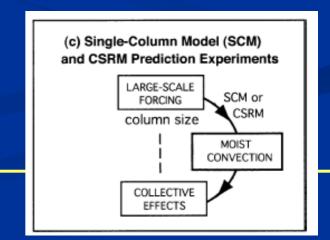
 $M_{Bj}$  = mass flux at base of cloud j

Task of Conv. Param

To calculate the collective effects of an ensemble of convective clouds in a model column



$$Q_{1C} = Q_1 - Q_R = L(\overline{c} - \overline{e}) - \frac{\partial \overline{\omega' s'}}{\partial p}$$

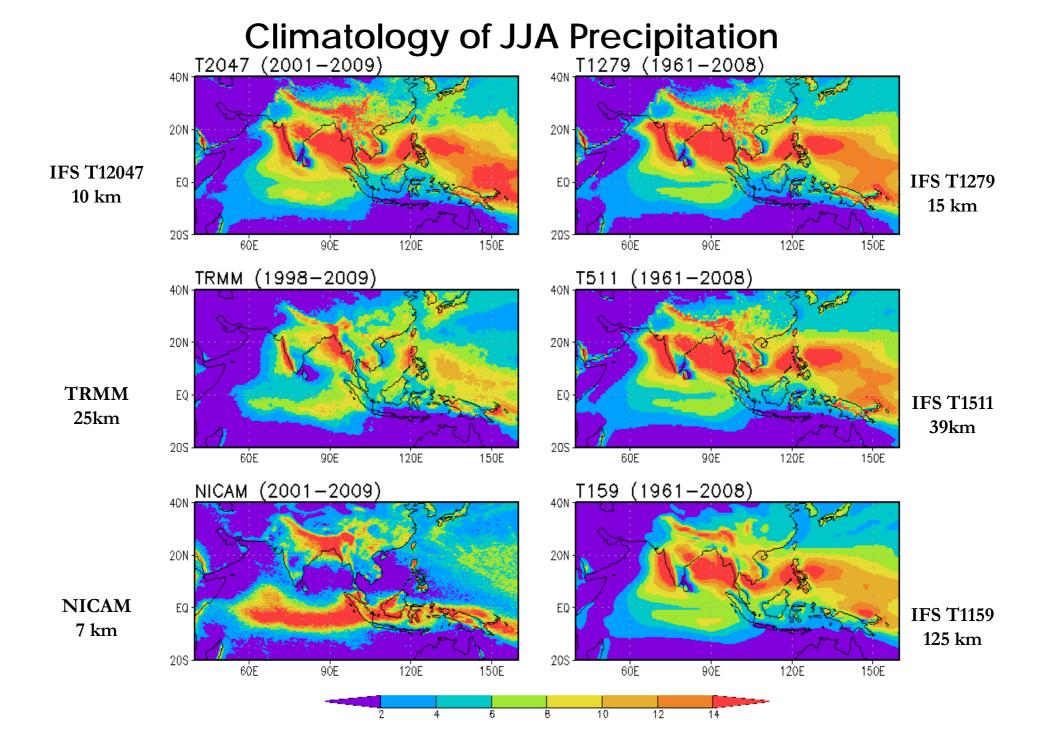


### **ISSUES**

- CFSv2 T126 shows colder Tropospheric temperature bias and colder SST bias
- CFSv2 T382 shows warmer Tropospheric temperature and warmer SST bias
- Inspite of contrasting bias, the rainfall bias in both the models are similar
- CFSv2T126 & CFSv2 T382 both produce too much frequency of lighter rainfall (drizzle) and shows dry bias over Indian land mass but northward propagation is reasonable in both.
- •CFSv2T126 & CFSv2 T382 both underestimates synoptic variance and overestimates ISO variance
- •Diurnal Convective lifecycle is equally incorrect in CFSv2T126 & CFSv2 T382. (Deep convection is lacking)

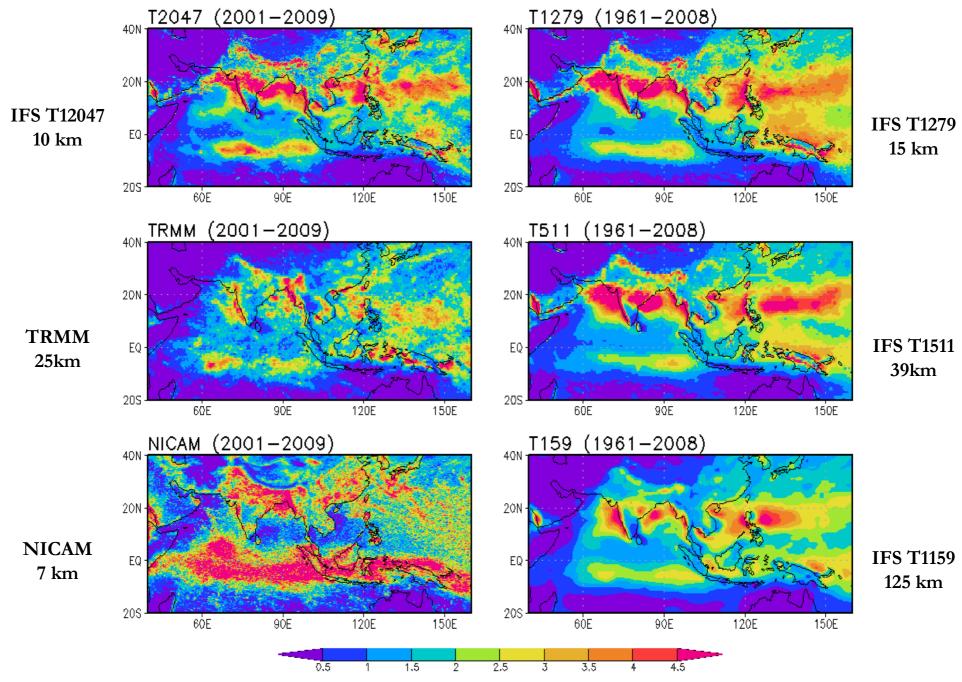
#### CFSv2 T126 bias CFSv2 T382 bias (a) Precip (mm/day): CFS\_T126 - TRMM 15N EQ **15S** 150W 6ÒE 150E 120W 3ÓE 15S (b) SST (deg C): CFS\_T126 - OISST 30N 15N 15N EQ **15S** 15S 180 150W 6ÒE (c) U850 (m/s): CFS T126 - NCEP 30N 15N EQ 15S **15S** 90E 120E 150E 180 150W 120W 6ÒE 30N (d) TT (K): CFS\_T126 - NCEP 15N 30N EQ 15S **15S** 30E 60E 120E 150E 180 150W 120W 90W 150W 90E 120W 9ÓW 30E 60E

Seasonal mean bias in a) precipitation (mm day-1), b) SST (°C), c) zonal wind at 850 hPa (m s -1) and d) tropospheric temperature (TT, K) relative to TRMM, TMI and CFSR respectively



Slide courtesy Emilia Jin, Athena Workshop, ECMWF, 7-8 June 2010

## Standard Deviation of JJA Precipitation Anomalies



Slide courtesy Emilia Jin, Athena Workshop, ECMWF, 7-8 June 2010

Fig. 4 Probability distribution function (PDF) of daily rainfall (mm day<sup>-1</sup>) during all JJAS seasons with a bin width of 5 mm day<sup>-1</sup> in percentage over a central India (CI), b Bay of Bengal (BoB), c Arabian Sea (AS) and eastern equatorial Indian Ocean (EEIO). The regions are marked by white boxes in Fig. 3b

CFSv2T382

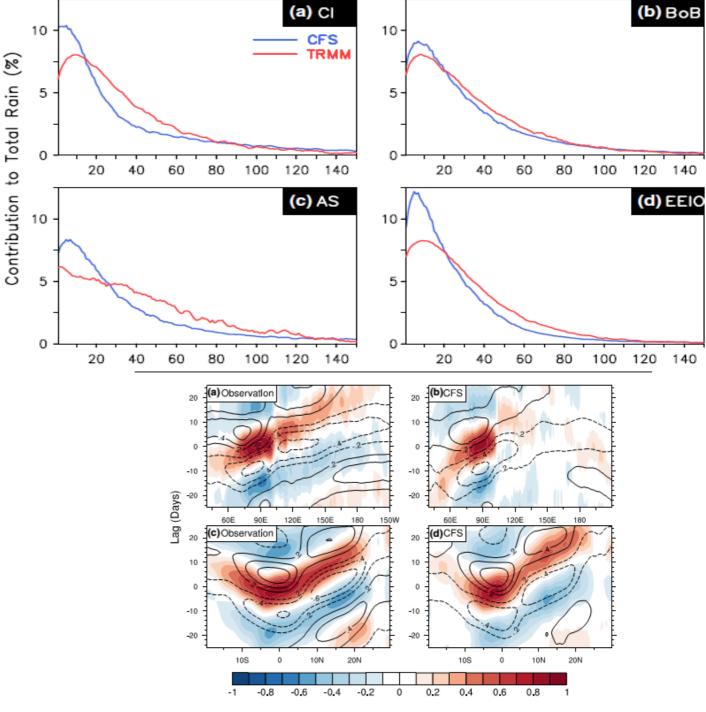
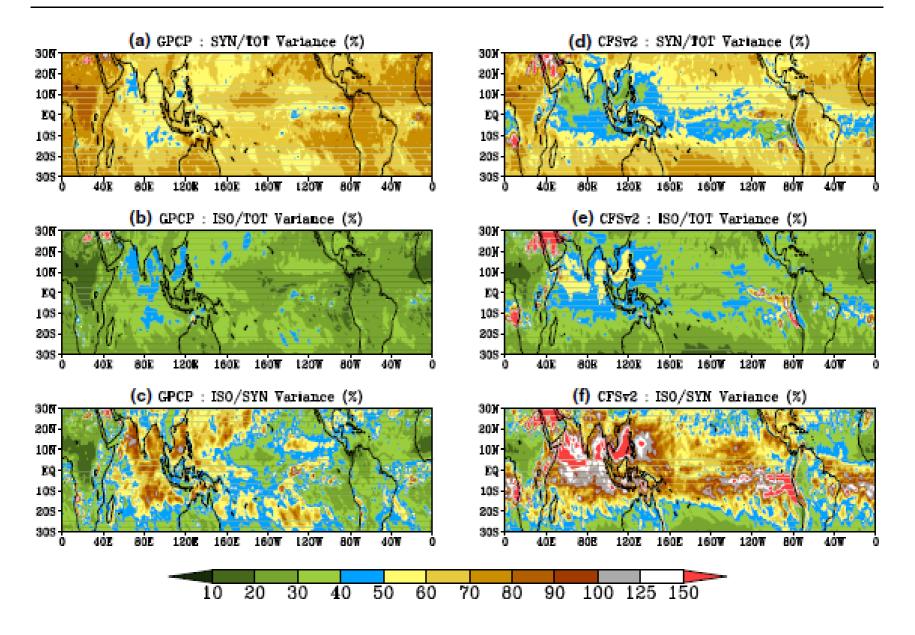


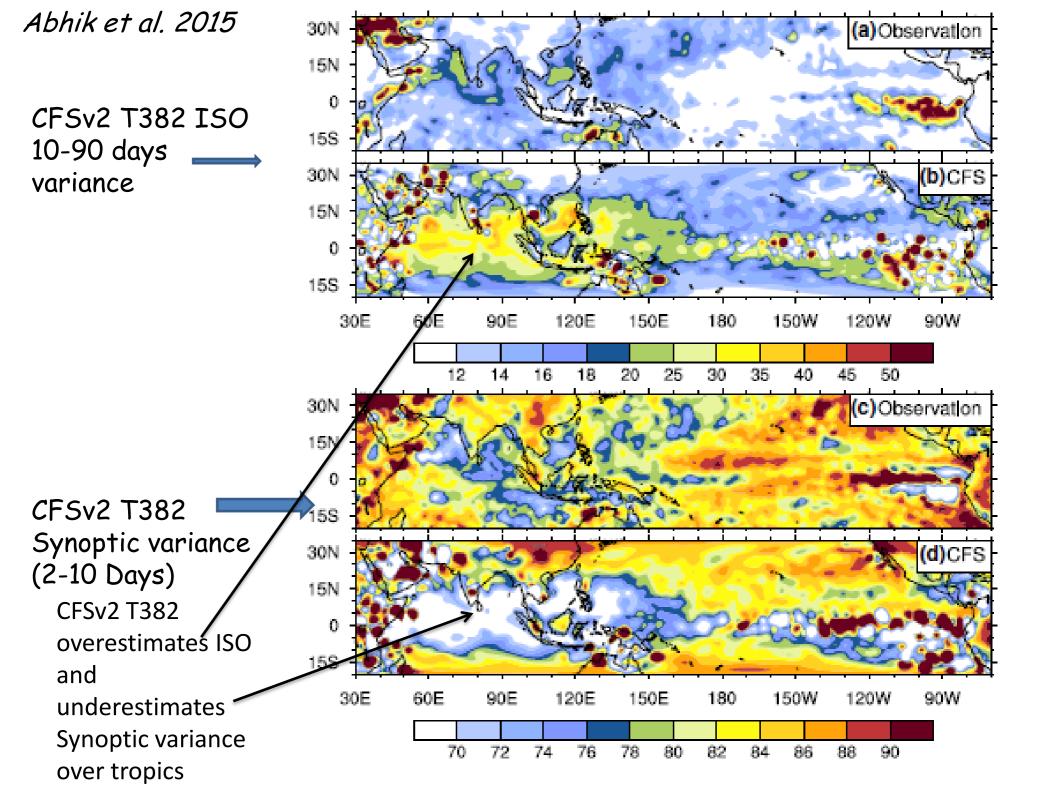
Fig. 8 a, b Longitude versus lag correlation and c, d latitude versus lag correlation of 20–100-day filtered precipitation (shaded) and U<sub>850</sub> (contour) with base 20–100-day filtered precipitation time series over

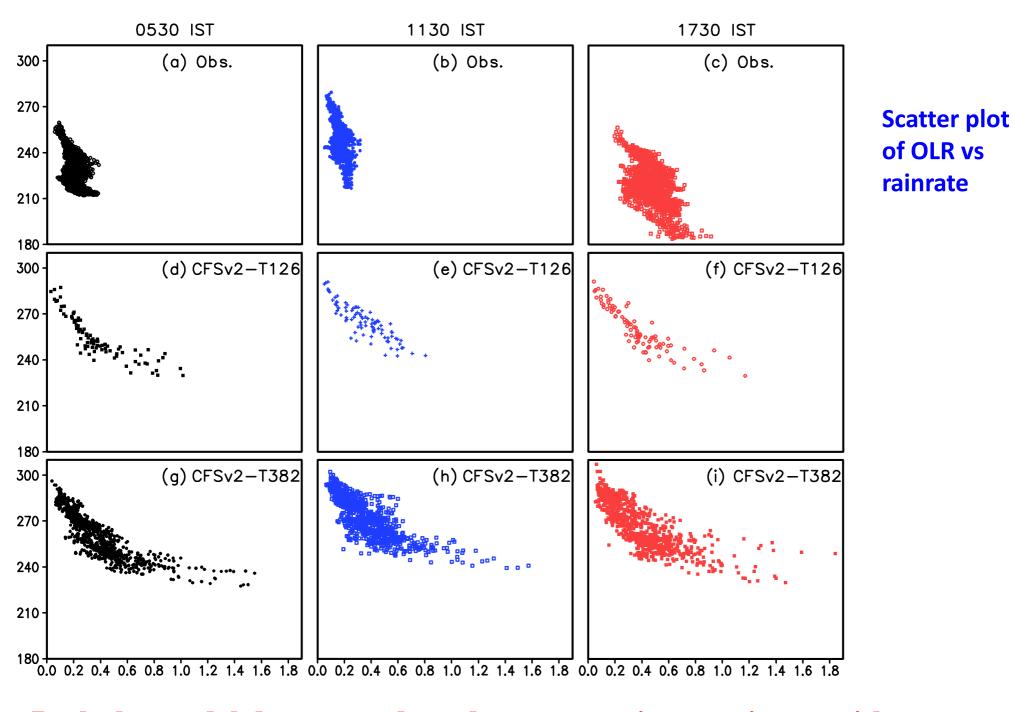
EEIO (10°S-5°N, 75°-100°E) for observation and CFS T382. For longitude-lag (latitude-lag) plot data are averaged between 70°E and 90°E (10°S and 10°N)

B. B. Goswami et al.



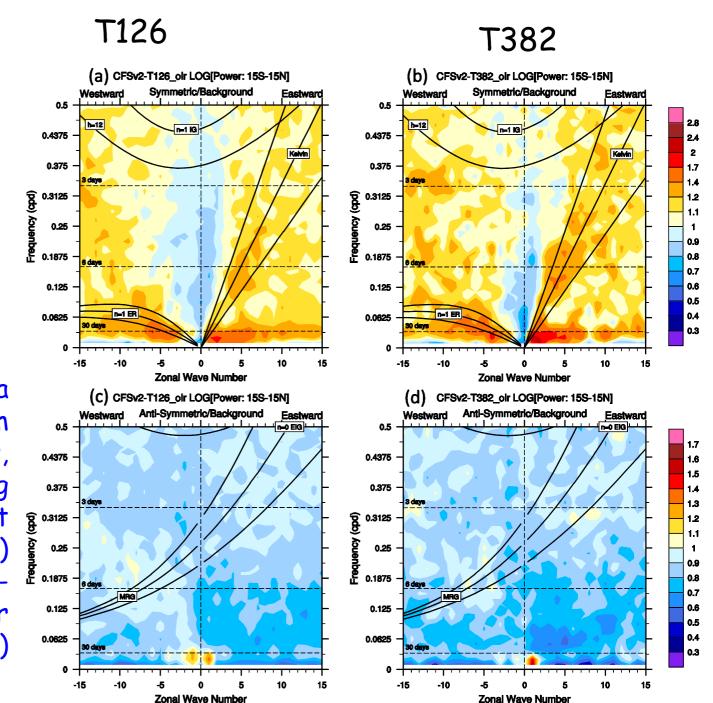
CFSv2 T126 (right panel) underestimates Synoptic variance and overestimates ISO variance over global tropics





Both the model do not produce deep convection consistent with too much of lighter precipitation

Ganai et al. 2015



Space-Time spectra (Wheeler-Kiladis diagram [Wheeler and Kiladis, 1999]) of OLR showing the symmetric component for (a) CFSv2-T126, (b) CFSv2-T382 and the antisymmetric component for (c) CFSv2-T126, (d) CFSv2-T382.

Goswami et al. 2015

#### CFSv2 T382

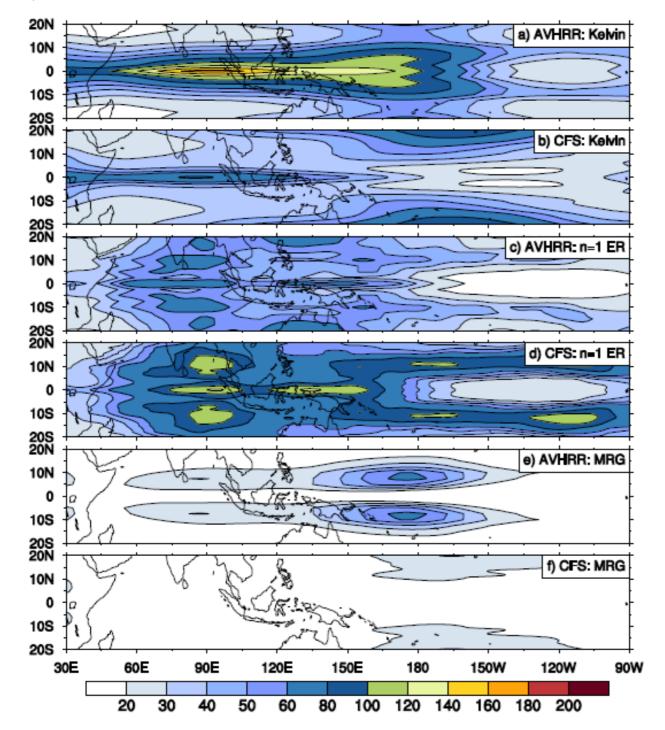


Fig. 13: Distribution of boreal summer time OLR variance (W<sup>2</sup> m<sup>-4</sup>) of (a), (b) Kelvin; (c), (d) n=1 ER and (e), (f) MRG waves for AVHRR and CFS.

Abhik et al., 2015

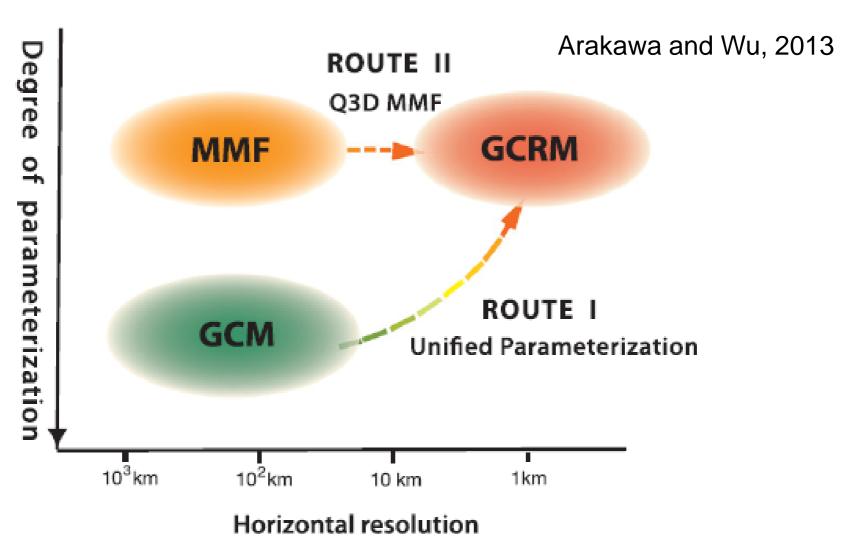


FIG. 3. Two routes for unifying the low- and high-resolution models.

Route II with 2D MMF: accomplished in IITM through development of SP-CFS

# Attempts of Improving the biases of CFSv2 through Superparameterized CFS (SP-CFS)

Bidyut B. Goswami, R. P. M. Krishna, P. Mukhopadhyay, Marat Khairoutdinov, and B. N. Goswami, 2015: Simulation of the Indian Summer Monsoon in the Superparameterized Climate Forecast System Version 2: Preliminary Results. J. Climate, 28, 8988-9012

10/11/2015

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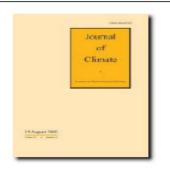
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Journal of Climate 2015 : e-View doi: http://dx.doi.org/10.1175/JCLI-D-14-00607.1

Simulation of the Indian Summer Monsoon in the Superparameterized Climate Forecast System version 2: Preliminary Results

BIDYUT B. GOSWAMI, R.P.M. KRISHNA, P. MUKHOPADHYAY, MARAT KHAIROUTDINOV, and B. N. GOSWAMI<sup>4</sup>

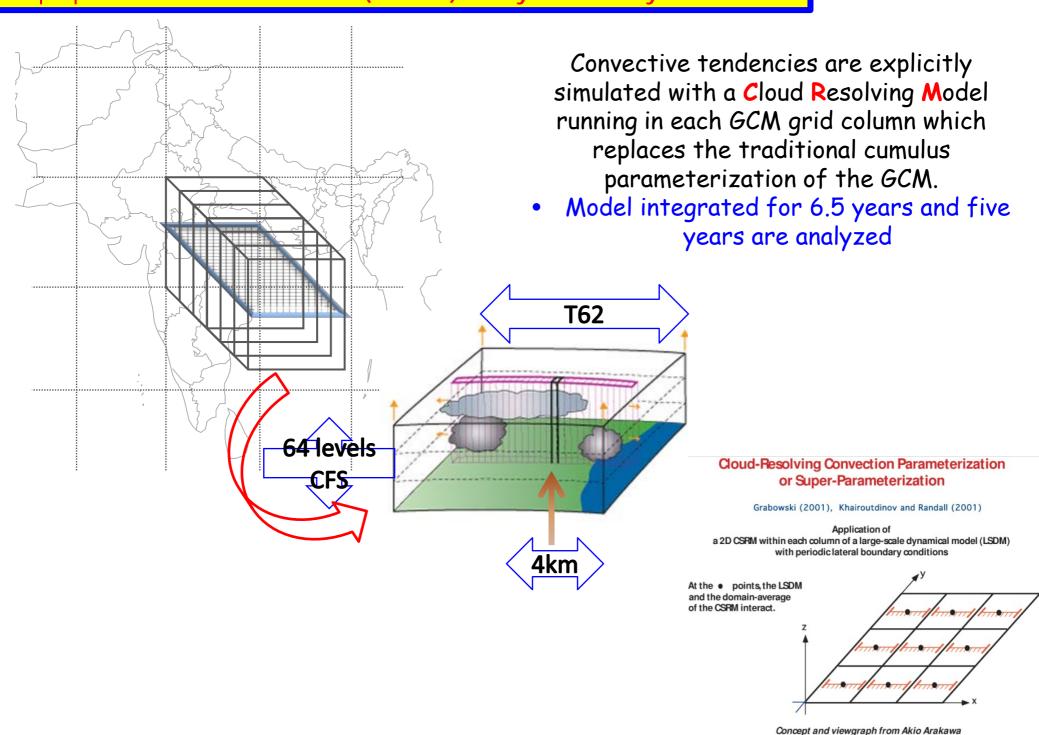
Department of Mathematics and Statistics, University of Victoria, Canada

<sup>2</sup> Indian Institute of Tropical Meteorology, Pune-411008, INDIA

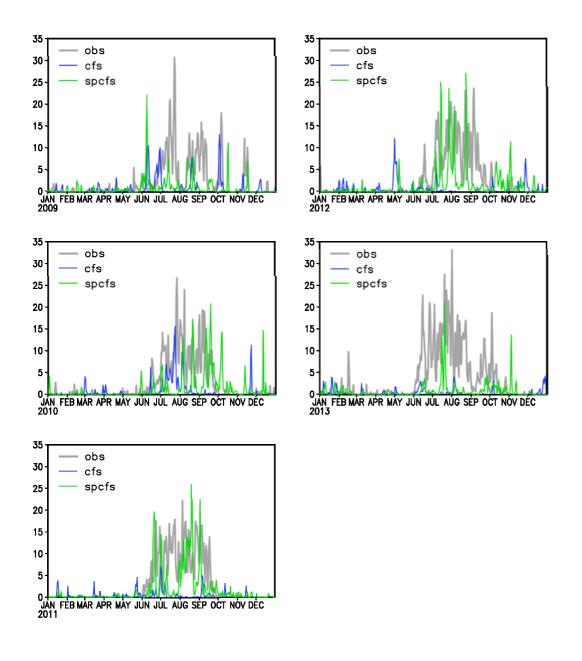
<sup>3</sup> School of Marine and Atmospheric Sciences, New York University, Stony Brook, USA

<sup>4</sup> Pisharoty Chair Professor, MoES, Indian Institute of Science Education and Research, Pune-411008, INDIA

## Superparameterized CFSv2-T62 (SPCFS) Analyses of 6.5 year free run



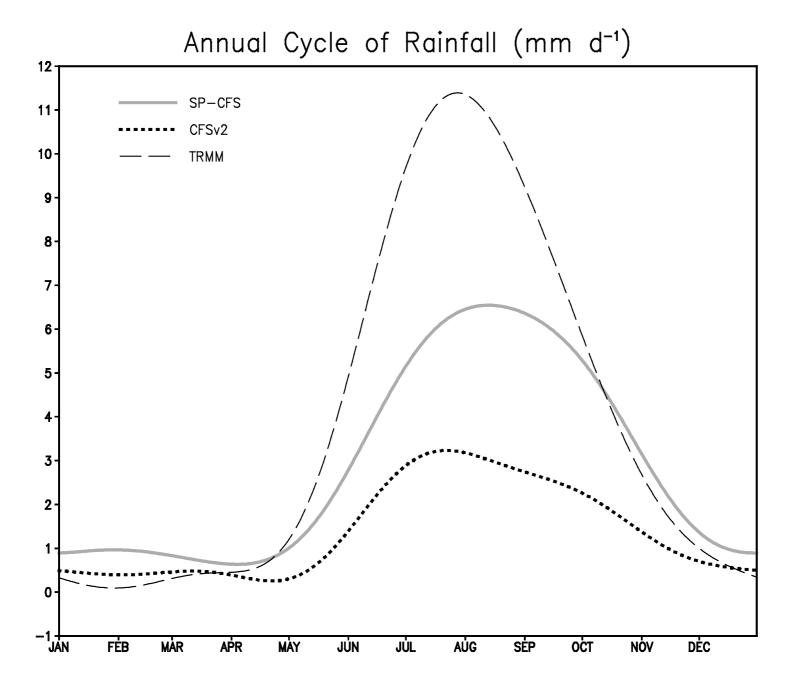
#### The rainfall is averaged over: 73-82E; 18-28N



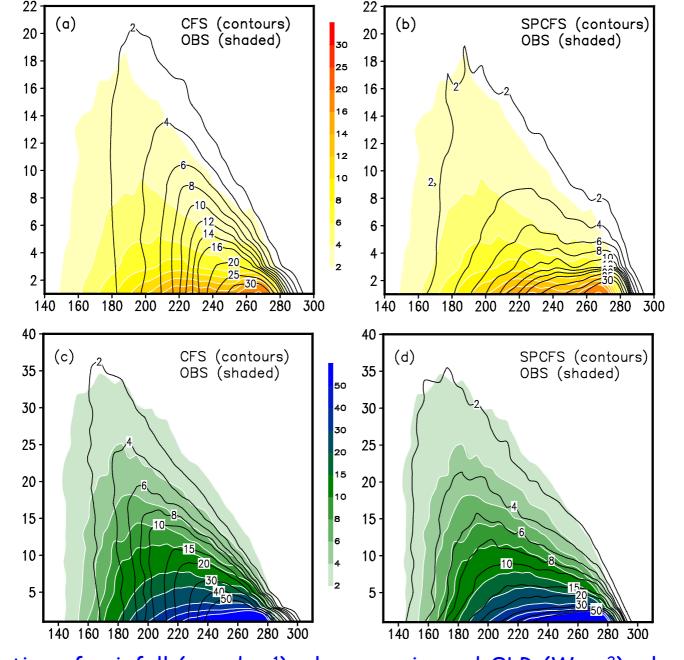
SP-CFS produces reasonable rain, CFS hardly rains

The Standard Dev for JJAS (5 years):

IMD=5.01 SPCFS=4.33 CFS=1.8



Annual cycle of the climatological mean rainfall (mm day<sup>-1</sup>) averaged over the area: 15°N-25°N; 75°E-90°E.

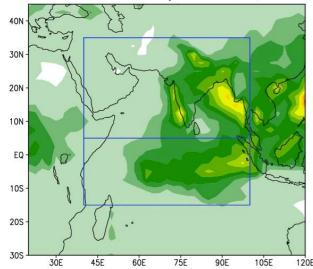


Joint distribution of rainfall (mm day<sup>-1</sup>), along y-axis, and OLR (W m<sup>-2</sup>), along x-axis, computed for each grid point, (a) & (b) over the monsoon domain bounded by  $15^{\circ}S-30^{\circ}N$  and  $50^{\circ}E-110^{\circ}E$  and (c) & (d) over the entire Tropics within  $15^{\circ}S-15^{\circ}N$ , for the 5 boreal summers (JJAS). For observation we have taken TRMM rainfall and NOAA OLR. Model simulated values are contoured and overlaid on observation (in shading). The values are in multiples of 100.

## North box : 40-100E; 5-35N

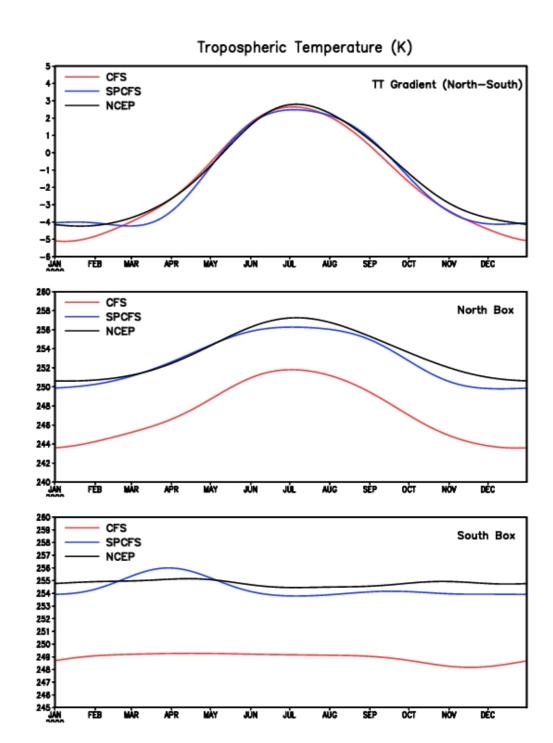
South box : 40-100E;15S-5N

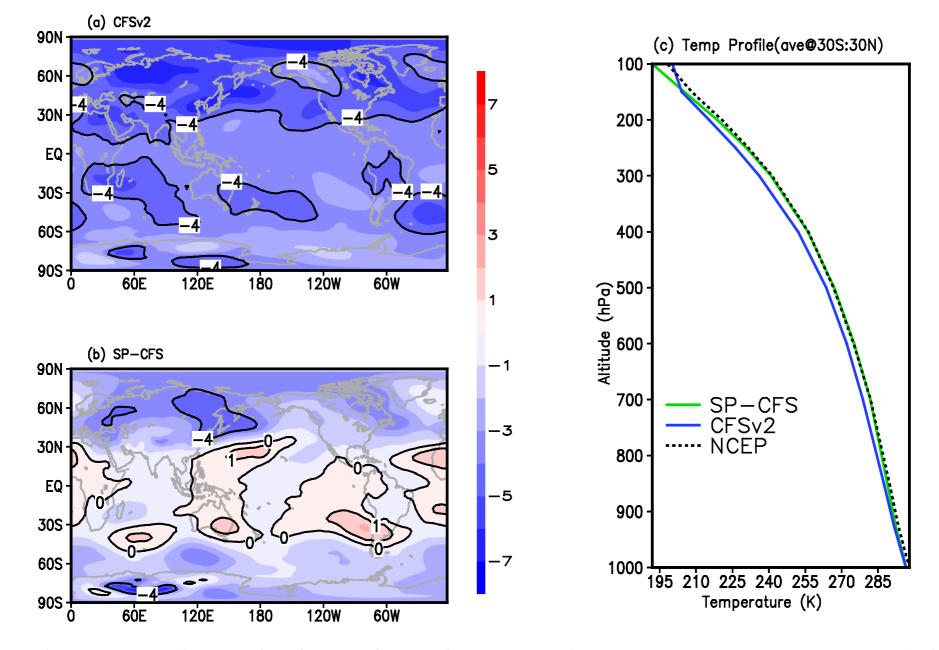
600-200hPa (Xavier et. al. 2007)



Improvement in tropospheric temperature bias is seen in TT gradient. Even though the Gradient looks reasonable in both CFS and SPCFS, but the bias is seen when we see the North and South boxes individually. The TT-gradient in a cooler background in CFS perhaps is consistent with reasonable circulation pattern (Fig-12 in manuscript) but deficient moisture (Fig-13b in manuscript) leading to dry monsoon.

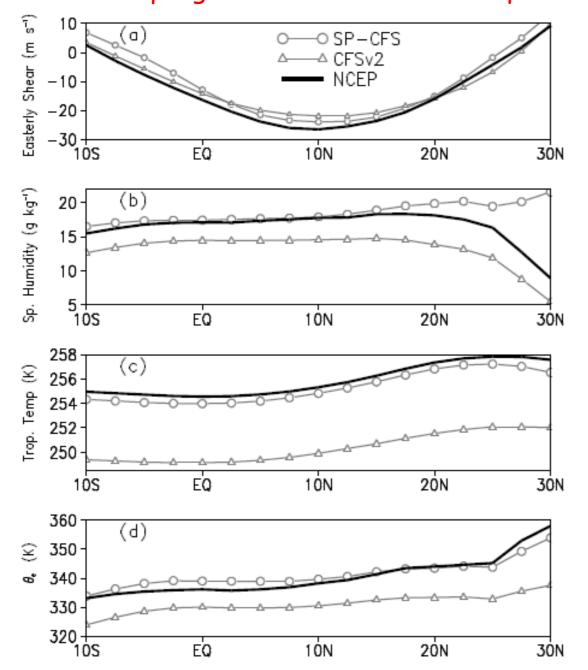
## Right result due to wrong reason in CFSv2



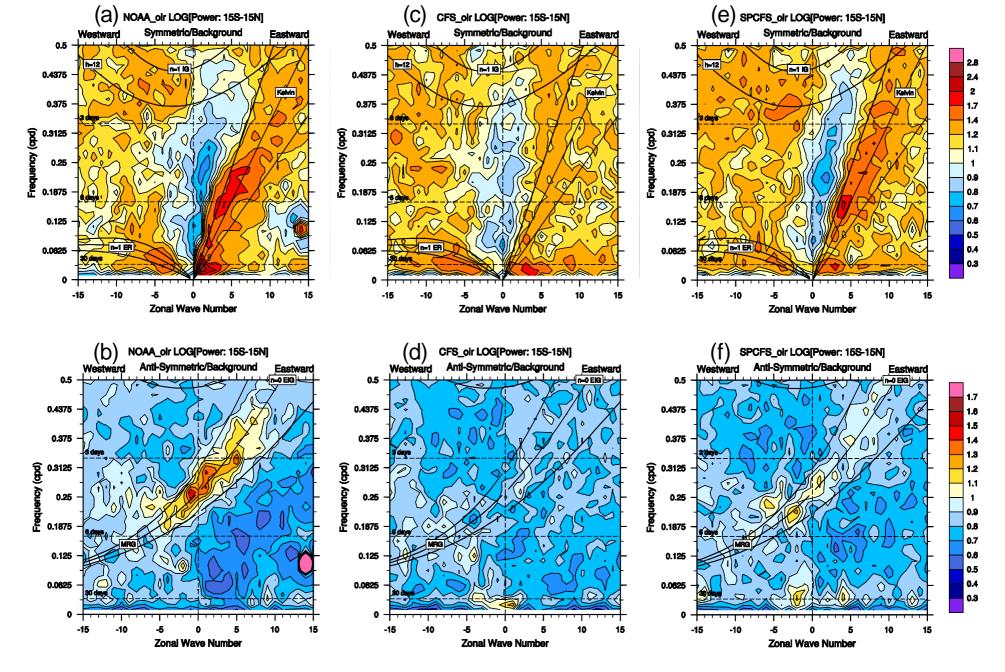


Boreal summer (JJAS) climatological Tropospheric temperature bias of (a) CFSv2 and (b) SP-CFS, relative to NCEP. (Averaged between 600hPa-300hPa). (c) Vertical profile of JJAS mean climatological temperature for tropics (30°S-30°N; 0°E-360°E).

# Mean state in SP-CFS has improved due to improvement in moist instability and convective coupling as evident in the subsequent slides

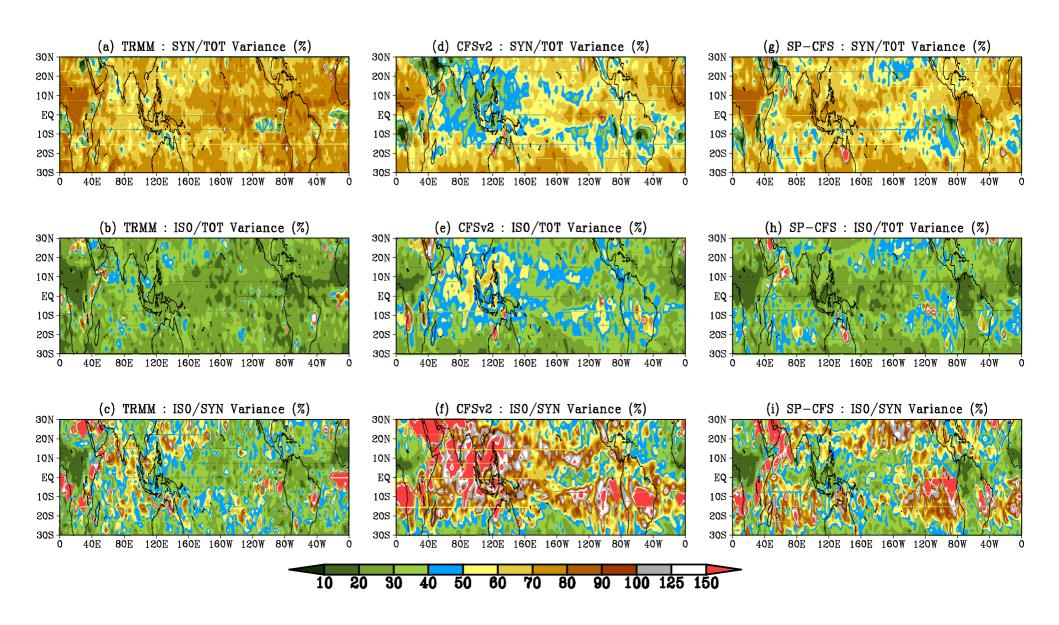


Climatological Seasonal meridional mean distribution of (a) easterly wind shear (U200-U850, m s-1), (b) surface specific humidity (g kg-1), (c) tropospheric (averaged temperature between 200 and 600 hPa) and (d) equivalent potential temperature (averaged between 1000 to 850 hPa and 65° to 95°E.

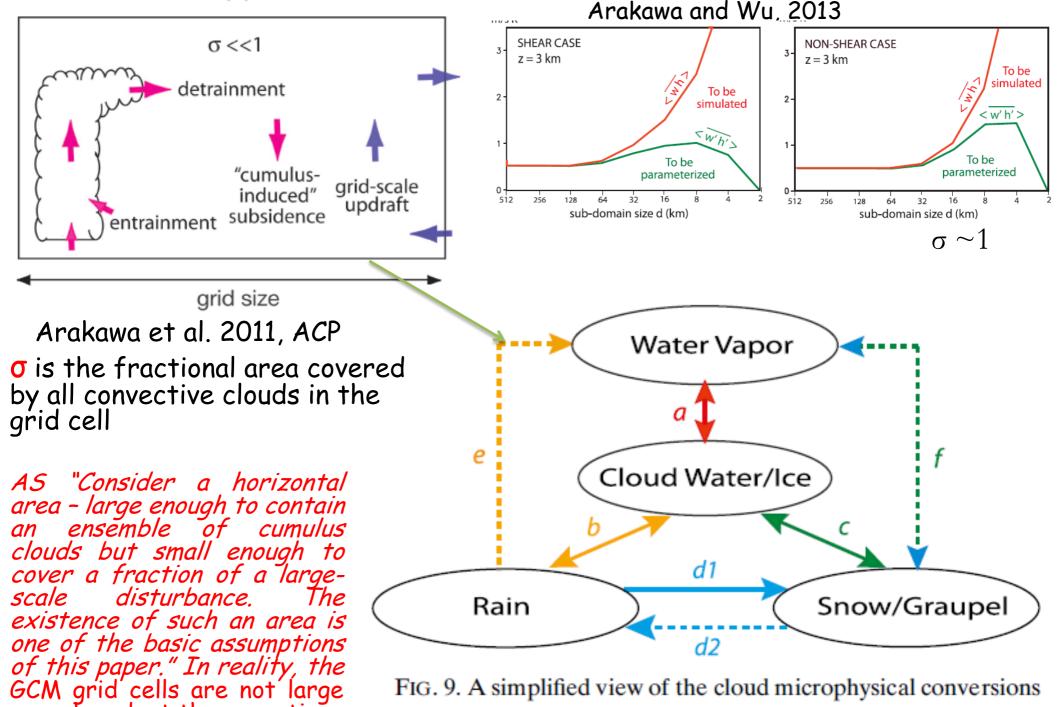


Space-Time spectra (Wheeler-Kiladis diagram [Wheeler and Kiladis, 1999]) of OLR showing the symmetric component for (a) NOAA OLR, (c) CFSv2 and (e) SP-CFS and the anti-symmetric component for (b) NOAA OLR, (d) CFSv2 and (f) SP-CFS.

### Ratio of Synoptic to ISO variance.



SP-CFS has improved the bias in synoptic and ISO variance



enough and, at the same time,

not small enough.

FIG. 9. A simplified view of the cloud microphysical conversions included in the CRM simulation used in this paper. See text for

Arakawa and Wu, 2014

#### Pros

- Superparameterized CFS (SP-CFS) demonstrates the merits of resolving cloud processes in a GCM
- Can be a good test bed
- Short run to for extreme events feasible (Li et al. 2014)

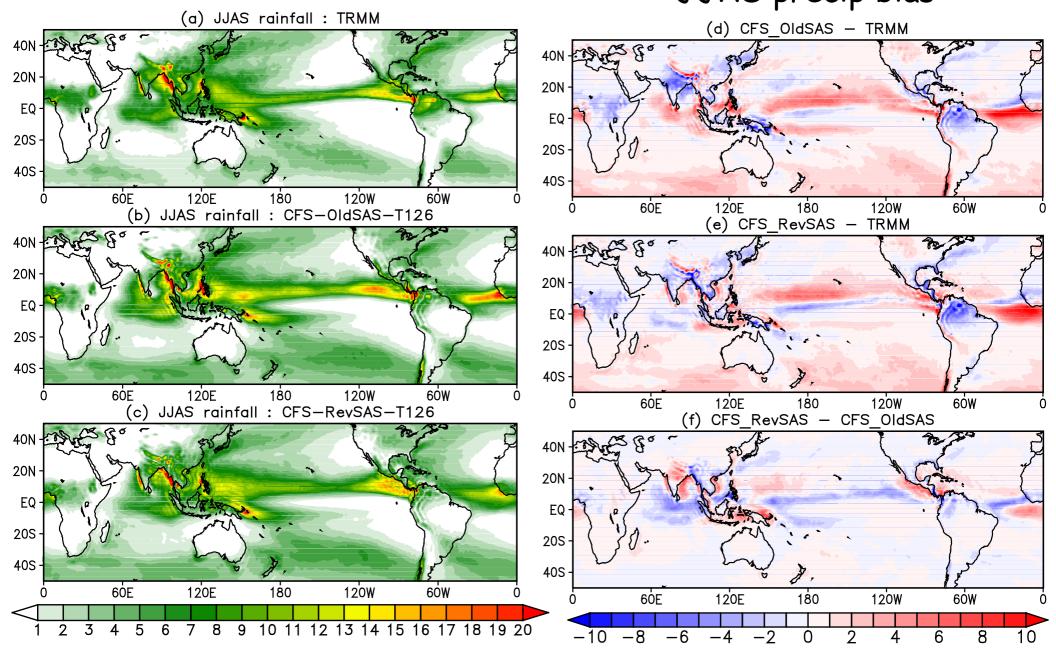
#### Cons

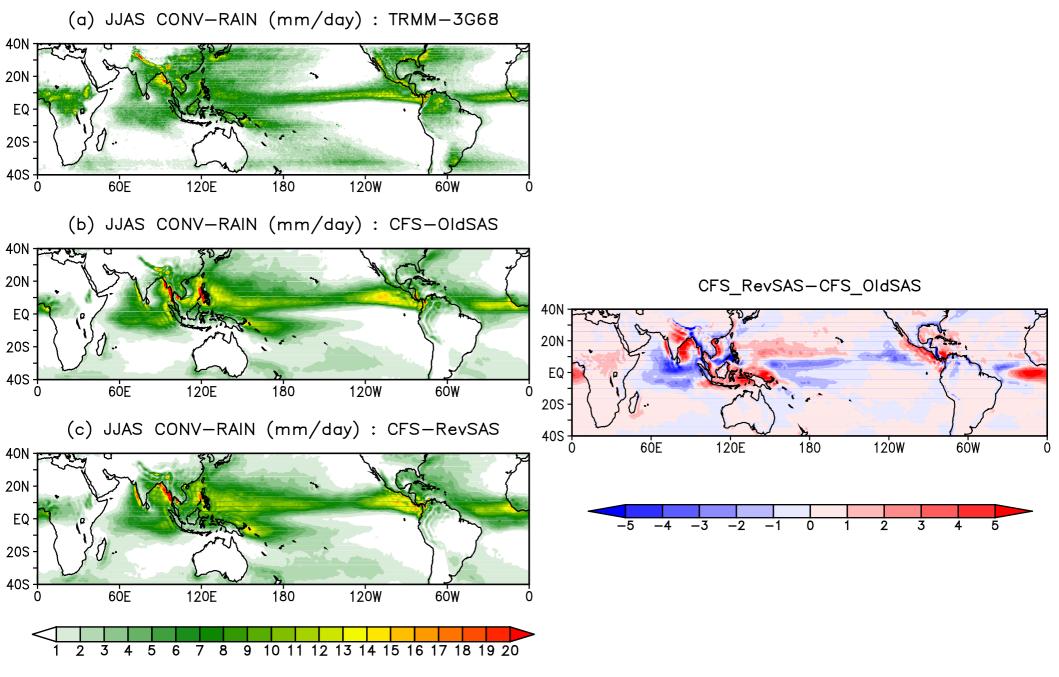
Operationally not feasible (ensemble will be a challenge)

Impact of Revising Subgrid scale convection only RevSAS

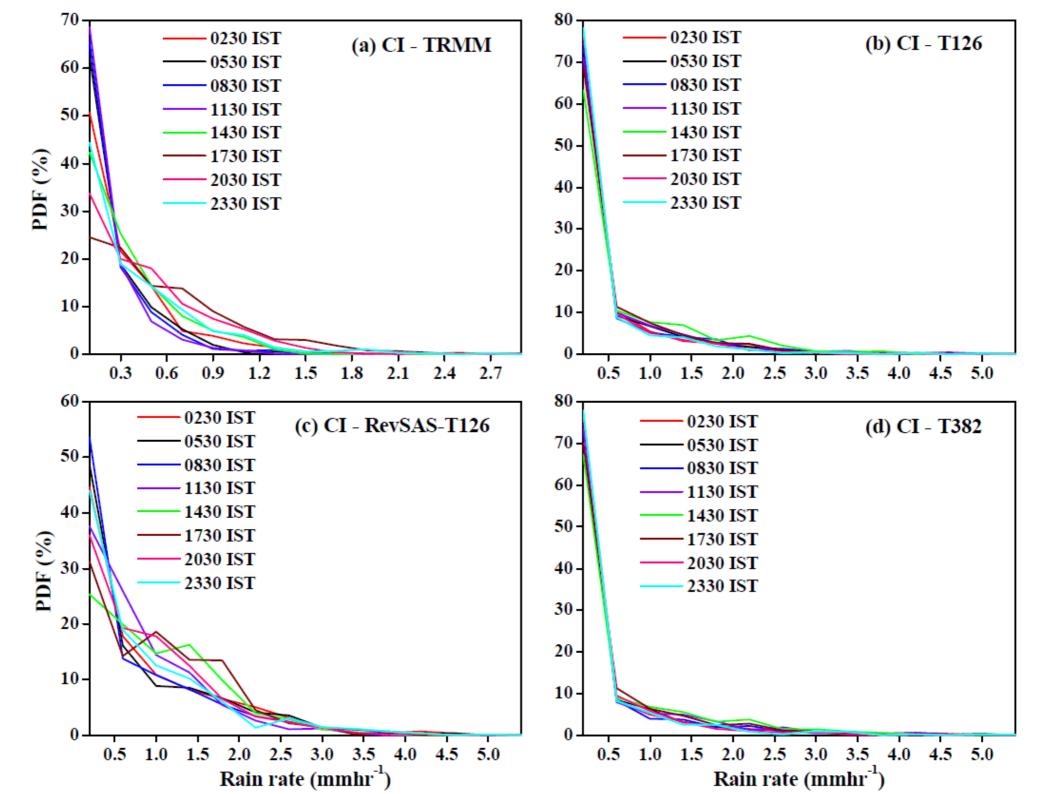
JJAS Mean precip

JJAS precip bias

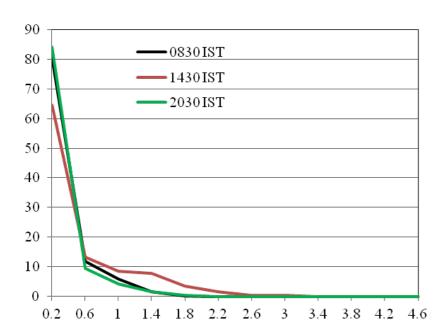




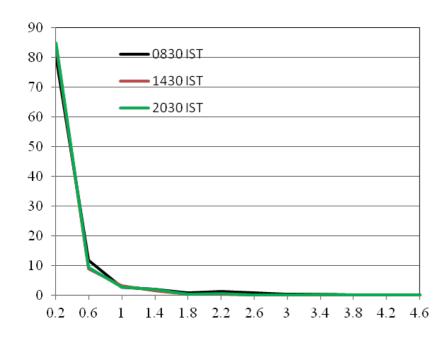
Convective Rain



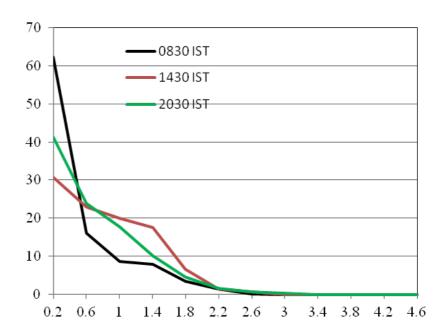
#### Convective-rain OldSAS



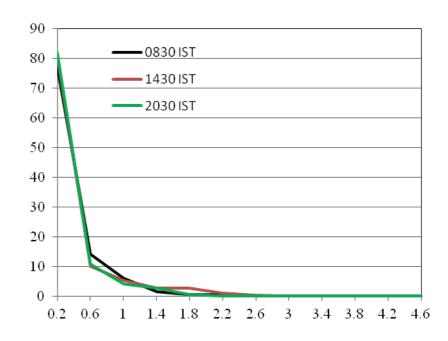
#### Stratiform-rain-OldSAS

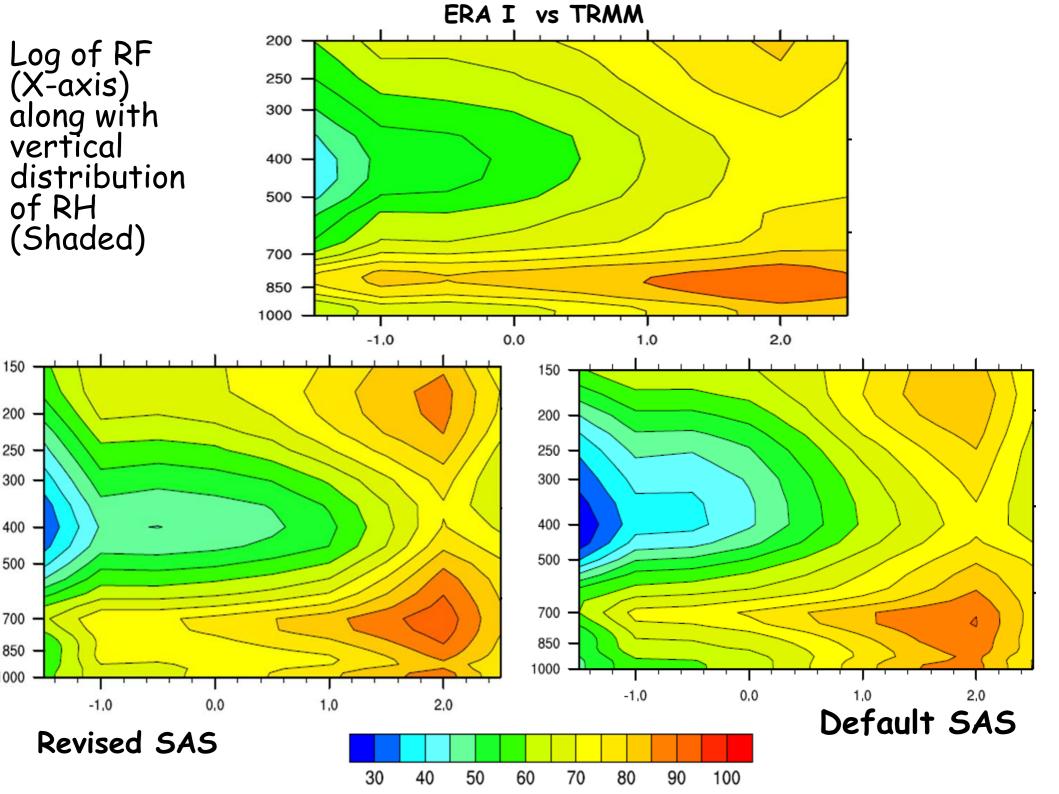


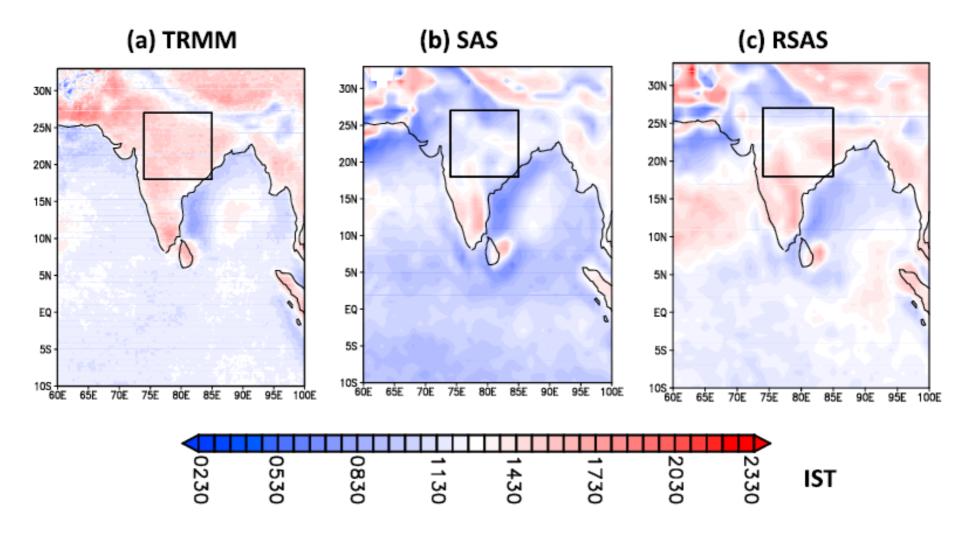
#### Convective-rain-RevSAS



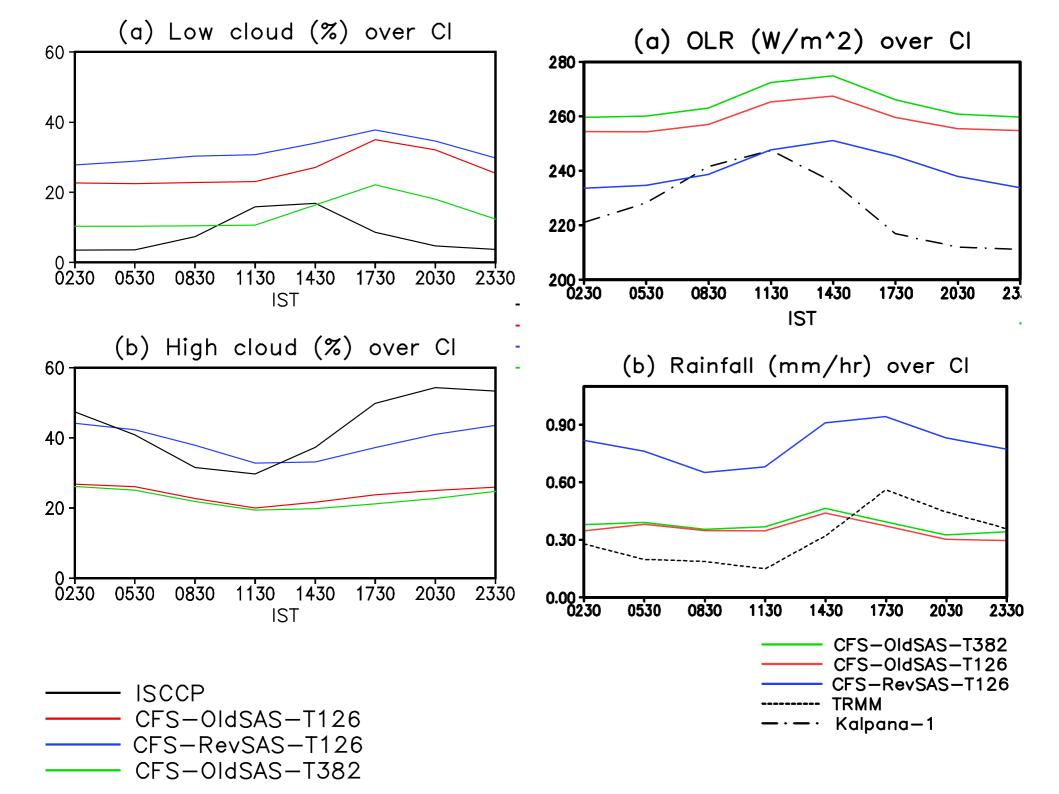
#### Stratiform-rain RevSAS

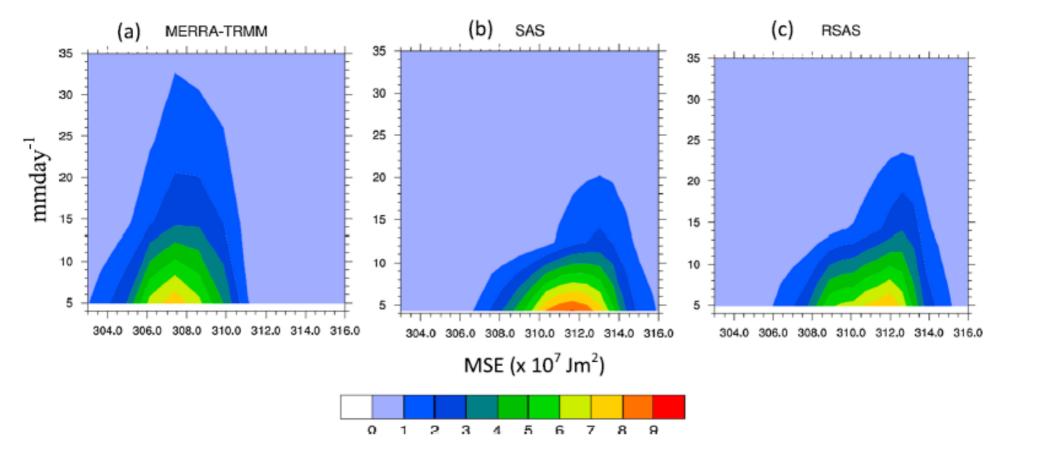






Seasonal (JJAS) mean distribution of diurnal phase (IST (h)) when maximum precipitation occurs for (a) TRMM, CFSv2 with (b) SAS and (c) RSAS scheme. Black box represents central India (CI) ( $18^{\circ}N-27^{\circ}N$ ,  $74^{\circ}E-85^{\circ}E$ ) region.

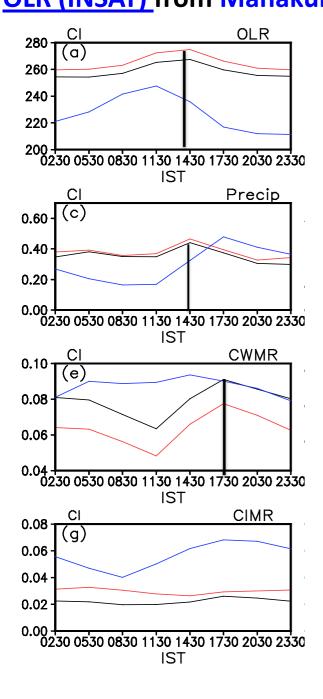


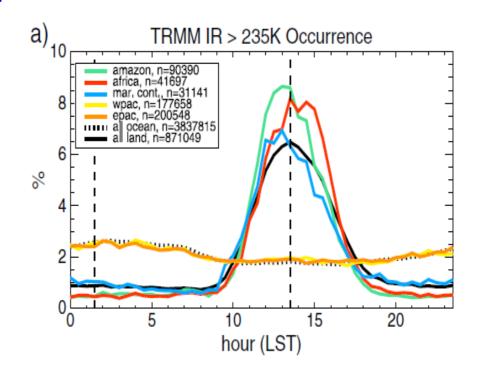


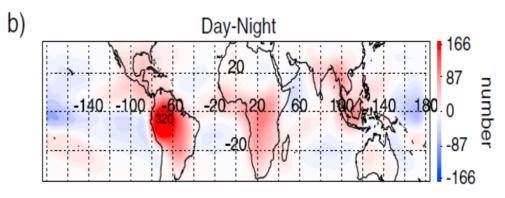
Joint probability distribution function of rainfall (mm d-1), along the y axis, and column integrated (surface to 100 hPa) MSE ( $\times$  10 7Jm-2), along the x axis, over CI region for (a) observation (TRMM and MERRA), CFSv2 with (b) SAS and (c) RSAS scheme during JJAS.

# WHAT NEXT

### CFST382 (RED line), CFST126 (BLACK line), TRMM & MERRA (Dotted BLUE line) **OLR (INSAT)** from Mahakur et al.



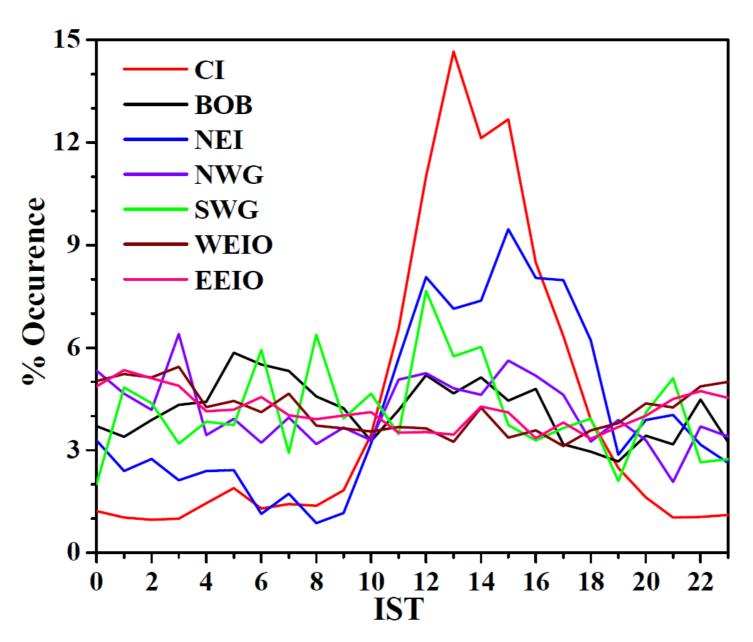




variation of population congestus different different regions. The bloadashed lines indicate the times CloudSat overpass.

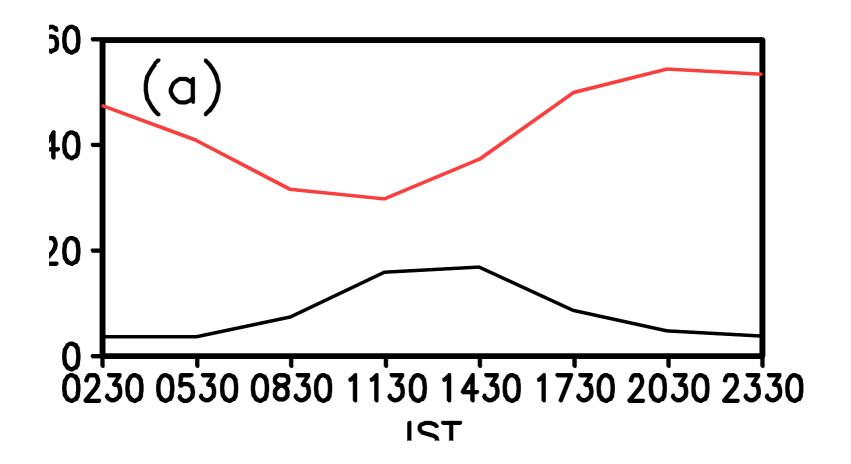
Wall et al., 2013

# Diurnal variation of population of TRMM VIRS congestus for different regions over Indian monsoon region

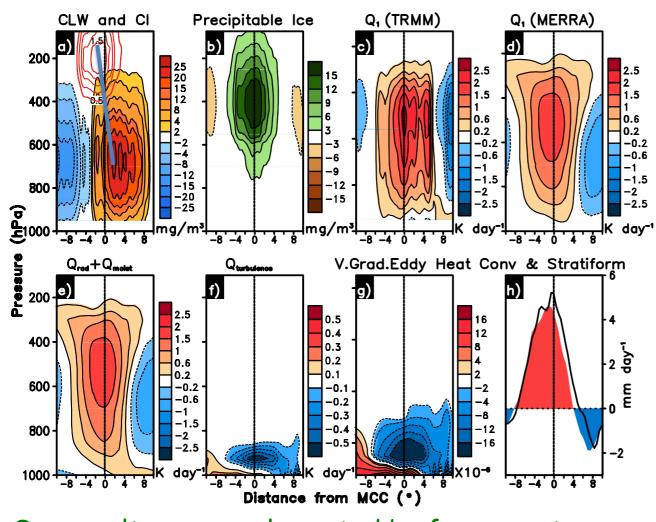


Mahakur et al.

ISCCP estimated low cloud (%) (black line) and high cloud (%) (Red line)



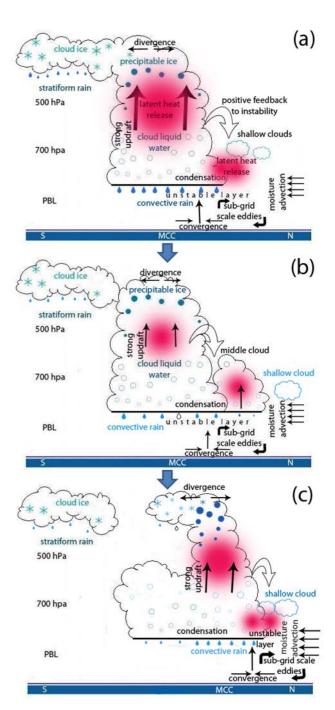
# Hypothesis based on observation for northward propagation BSISO (Abhik et al, 2013)



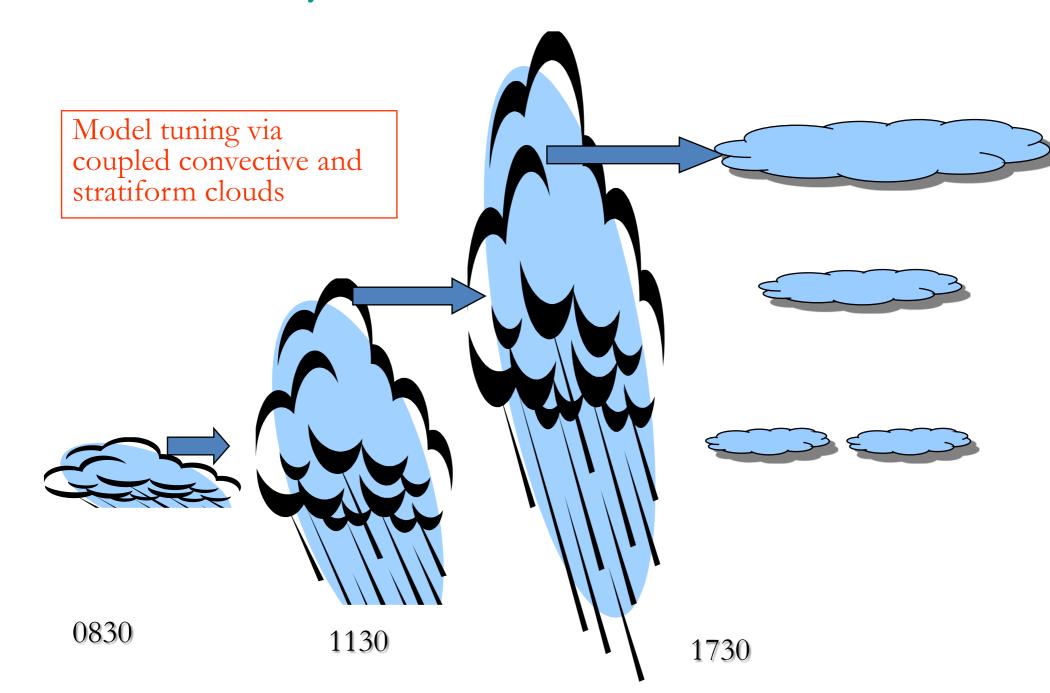
Our results are supplemented by few recent studies e. g.

Preconditioning Deep Convection with Cumulus Congestus by Hohenegger and Steven, 2013

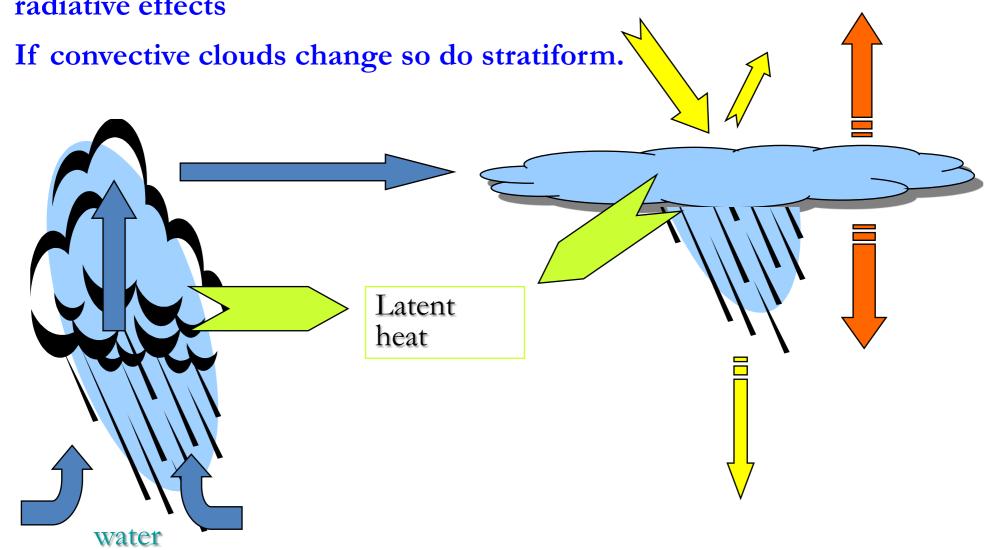
A climatology of tropical congestus using CloudSat by Wall et al. 2013



To simulate better stratiform clouds a spectrum of cumulus clouds is necessary.



Separation into stratiform and convective clouds, detraining water from convective clouds is source for stratiform clouds => radiative effects



# Two-moment bulk stratiform cloud microphysics in the GFDL AM3 GCM: description, evaluation, and sensitivity tests

M. Salzmann<sup>1,\*</sup>, Y. Ming<sup>2</sup>, J.-C. Golaz<sup>2</sup>, P. A. Ginoux<sup>2</sup>, H. Morrison<sup>3</sup>, A. Gettelman<sup>3</sup>, M. Krämer<sup>4</sup>, and L. J. Donner<sup>2</sup>

#### NCEP Initiative

Sun and Han, AGU 2014 "Zhao and Carr microphysics scheme has been implemented into the NCEP Global Forecasting System (GFS) for many years. It predicts total cloud condensate (cloud water or ice). We are testing several sophisticated microphysics schemes from the Weather Research and Forecasting Model (WRF) in the GFS. These schemes have more cloud species and more physically-based parameterized processes."

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VOLUME 59

### Interactions between Cloud Microphysics and Cumulus Convection in a General Circulation Model

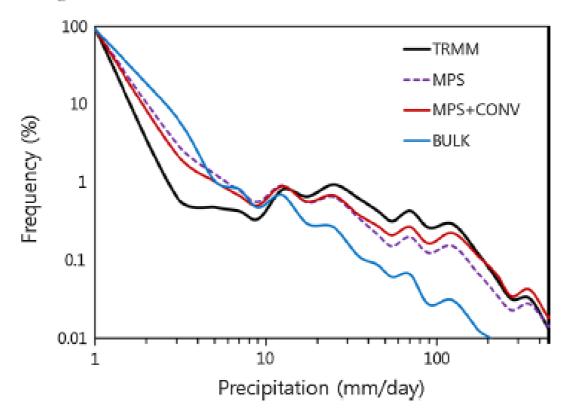
Laura D. Fowler and David A. Randall

Department of Atmospheric Science, Colorado State University, Fort Collins, Colorado

(Manuscript received 1 May 2001, in final form 17 May 2002)

### GCMs with implicit and explicit representation of cloud microphysics for simulation of extreme precipitation frequency

In-Sik Kang · Young-Min Yang · Wei-Kuo Tao



### 649

### A new prognostic bulk microphysics scheme for the IFS

Richard M. Forbes<sup>1</sup>, Adrian M. Tompkins<sup>2</sup> and Agathe Untch<sup>1</sup>

Research Department

<sup>1</sup>ECMWF <sup>2</sup>ICTP, Italy

September 2011

This paper has not been published and should be regarded as an Internal Report from ECMWF.

Permission to quate from it should be abtained from the ECMWF.



European Centre for Medium-Range Weather Forecasts
Europäisches Zentrum für mittelfristige Wettervorhersage
Centre européen pour les prévisions météorologiques à moyen terme

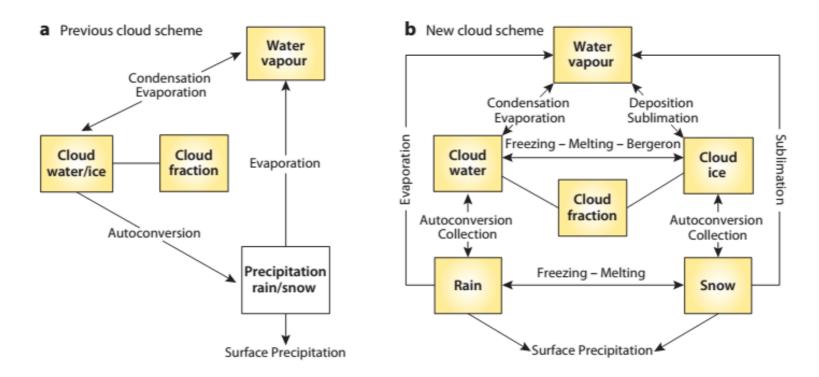


Figure 1: Schematic of the IFS cloud scheme: (a) the Tiedtke scheme with three moisture related prognostic variables operational from 1995 to 2010 (before IFS Cy36r4) and (b) the new cloud scheme with six moisture related prognostic variables (Cy36r4 onwards). Yellow boxes indicate prognostic variables.

#### 2.2 Numerical framework

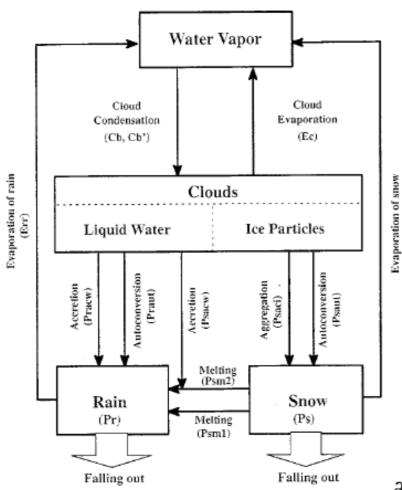
The new scheme is a multi-species prognostic microphysics scheme, with m = 5 prognostic equations for water vapour, cloud liquid water, rain, cloud ice and snow. The equation governing each prognostic cloud variable within the cloud scheme is

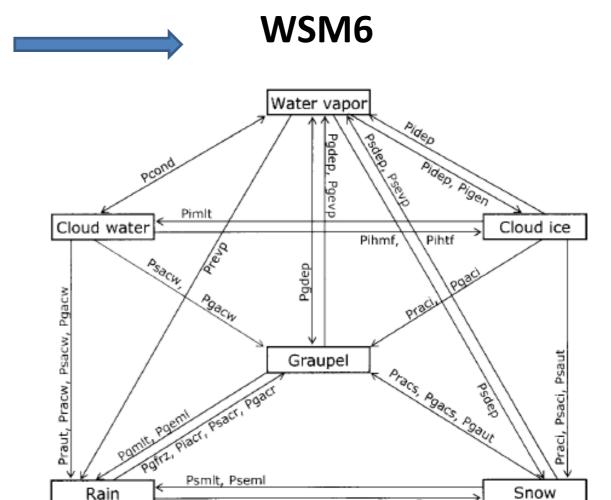
$$\frac{\partial q_x}{\partial t} = S_x + \frac{1}{\rho} \frac{\partial}{\partial z} (\rho V_x q_x), \qquad (1)$$

where  $q_x$  is the specific water content for category x (x = 1 for cloud liquid droplets, x = 2 for rain, and so on),  $S_x$  is the net source or sink of  $q_x$  through microphysical processes, and the last term represents the sedimentation of  $q_x$  with fall speed  $V_x$ .

#### Zhao & Carr 1997

#### **Default CFS Microphysics**





**Tendencies** 

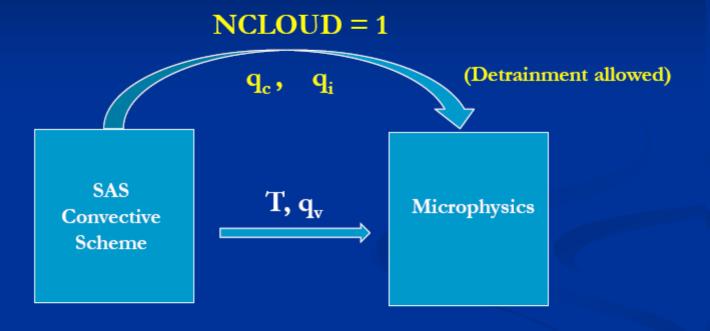
Piacr, Psacr

$$\frac{\partial}{\partial t}(n) = ADV(n) + TURB(n) + SEDIM(n) + SOURCE(n)$$

where n =  $[n_r, n_i, n_s, n_{c/w}, n_g, n_v]$  represents the concentration of rain, ice crystals, snow, graupel, cloud water, water vap.

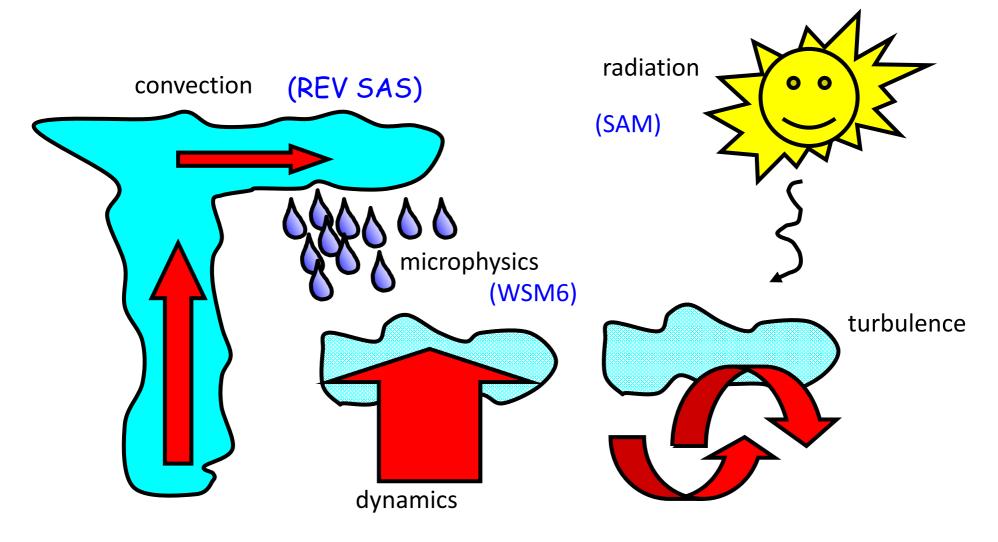
Rain

# The Convective and Microphysical Scheme Communication





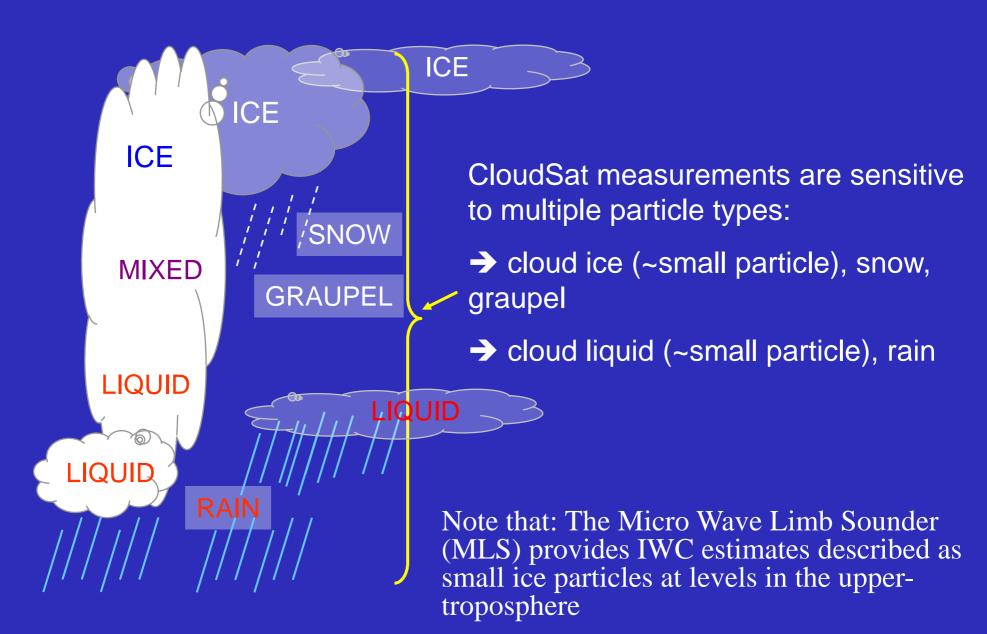
# Revised Cloud-Convective-Radiation in CFSv2 T126



Clouds are the result of complex interactions between a large number of processes SAM: System of Atmospheric Model

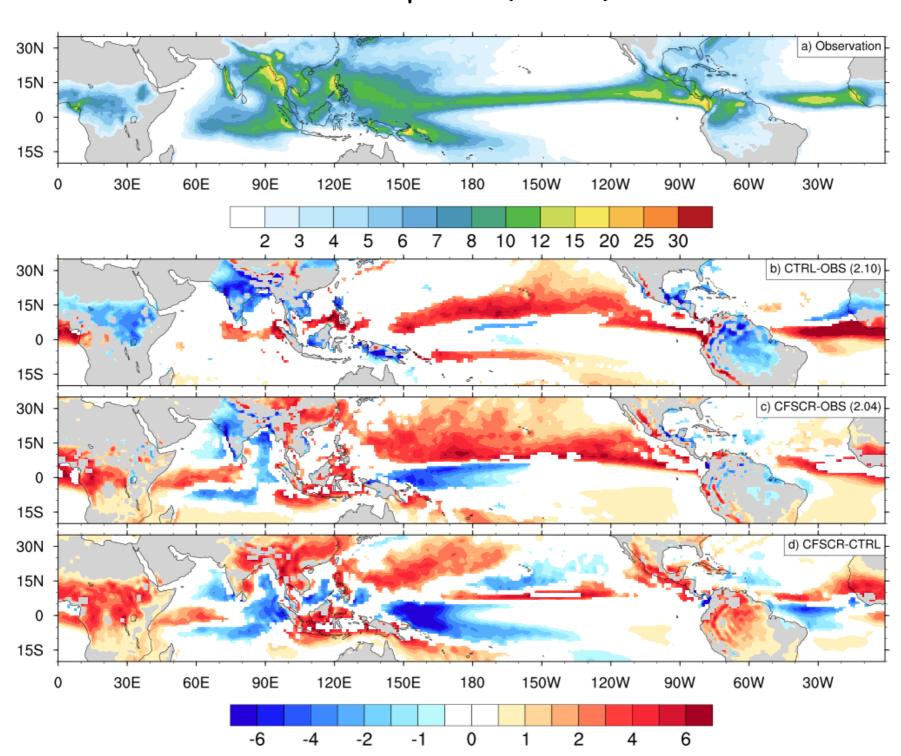
### CloudSat IWC/LWC Retrieval



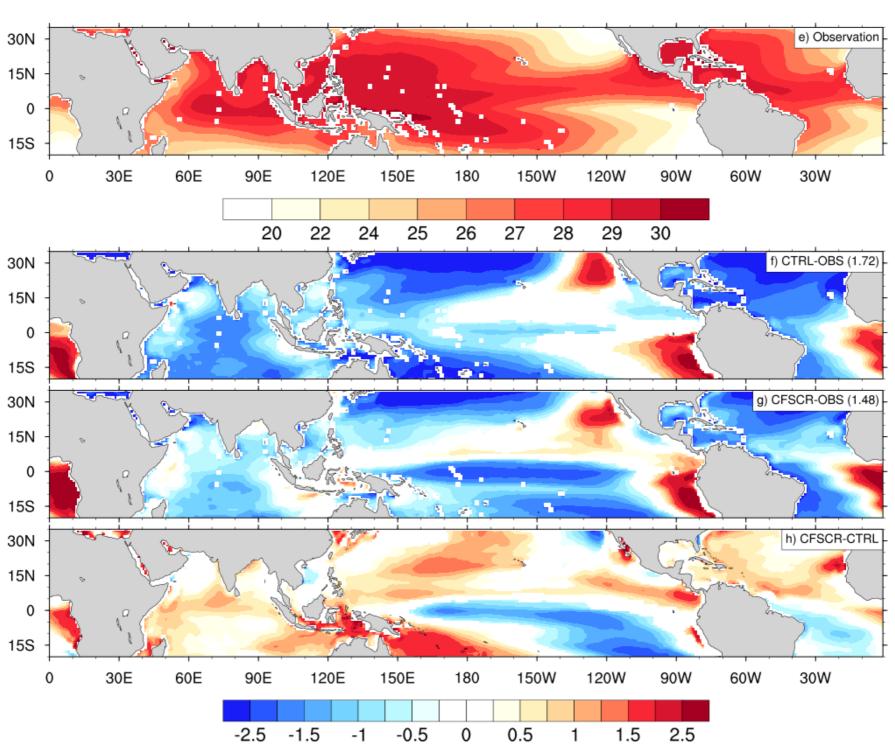


Slide Courtesy: Frank Li, JPL

### Precip bias (JJAS)

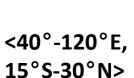


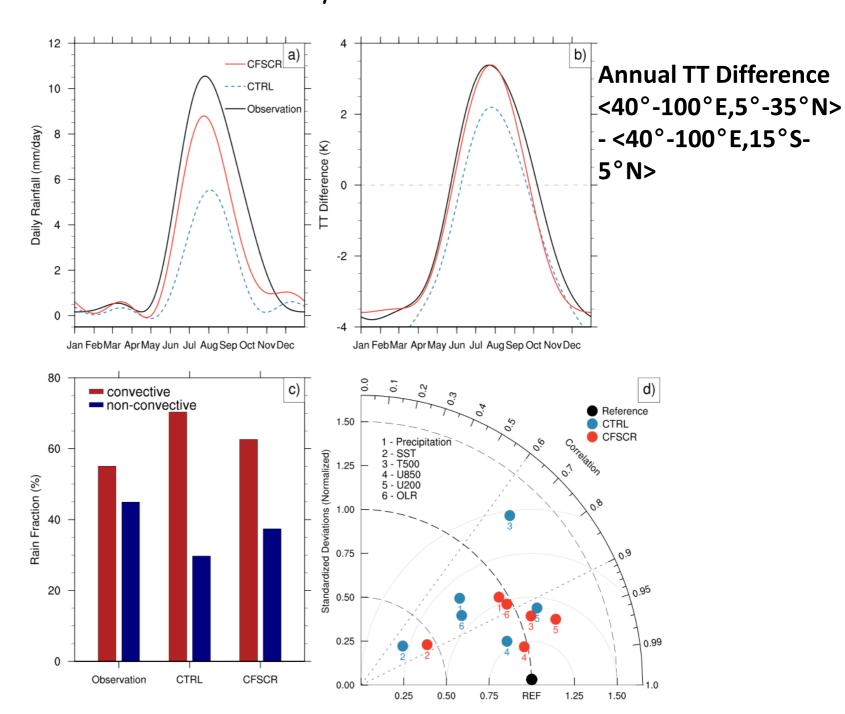
### SST bias

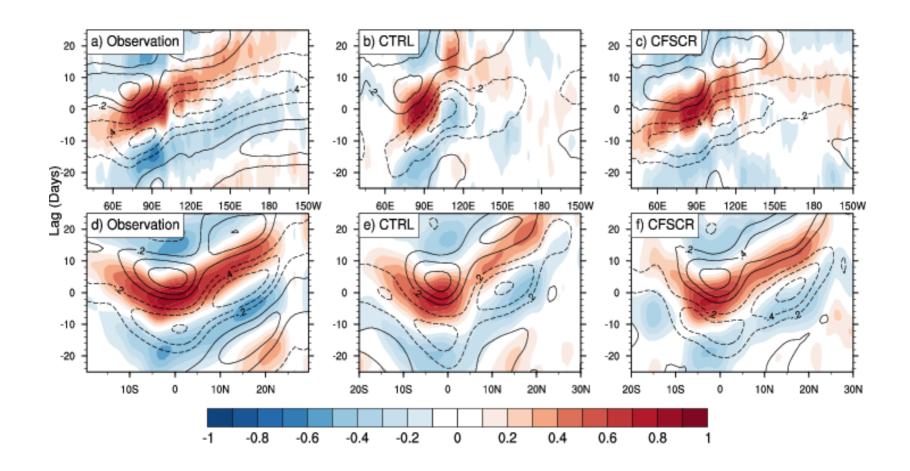


Revised convection, modified microphysics and radiation is able to improve the mean state and Intraseasonal variability of CFSv2T126

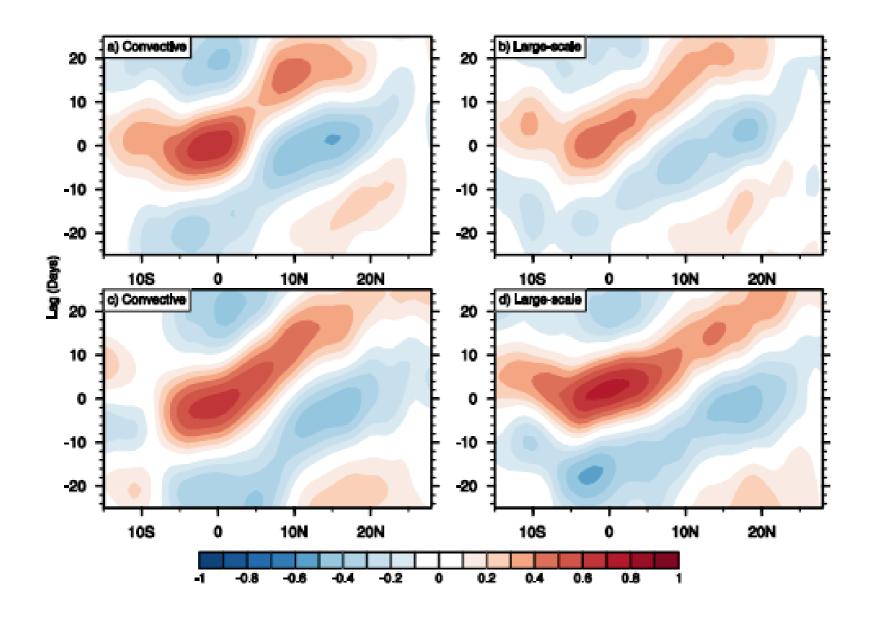
Annual Rainfall Cycle <73°-85°E,15°-25°N>



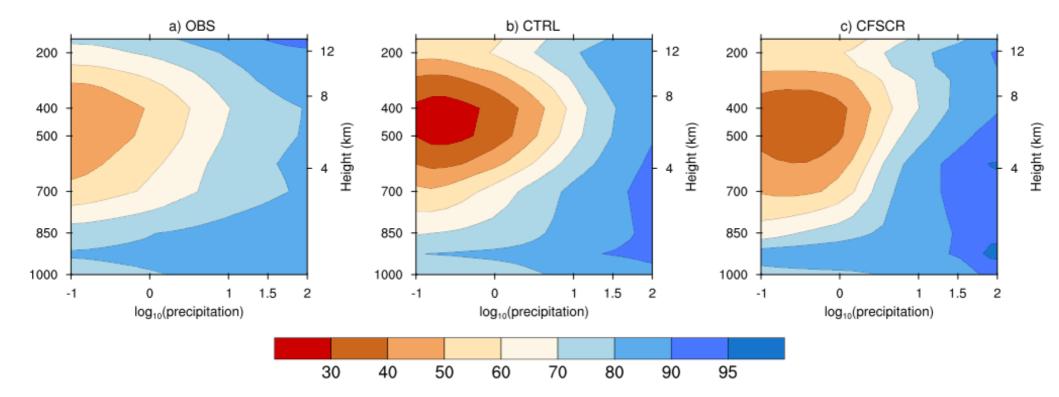




Longitude (Latitude) vs lag correlation of 20–100-day filtered precipitation (shaded) and  $U_{850}$  (contour) with base 20–100-day filtered precipitation time series over EEIO (10°S–5°N, 75°–100°E).



Longitude (Latitude) vs lag correlation of 20–100-day filtered convective and large-scale precipitation from CTRL and CFSCR



Composite profile of relative humidity as a function of rain rate over ISM domain (40°-120° E, 15° S-30° N) during all JJAS seasons from (a) observation (ERA-Interim vs TRMM), (b) CTRL and (c) CFSCR. The rain rate at the x-axis is plotted in  $\log_{10}$  scale.

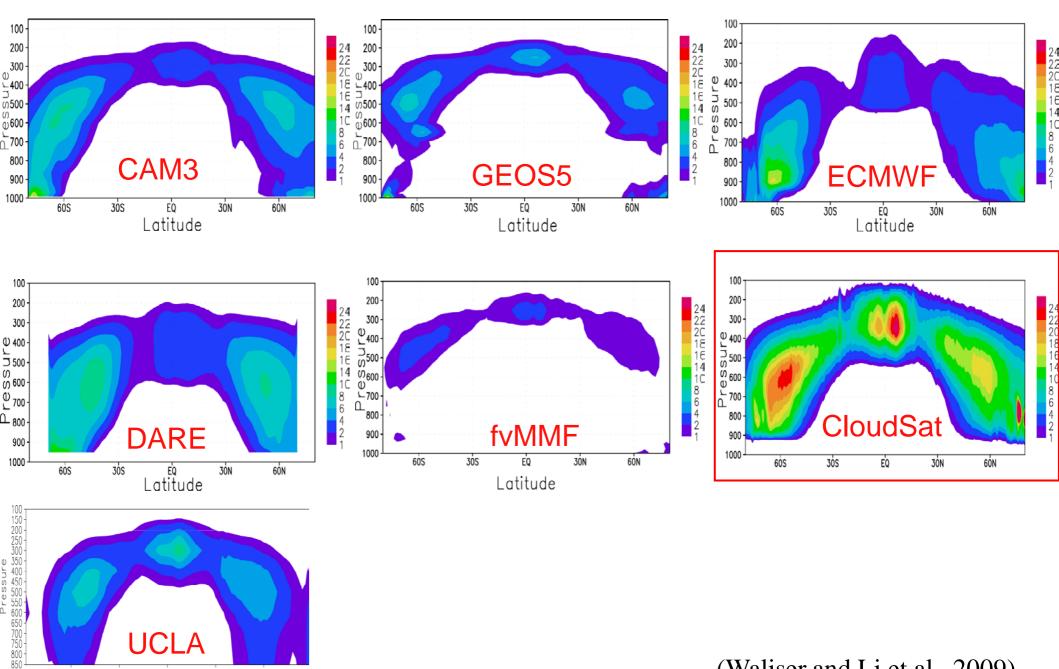
#### GCM CLOUD ICE WATER CONTENT (IWC) ANNUAL MEAN VALUES

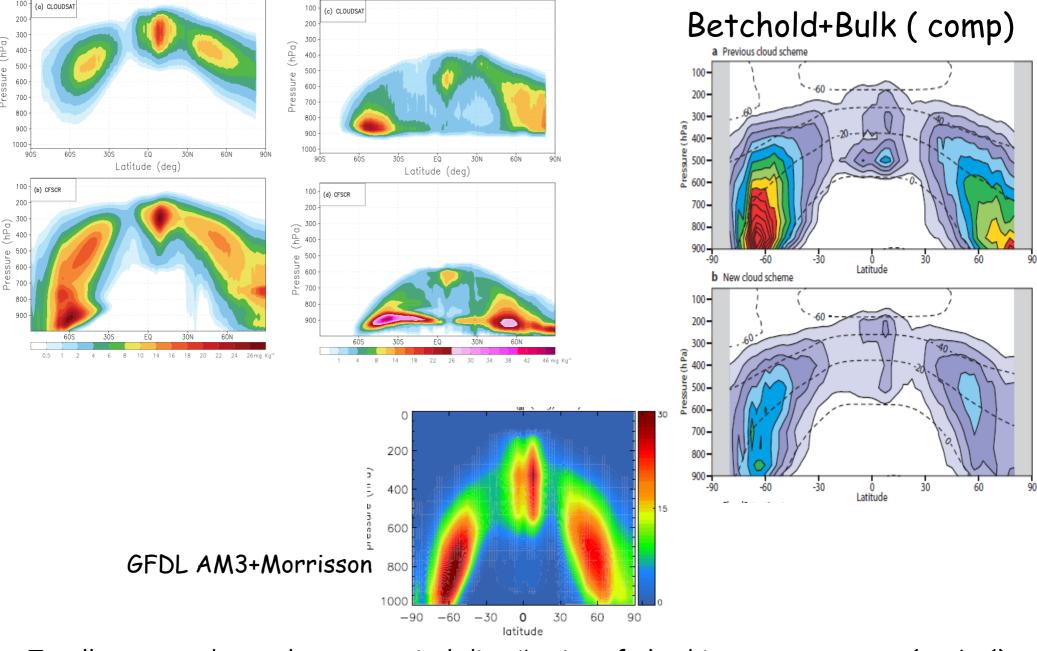
**UCLA** 

EQ

30N

305





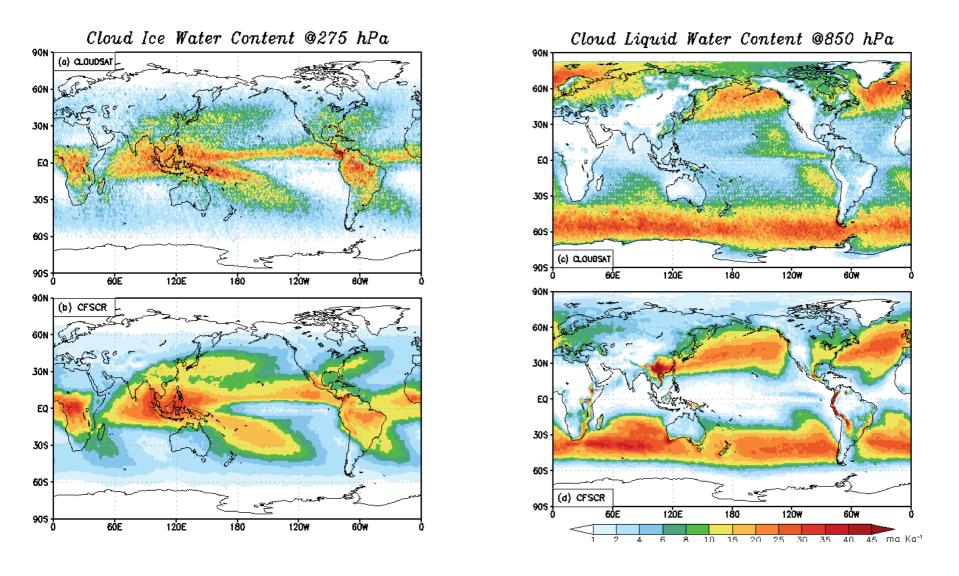
Cloud Liquid Water Content

Cloud Ice Water Content

ECMWF IFS cloud ice

Zonally averaged annual mean vertical distribution of cloud ice water content (mg kg<sup>-1</sup>) obtained from (a) CFSCR; and cloud liquid water content (mg kg<sup>-1</sup>) from (b) CFSCR model.

CFSCR: Modified CFSv2 with revised Cloud Microphysics, Convection and radiation



Annual mean isobaric distribution of cloud ice water content (mg kg<sup>-1</sup>) obtained from (a) CloudSat 2B-CWC-RO, (b) CFSCR (at 271 hPa model level); and cloud liquid water content (mg kg<sup>-1</sup>) from (c) CloudSat, (d) CFSCR (858 hPa).

Cloud Ice Water Content Cloud Liquid Water Content 100 (a) CLOUDSAT 200 100 (a) CLOUDSAT 200 300 Pressure (hPd) 300 400 Pressure (hPd) 400 500 500 600 600 700 700 800 800 900 900 1000 1000 305 EQ 50% 30N 905 60N 90N 305 EQ. 605 30N 60N 905 90N Latitude (deg) Latitude (deg) 100 -1.00 blas=2.1, nmse=6.7 blas=0.1, rmse=8.0 (b) CFSCR (d) CFSCR 200 200 500 300 -Pressure (hPa) Pressure (hPa) 400 480.5 500 500 600 500 700 700 800 800 -900 900 -1000

1000 -905

505

305

EQ

10.

30N

CON

905

60%

305

EQ

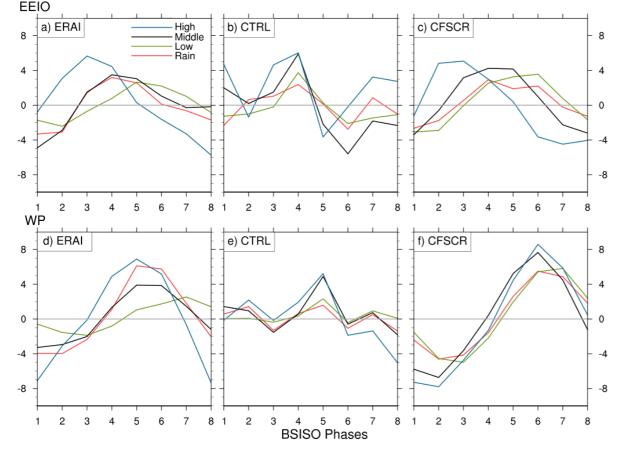
30N

30

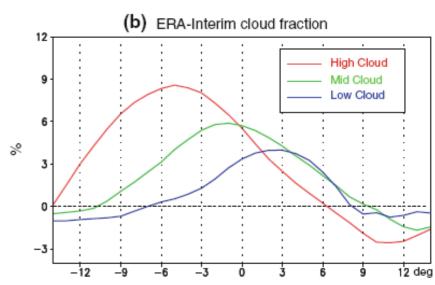
60N

18

90N

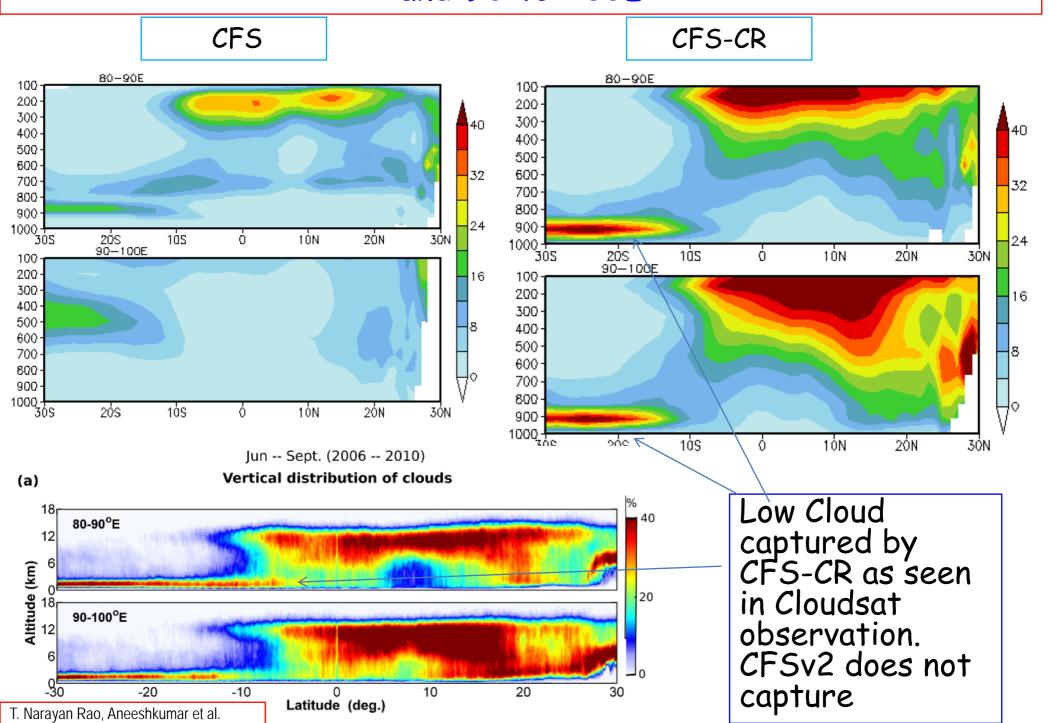


Evolution of anomalous low, middle and high cloud fractions (%, left axis) and rainfall anomalies (mm day-1, right axis) associated with BSISO1 convection over EEIO (top panels) and WP (bottom panels) for (a-b) observation, (c-d) CTRL and (e-f) CFSCR.



Jiang et al. 2011

## Vertical distribution of clouds (JJAS) averaged over 80-90E and 90 to 100E



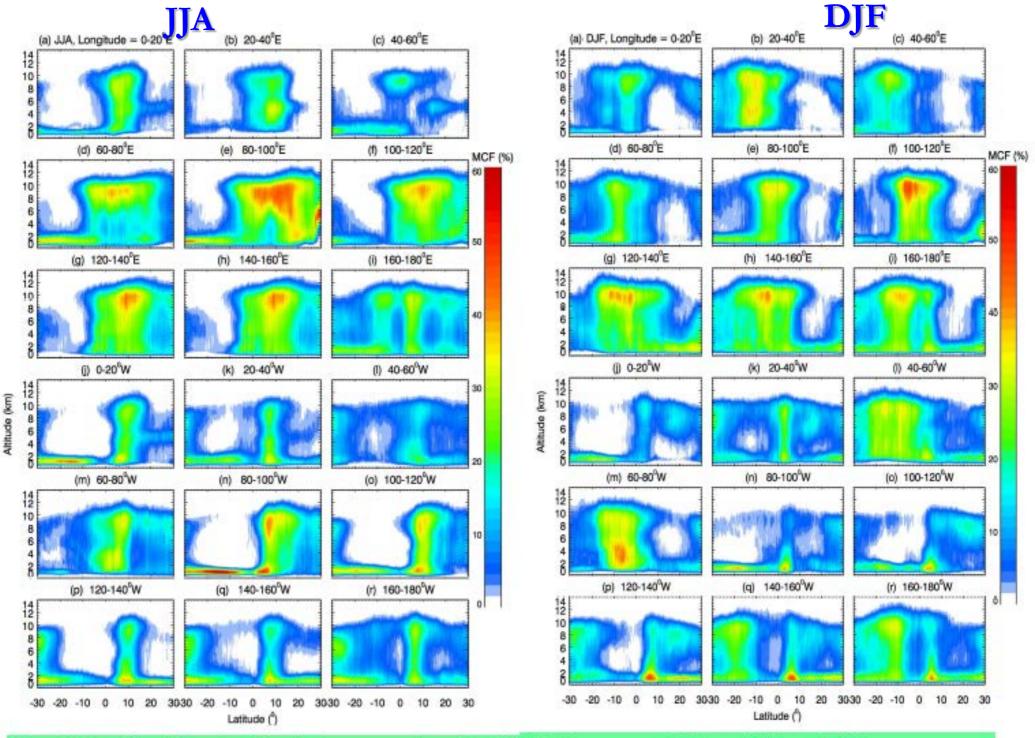
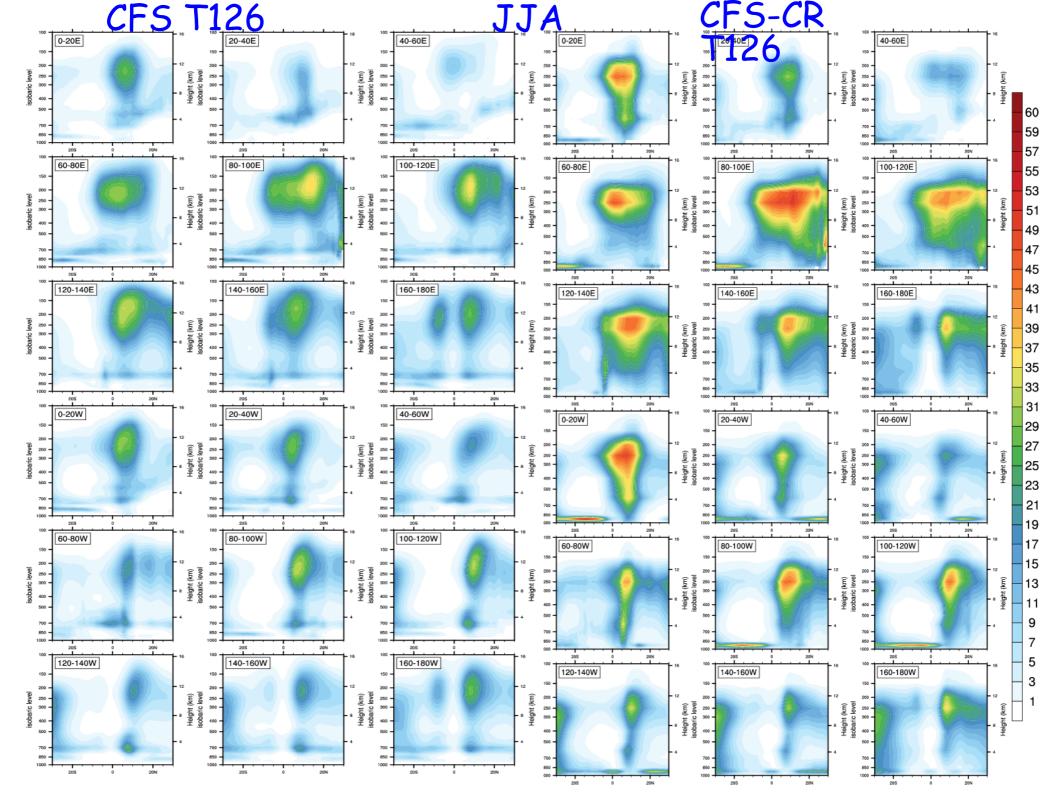
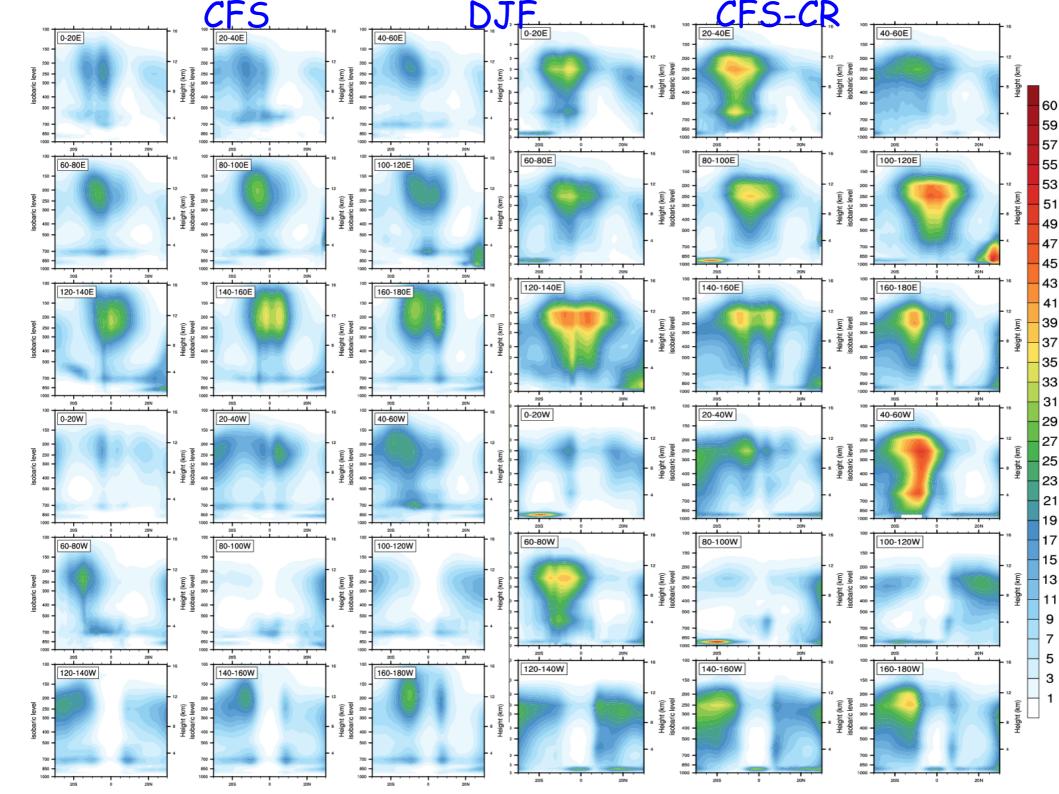
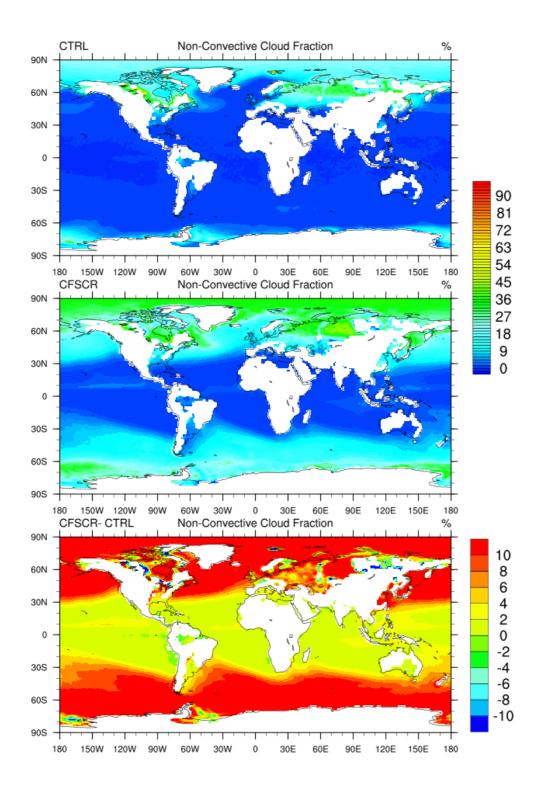


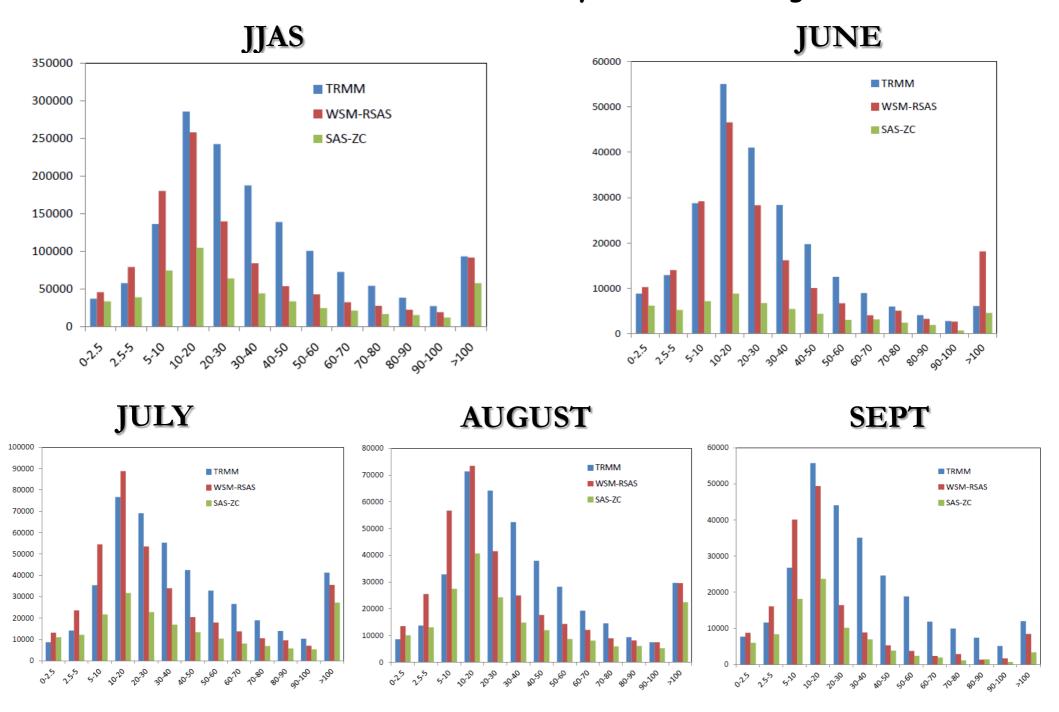
Fig. 1 Mean Cloud Fraction (longitude cross-section of 0-20°) during summer (JJA) and Winter (DJF) seasons represents the characteristic features associated with Ascending and Descending limbs of Walker and Hadley Cells.



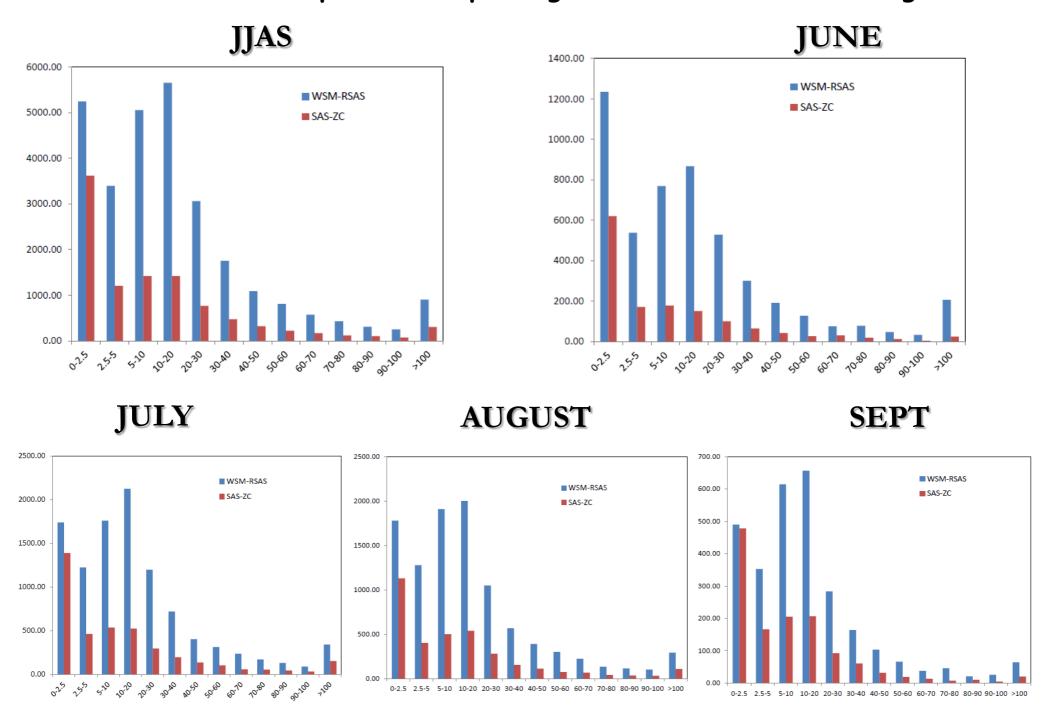


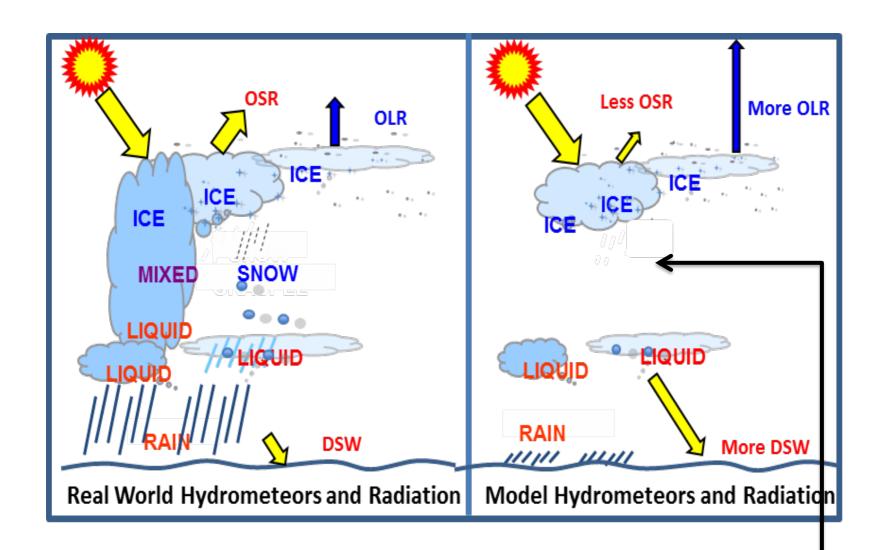


#### Amount of total rainfall contributed by different categories



#### Total cloud water path corresponding to different rainfall categories





Bridging the Gap in CFSv2 using modified Microphysics: WSM6

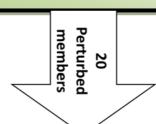
### SCHEMATIC OF GEFS (SL) T574 L64 RUNNING AT IITM

**GDAS** 

EnKF – GSI Hybrid Data Assimilation System Ensemble Initialization method: Ensemble Kalman Filter (EnKF) scheme.

The 6-hr forecasts from the previous cycle





Tropical Storm Relocation (if storm is present)

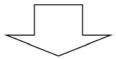
Centering of the perturbations on the ensemble control analysis (Distributes the spread around analysis instead of Ensemble Mean



FORECAST: 21 members runs for 192 hrs (8 days)

GFS Semi-Lagrangian T574 (approx 35 km at equator); L64 vertical resolution

The stochastic total tendency perturbation (STTP) to enhance model uncertainty



POST PROCESSING: 192 hr (8 days) forecast

Resolution:  $0.5^{\circ} \times 0.5^{\circ} / 1^{\circ} \times 10^{\circ} / 2.5^{\circ} \times 2.5^{\circ}$ 

Products: 21 members, Ensemble mean, Ensemble Spread

### SCHEMATIC OF GFS (SL) T1534 L64 RUNNING AT IITM

**GDAS** 

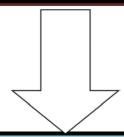
EnKF - GSI Hybrid Data Assimilation System

Analysis

FORECAST: GFS Semi-Lagrangian T1534 (approx 12 km at equator)

L64 vertical resolution

Runs for 240 hrs (10 days)



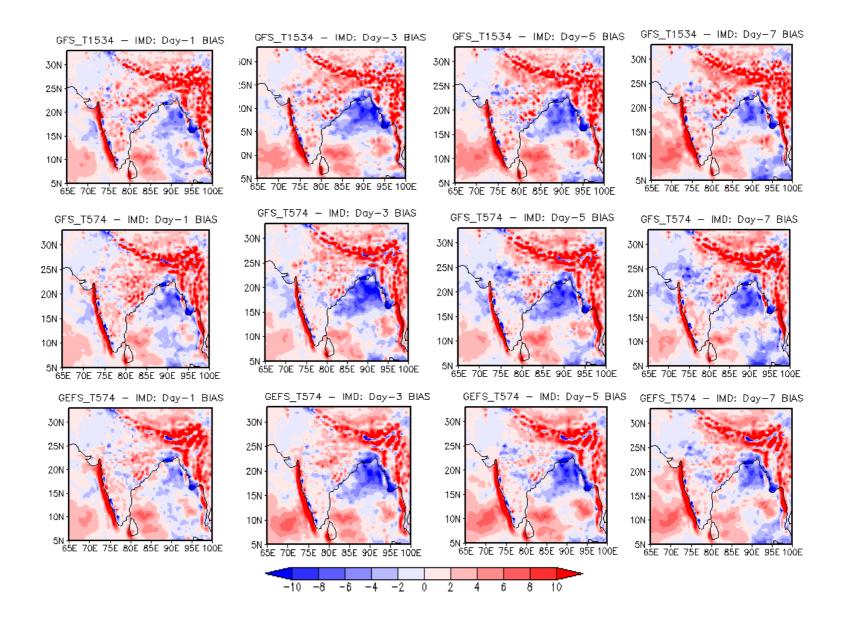
#### POST PROCESSING

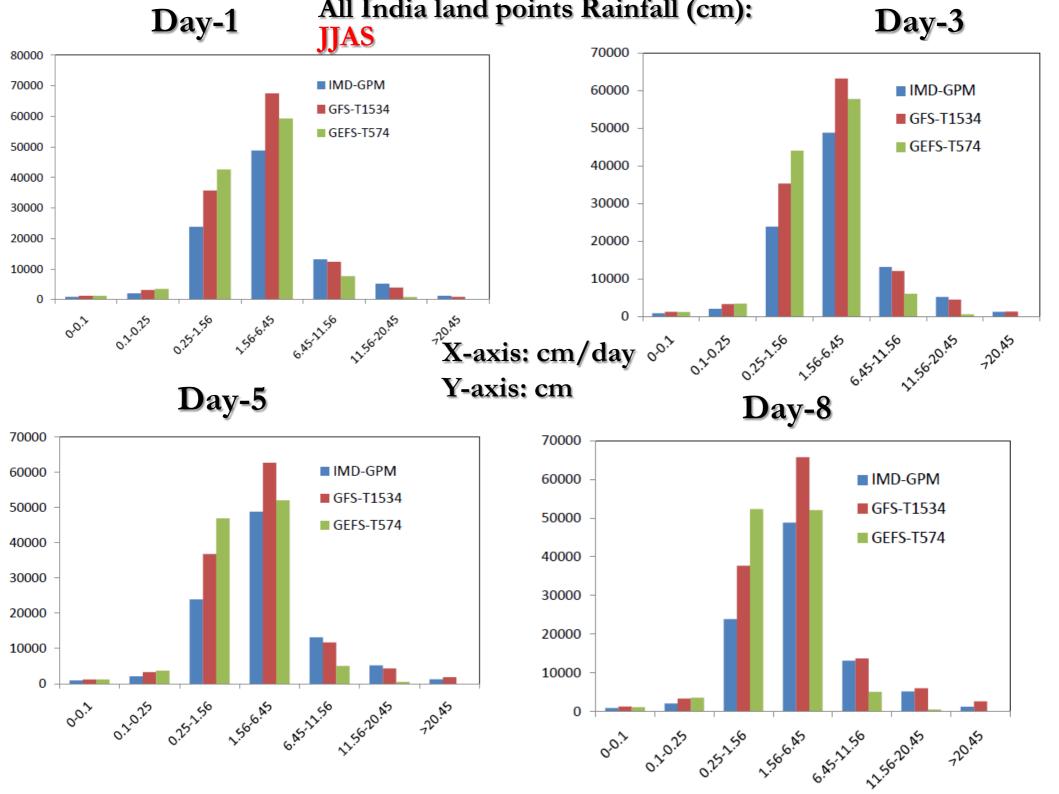
120hr (10 days) forecast

Resolution: Regular grid and Gaussian grid at different resolutions

 $0.125^{\circ} \times 0.125^{\circ} / 0.25^{\circ} \times 0.25^{\circ} / 0.5^{\circ} \times 0.5^{\circ} / 1^{\circ} \times 1^{\circ}$ 

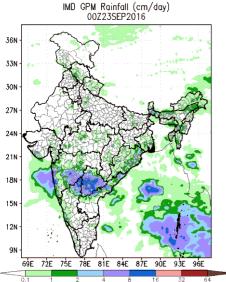
### JJAS 2016 rainfall BIAS (mm/day)





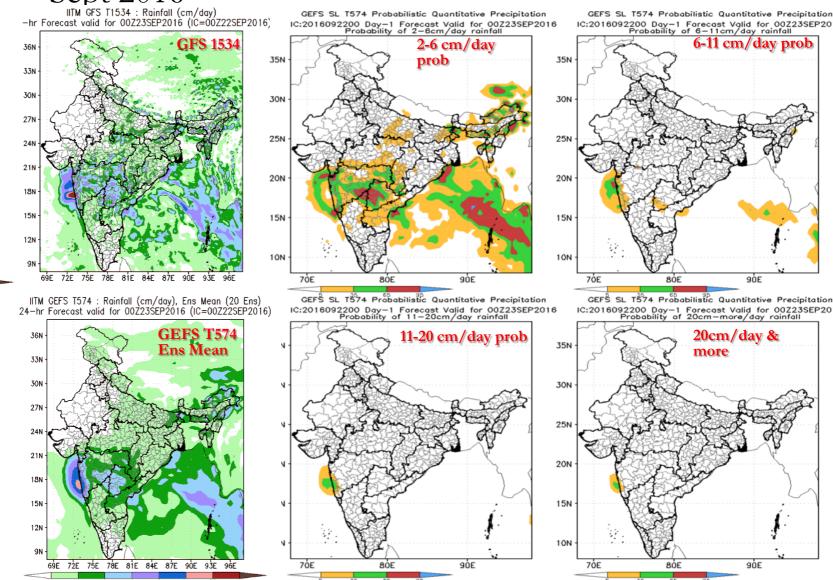
### Recent very heavy rainfall over west coast predicted by GEFS T574 (30 km) with the probability

#### **OBSERVATION**



As per the IMD records, the 24 hr accumulated station rain on 23 Sep 2016 over
Colaba
(18.9067 N/ 72.8147 E) 45mm
SantaCruz
(19.0823 N/72.8407 E) 40 mm
Harnai, Goa
(17.8154 N/73.0981 E): 136.7 mm

### Forecast Valid for 23 Sept 2016 IC:22 Sept 2016



### Summary

Improvement of GCM through better representation of cloud processes

Mean as well as Intraseasonal variability improved

Efficient in capturing the low clouds

GFS T1534 shows promise in capturing heavy precip

GEFS T574 efficient in providing probability of heavy rain

# Thank You!



